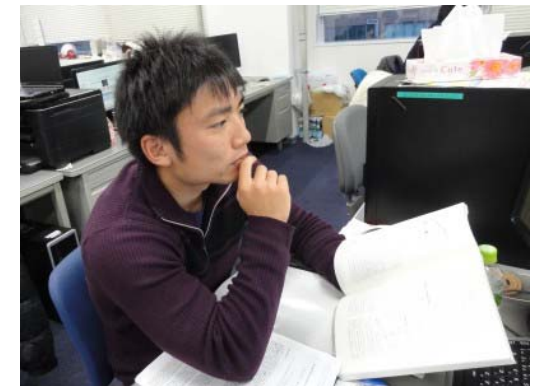


Micromechanical Aspects of Natural Formation Process of Geological Grains

Takashi Matsushima
Kan Sato
University of Tsukuba



Introduction

Geomaterials are natural materials formed by **long-time geological process**.

Their **physical and mechanical properties** differ in different places and change in time as well.

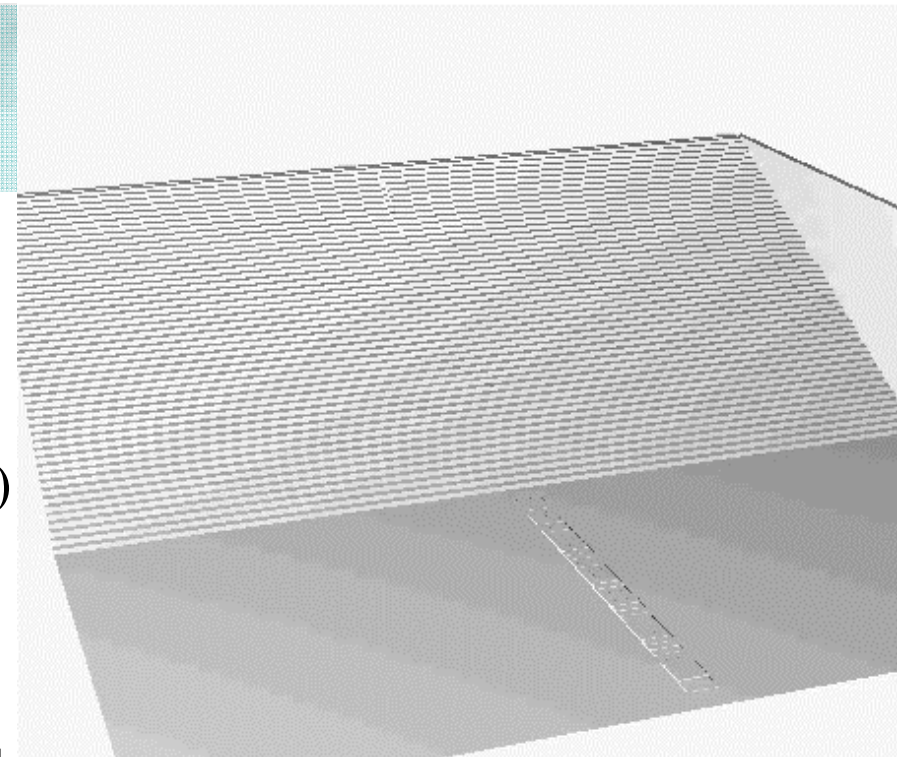


A big challenge in geomechanics is to develop a **simulation tool** for quantitatively evaluating such properties.

Erosion and Sedimentation

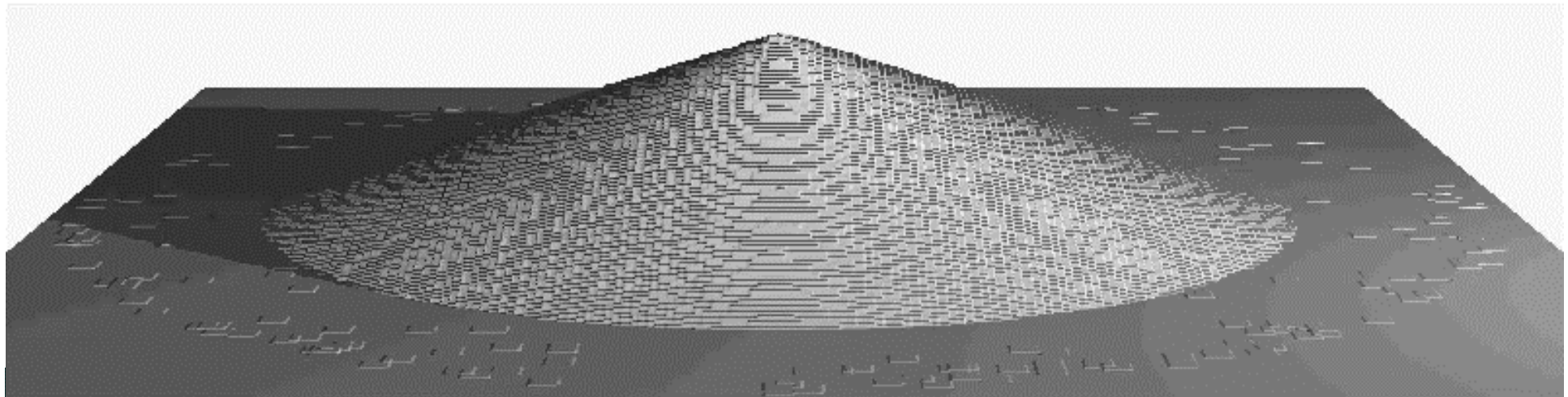
Geological process simulation
by **depth-integrated particle method**
considering **grain segregation**
(Matsushima, JGS annual meeting, 2014,2015)

The erosion/sedimentation model is
based on the experimental observation,
but evolution of GSD is not considered.

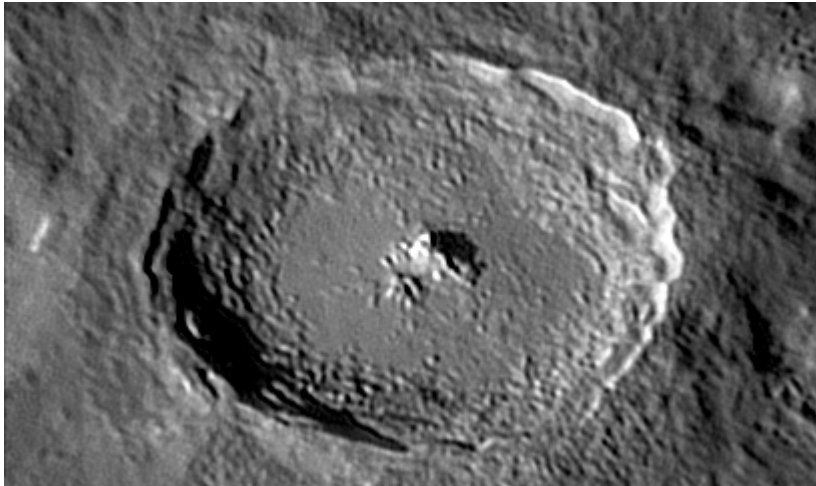


↑ Formation of alluvial fan

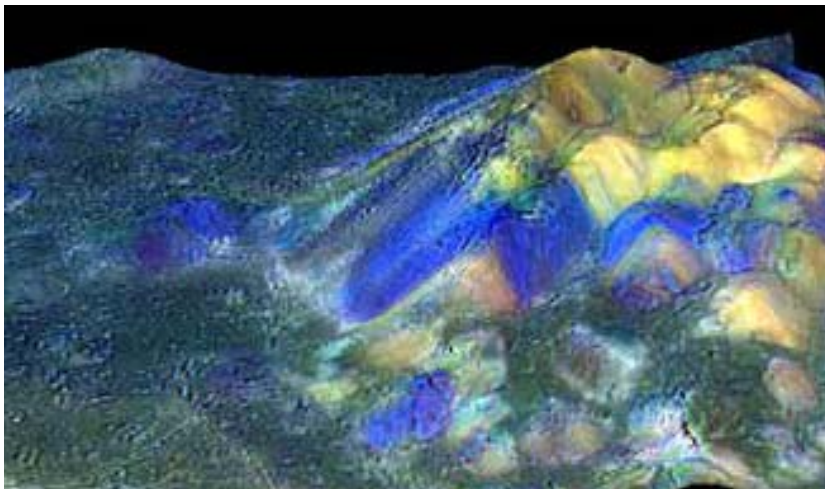
↓ Formation of erosion valley



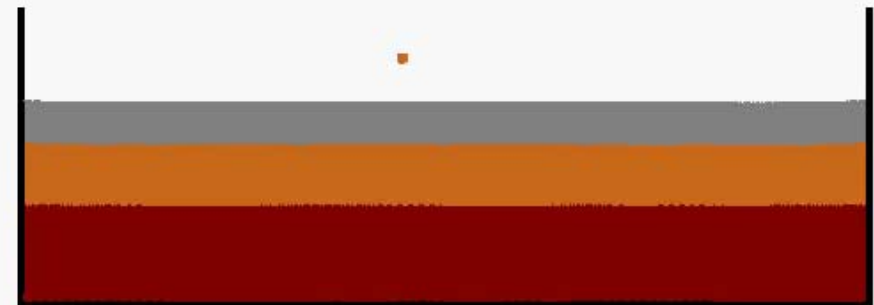
Impact cratering



Tycho Crater (D=85km)



Ohtake et al.: Nature, 461,10,2009



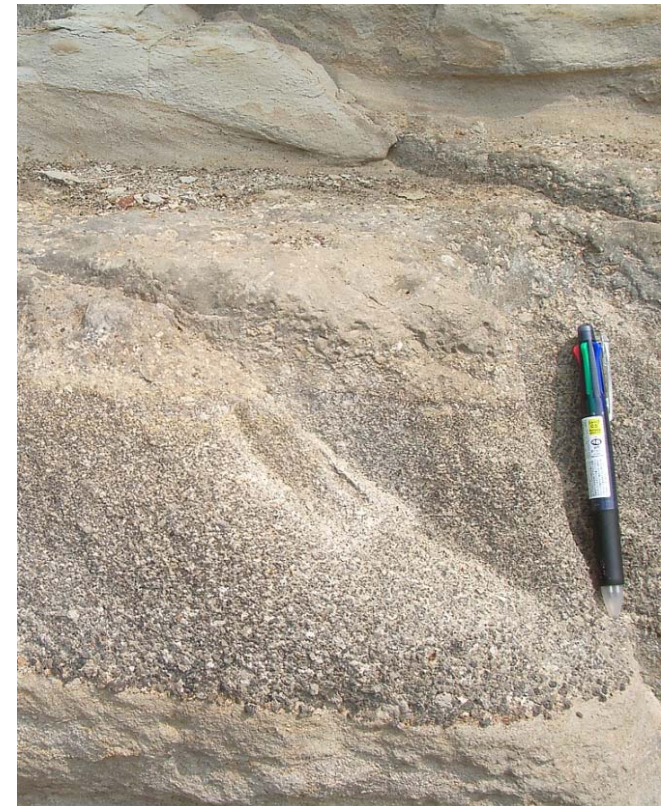
Formation of central peak by
SPH (*unpublished*)
CE: Drucker-Prager EP model

Importance of grain scale processes

On solid planetary surfaces, geomaterials are composed of a lot of **solid geological particles** of various sizes and shapes.

Such particles were formed and evolved either by **crushing**, **agglomeration** or **solidification** process under a certain natural environment.

These processes greatly affect the **bulk mechanical properties** of the layers



grain segregation
in geological layers

Grant-in-Aid for Scientific Research (B) 2014-16

“Micromechanics of compression, shear and solidification of geo-materials under high pressure”

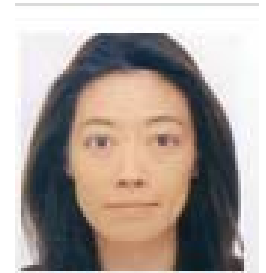
Takashi Matsushima (PI)

Takahiro Hatano (University of Tokyo)

Keiko Watanabe (University of Ritsumeikan)

Masuhiko Beppu (National Defense Academy)

Hiroko Kitajima (Texas A&M University)



Various experiments with **common geo-materials**

Materials used



Gifu sand
mountain sand
90.2% SiO_2
 $D_{mean} = 2.38\text{mm}$
angular



Kashima sand
river sand
96.3% SiO_2
 $D_{mean} = 2.08\text{mm}$
round

Experimental program

today's talk

[1] **Single grain crushing** test

Sato

[2] **One-dimensional compression (ODC)** test

Sato

[3] **Rotary shear (RS)** test

Kitajima, Sato

[4] High-speed projectile impact test

Watanabe

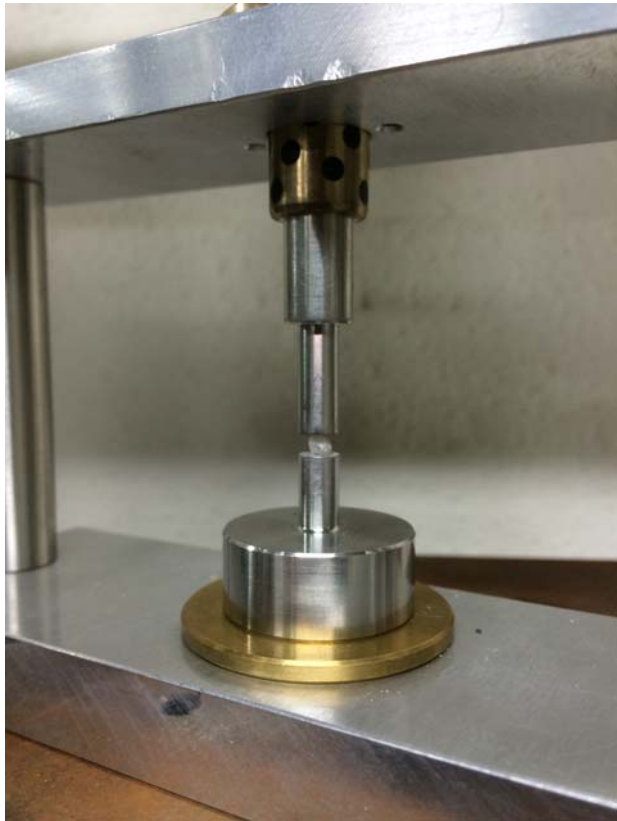
[5] Explosion test

Beppu

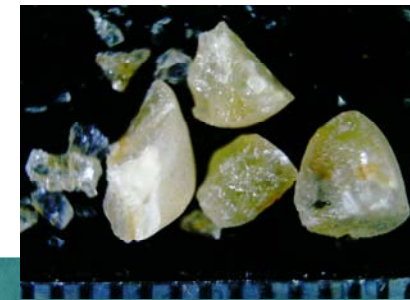
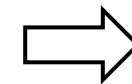
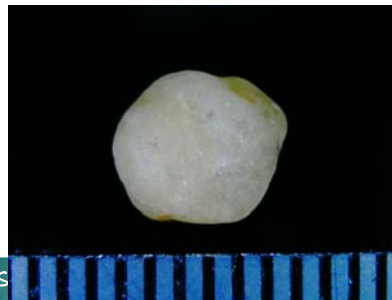
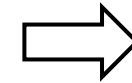
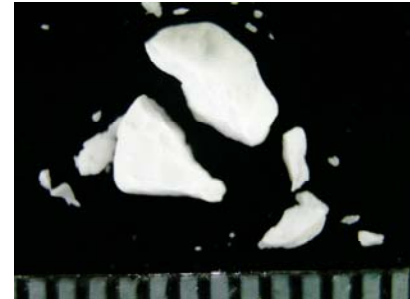
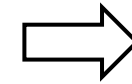
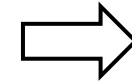


[1] Single grain crushing test

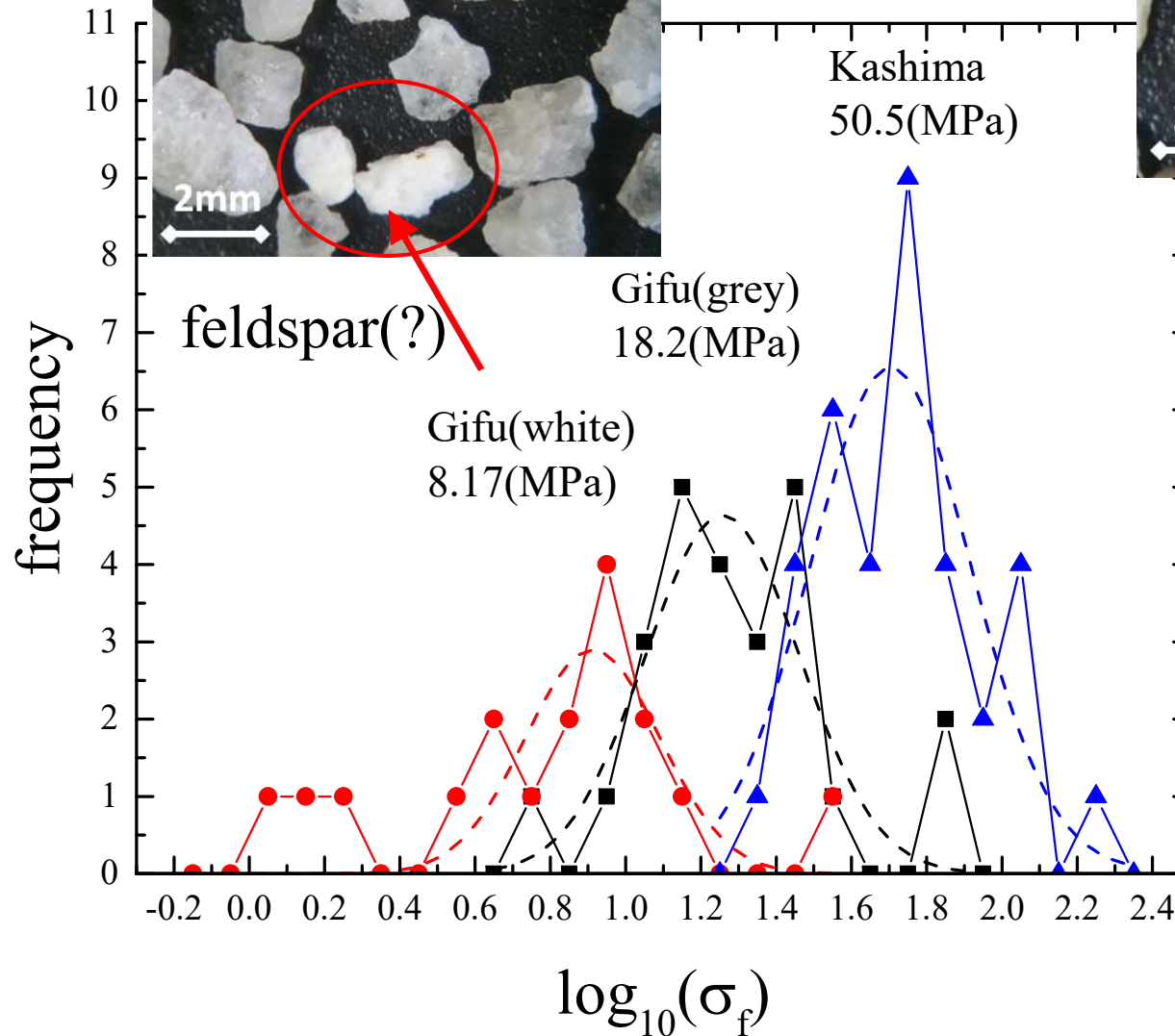
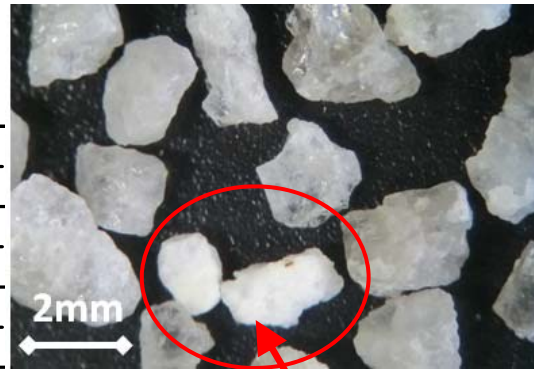
[1] Single grain crushing test



Measure single grain crushing stress



Grain crushing stress



$$\sigma_f = \frac{F_f}{h^2}$$

Strength distribution is fitted by log-normal one.

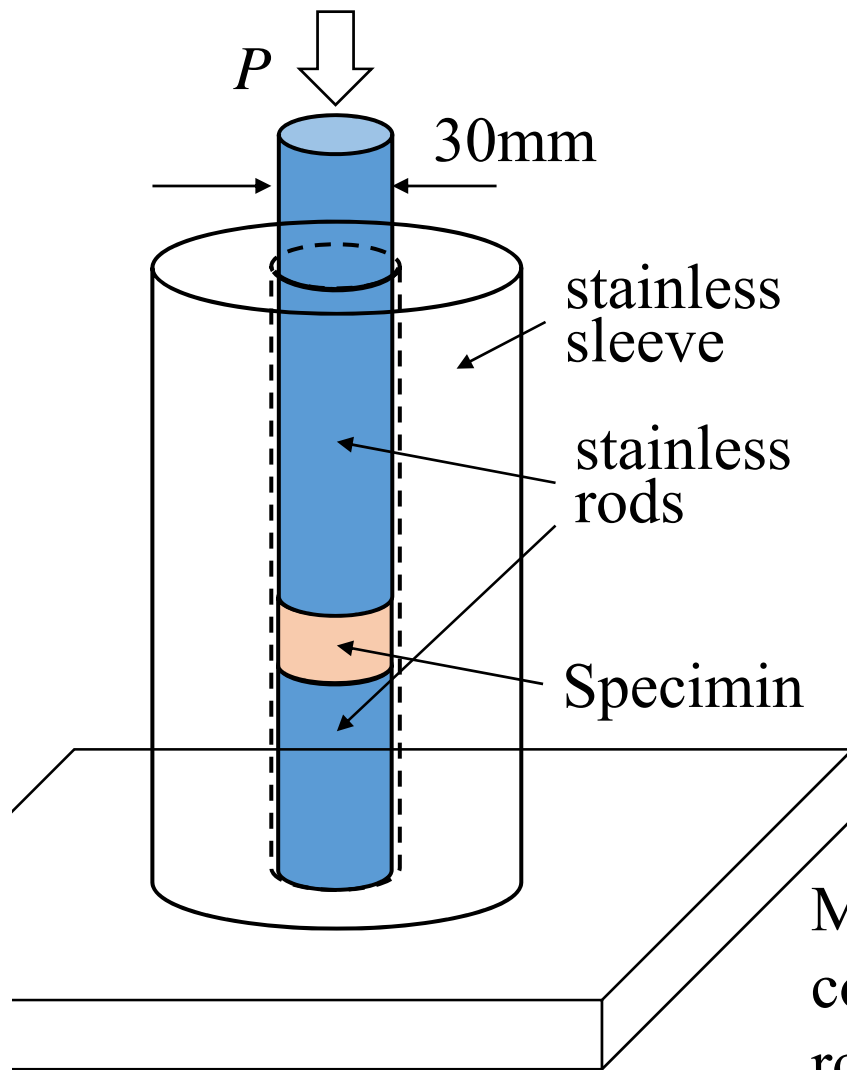
Gifu < Kashima



[2] One-dimensional compression test

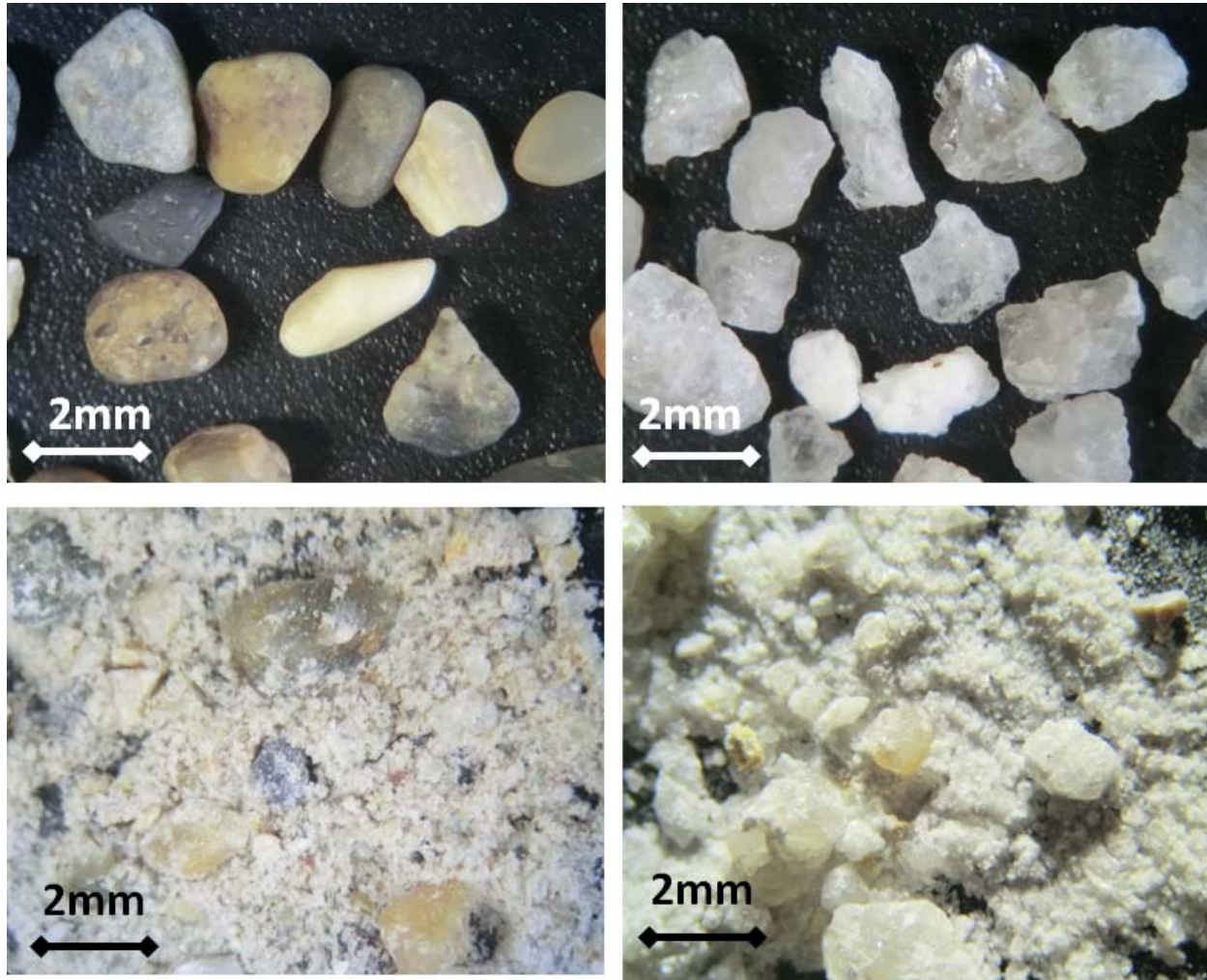
[2] One-dimensional compression test

Sato et al. KKHTCNN, 2015



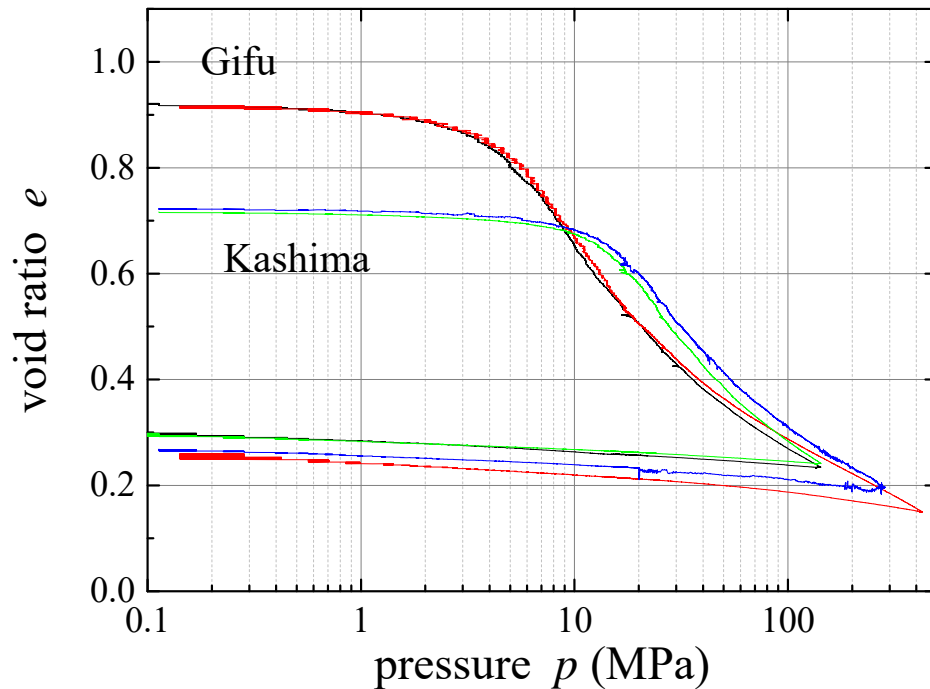
Measured volumetric strain is modified considering the deformation of stainless rods, sleeve and road cell.

[2] One-dimensional compression test

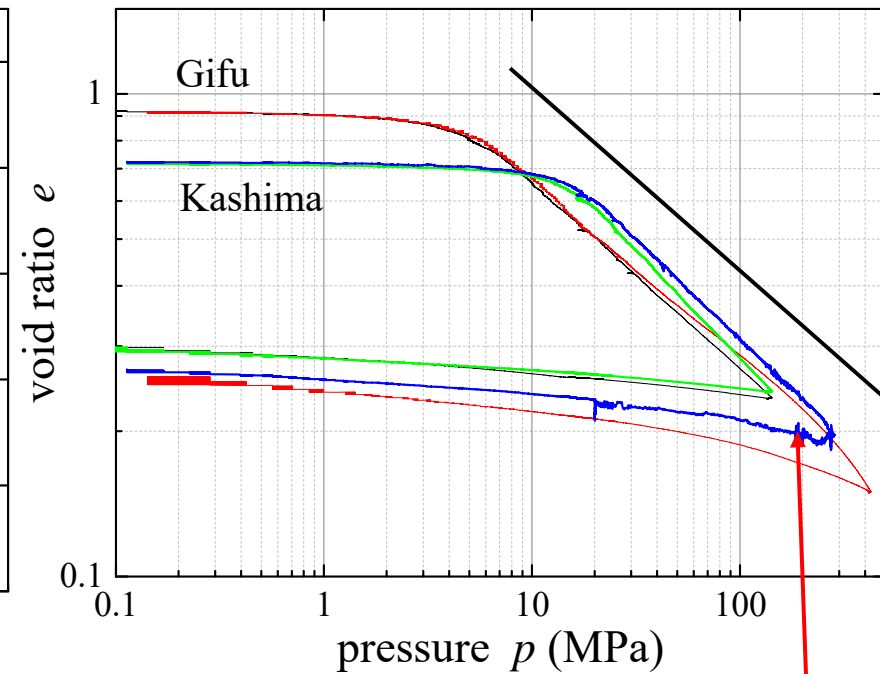


After 400(MPa) loading

Void ratio – Pressure relation



Conventional e - $\log(p)$

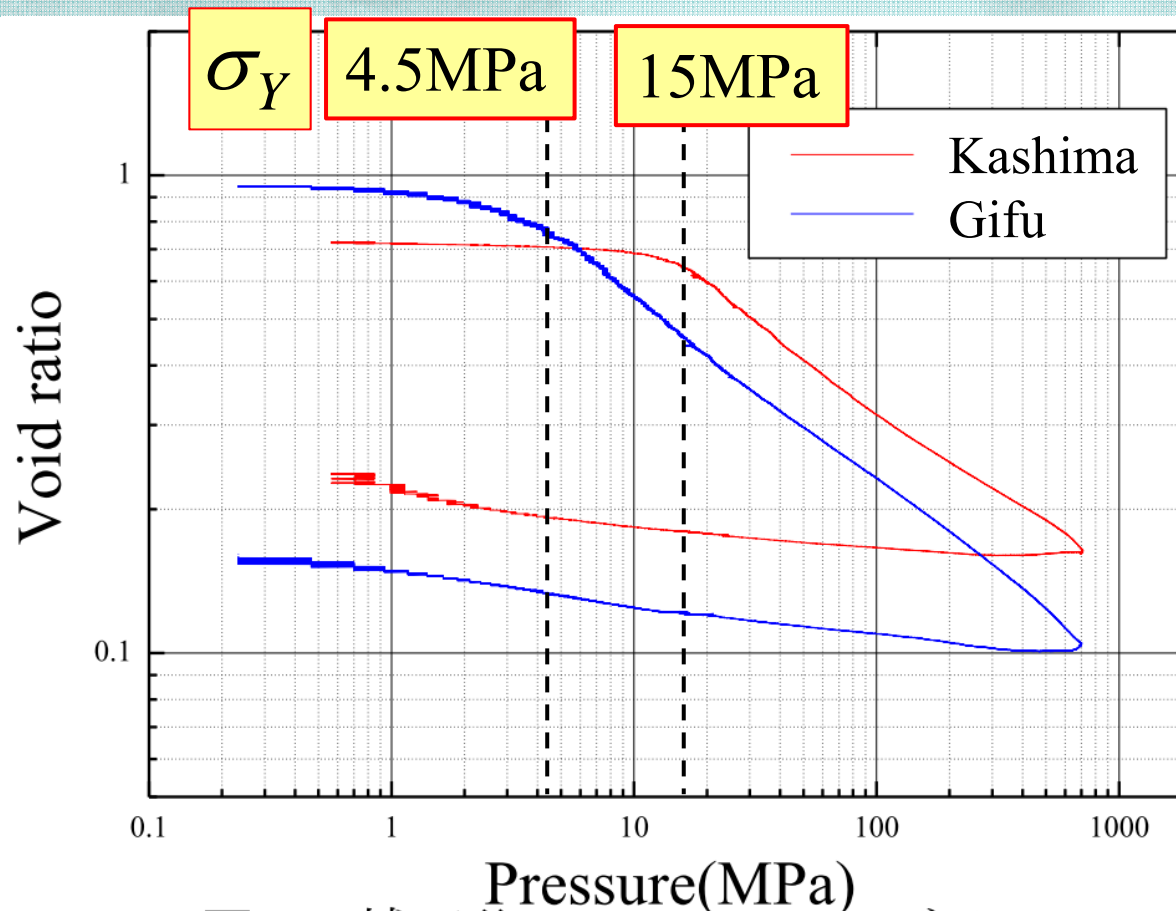


$\log(e)$ - $\log(p)$

Plastic compression regime
is well modeled by a straight line

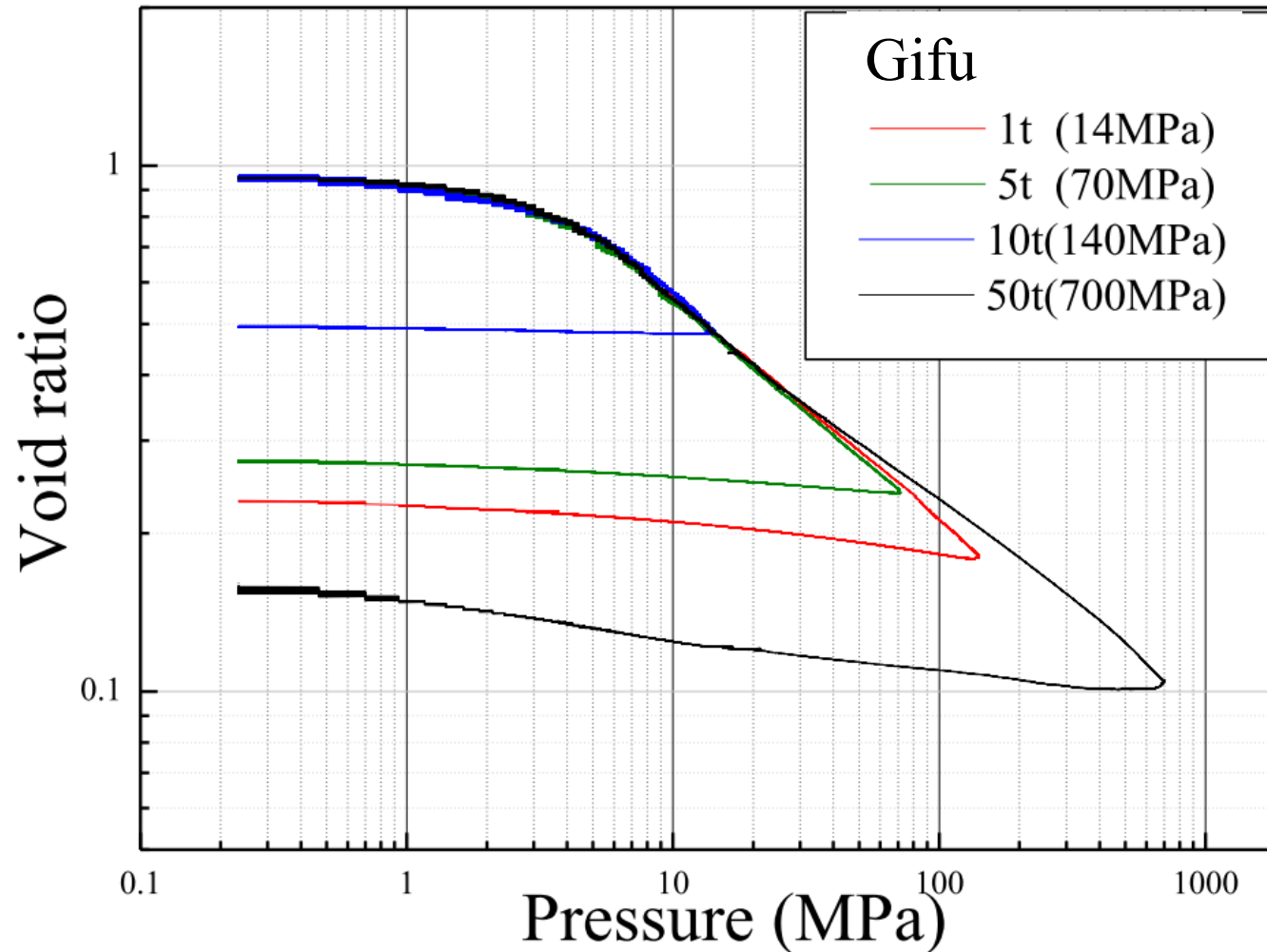
Butterfield, Geotechnique, 1979

log (e) – log (p) relation



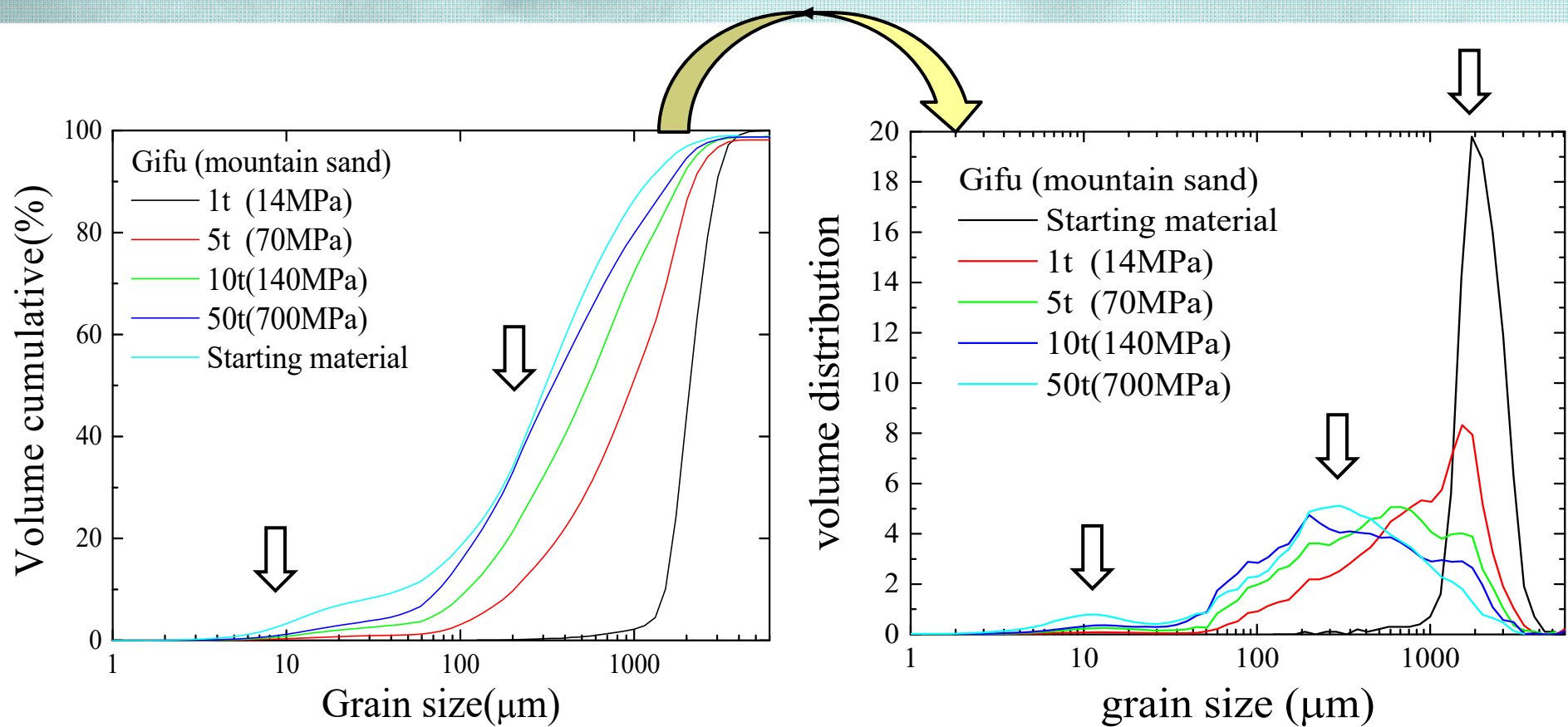
- * σ_Y ratio (15MPa : 4.5MPa) \sim σ_f ratio (50MPa : 18MPa)
- * The difference of the value itself must be due to **heterogeneous stress transmission (force chain)**.
- * Both have the similar **power** in plastic compression regime

log (e) – log (p) relation



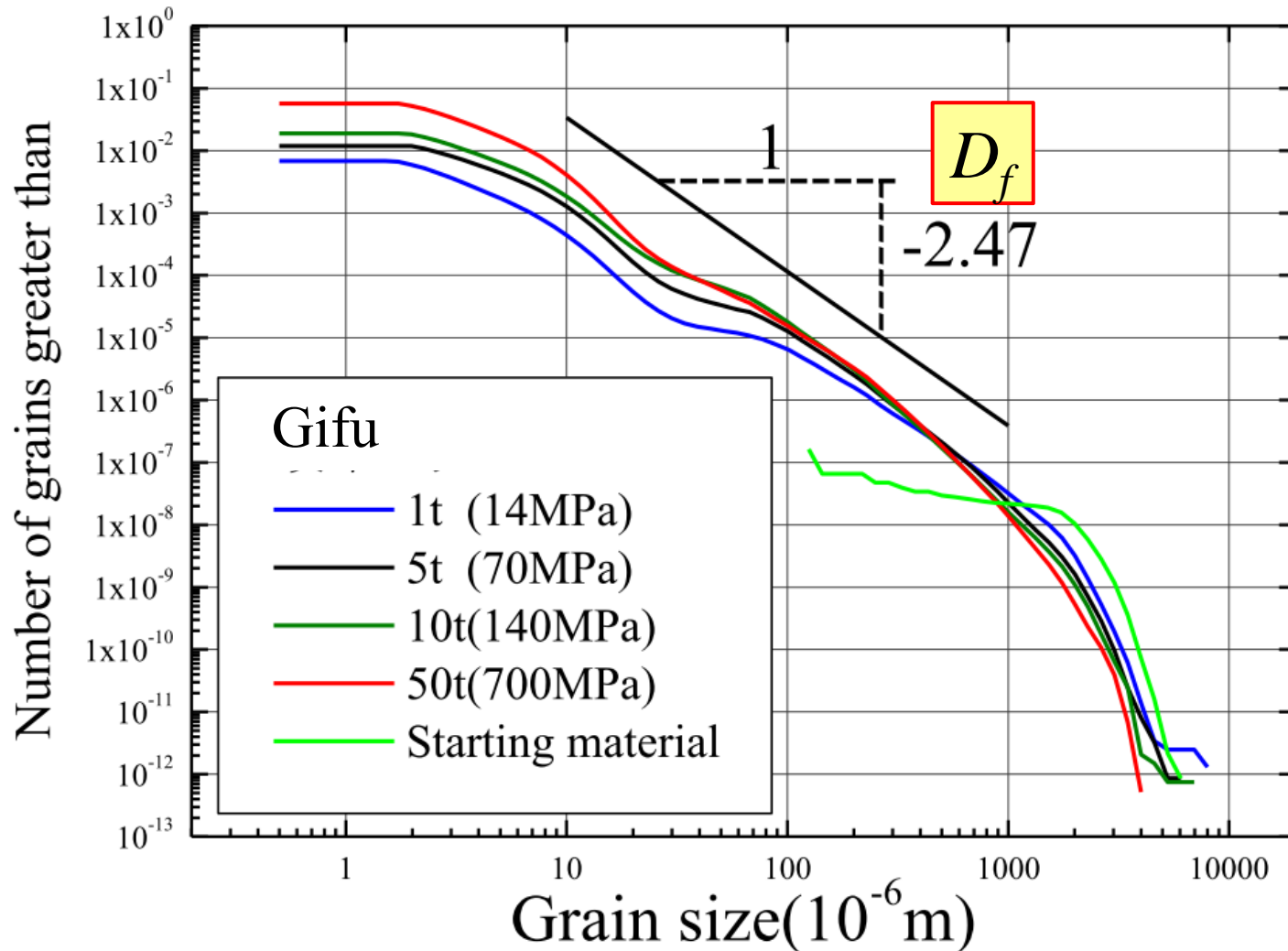
Various loading level → Evolution of grain size distribution

Evolution of grain size distribution



Steady peaks appear during crushing process

Grain size distribution (cumulative number)



$$P(d) = A \cdot d^{-D_f}$$

Fractal (or power law) distribution

TABLE 1. Fractal Dimensions for a Variety of Fragmented Objects

| Object | Reference | Fractal Dimension D |
|---|-----------------------------------|-----------------------|
| Projectile fragmentation of gabbro with lead | <i>Lange et al.</i> [1984] | 1.44 |
| Projectile fragmentation of gabbro with steel | <i>Lange et al.</i> [1984] | 1.71 |
| Meteorites (Prairie Network) | <i>McCrosky</i> [1968] | 1.86 |
| Artificially crushed quartz | <i>Hartmann</i> [1969] | 1.89 |
| Plane of weakness model | this paper | 1.97 |
| Disaggregated gneiss | <i>Hartmann</i> [1969] | 2.13 |
| Disaggregated granite | <i>Hartmann</i> [1969] | 2.22 |
| FLAT TOP I (chemical explosion, 0.2 kt) | <i>Schoutens</i> [1979] | 2.42 |
| Asteroids (theory) | <i>Hellyer</i> [1971] | 2.48 |
| PILED RIVER (nuclear explosion, 61 kt) | <i>Schoutens</i> [1979] | 2.50 |
| Broken coal | <i>Bennett</i> [1936] | 2.50 |
| Interstellar grains | <i>Mathis</i> [1979] | 2.50 |
| Asteroids (theory) | <i>Dohnanyi</i> [1969] | 2.51 |
| Projectile fragmentation of basalt | <i>Fujiwara</i> [1977] | 2.56 |
| Sandy clays | <i>Hartmann</i> [1969] | 2.61 |
| Terrace sands and gravels | <i>Hartmann</i> [1969] | 2.82 |
| Pillar of strength model | <i>Allègre et al.</i> [1982] | 2.84 |
| Glacial till | <i>Hartmann</i> [1969] | 2.88 |
| Stony meteorites | <i>Hawkins</i> [1960] | 3.00 |
| Asteroids | <i>Donnison and Sugden</i> [1984] | 3.05 |
| Ash and pumice | <i>Hartmann</i> [1969] | 3.54 |

Fractal GSD (Turcotte 1986)

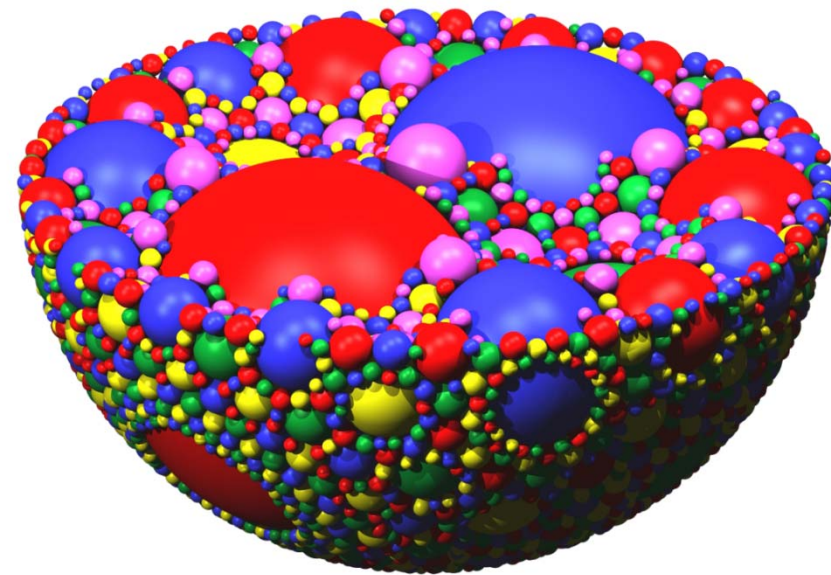
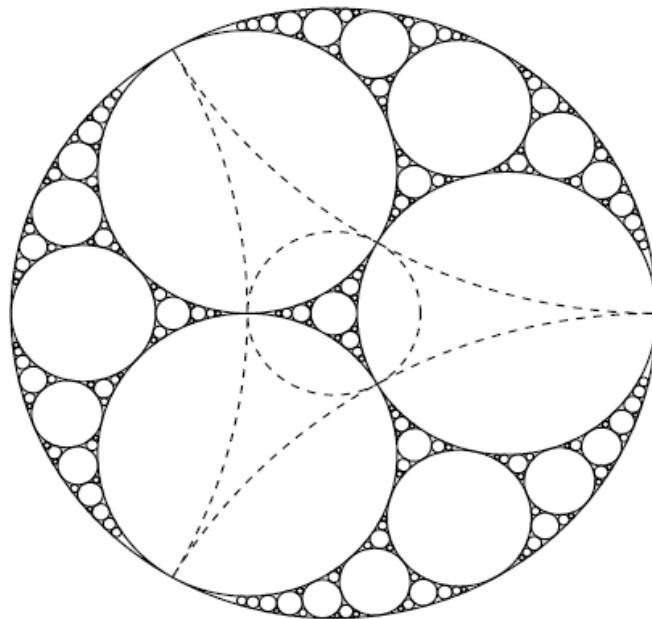
$$D_f = 1.4 \sim 3.5$$

Apollonian sphere packing

The observed power D is close to

that in **Apollonian sphere packing** ($D_f = -2.47$)

(Borkovec, M., De Paris, W., Peikert, R., *Fractals*, Vol. 2, No. 4, 521–526, 1994.)

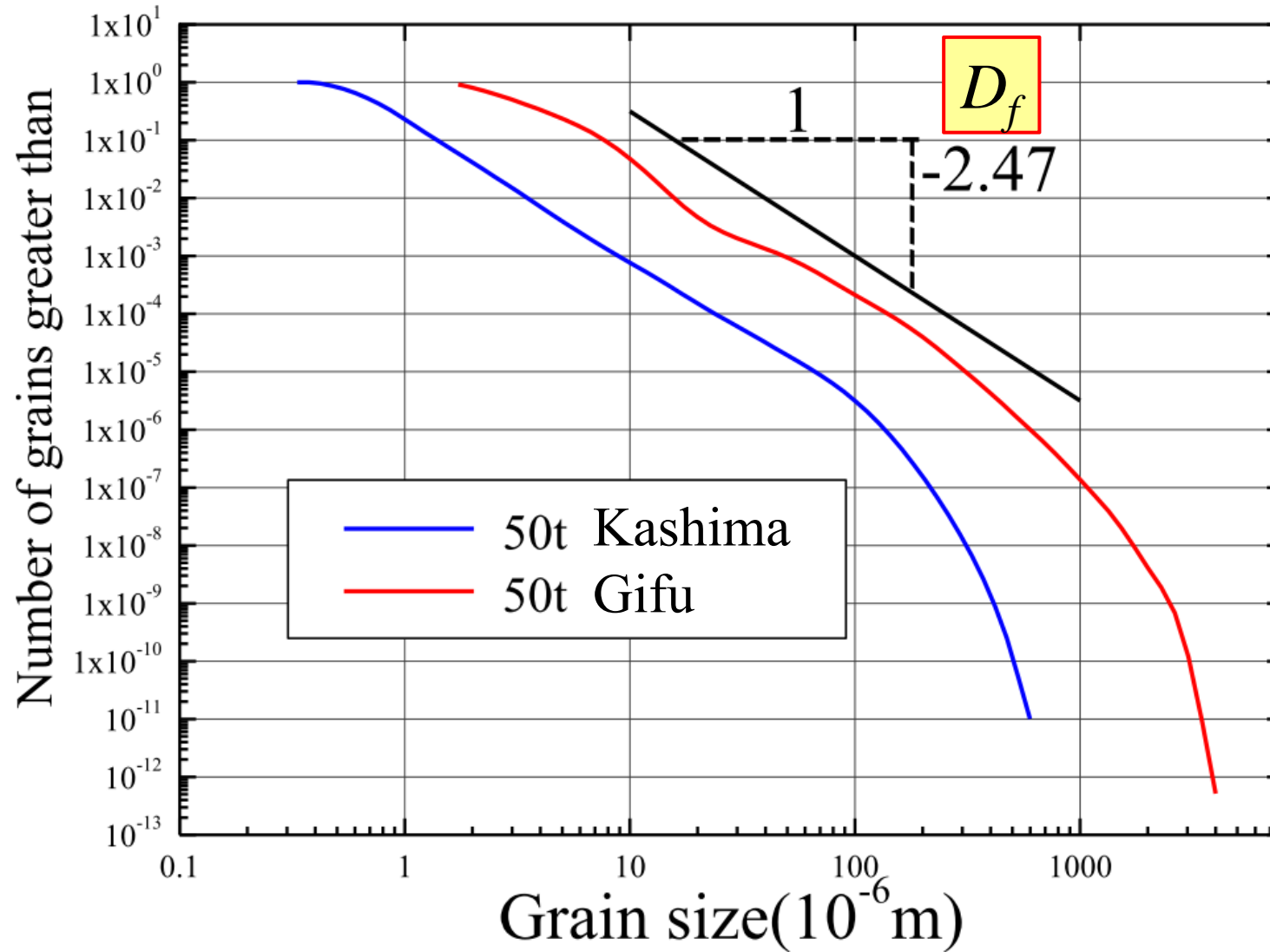


Possible comminution mechanism under **confinement**

→ The Largest particles fitting in the pore survives

→ Second and third peaks appear

Grain size distribution (cumulative number)

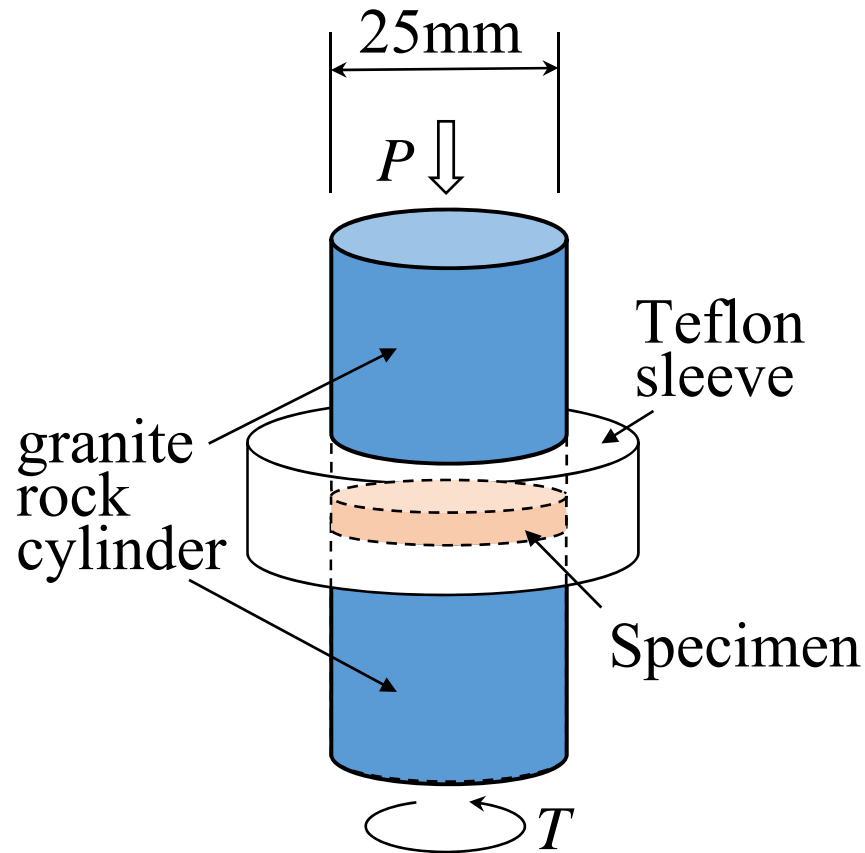




[3] Rotary shear test

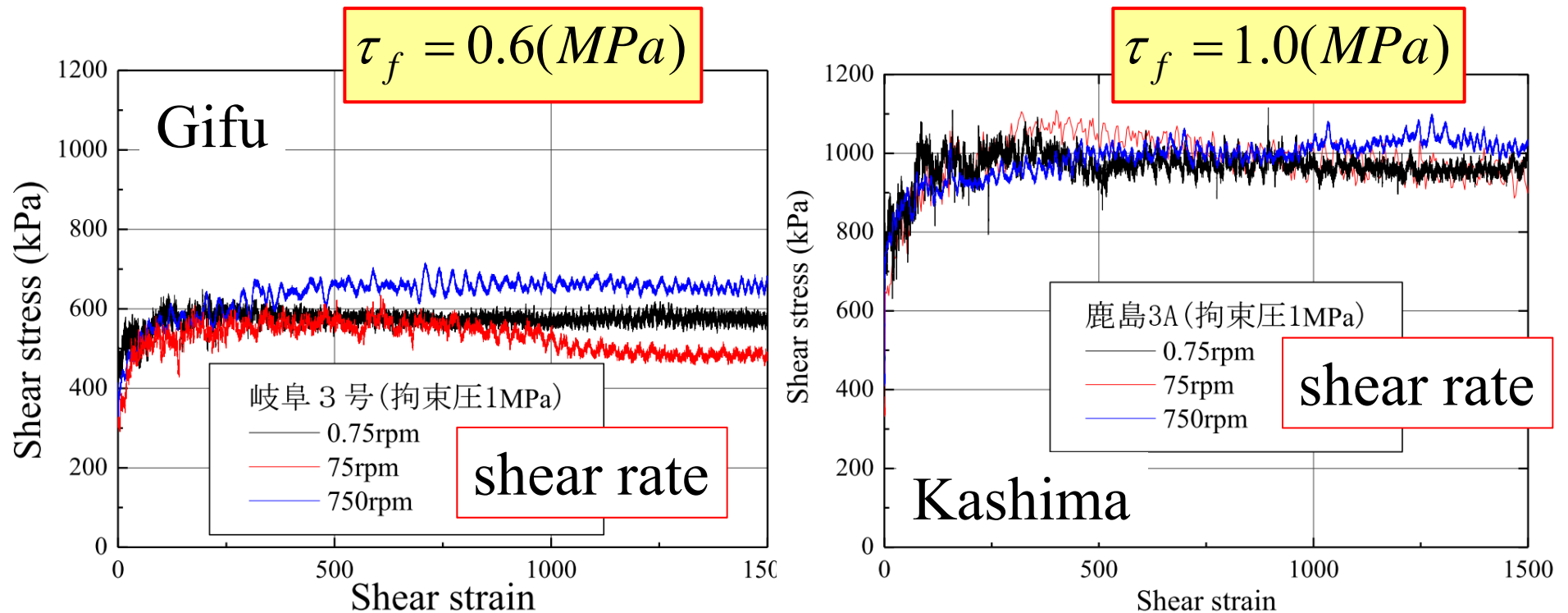
[3] Rotary shear test

Sato et al. KKHTCNN, 2015



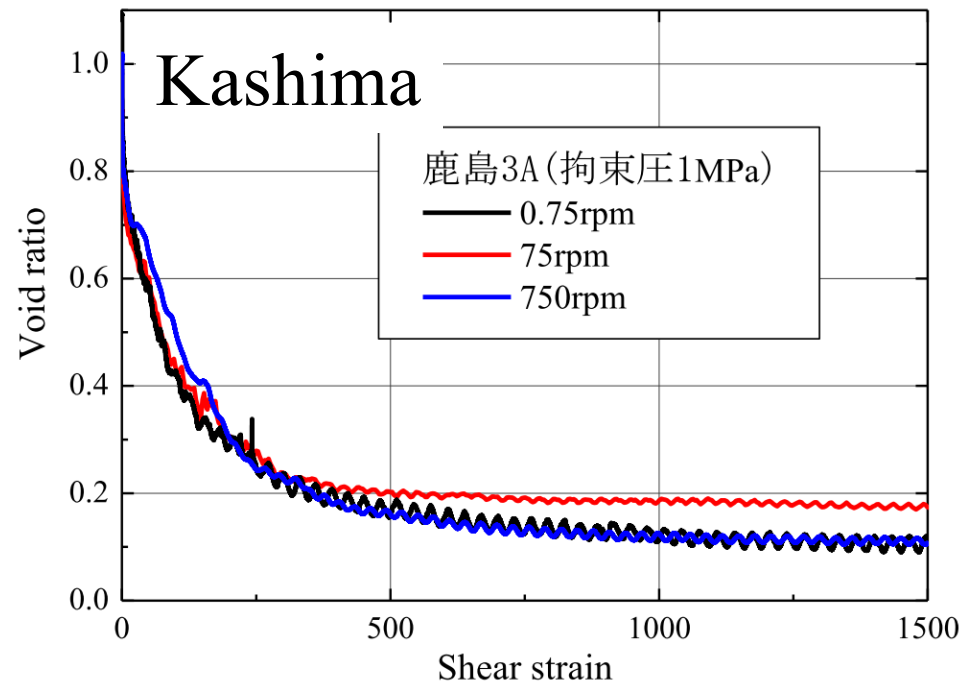
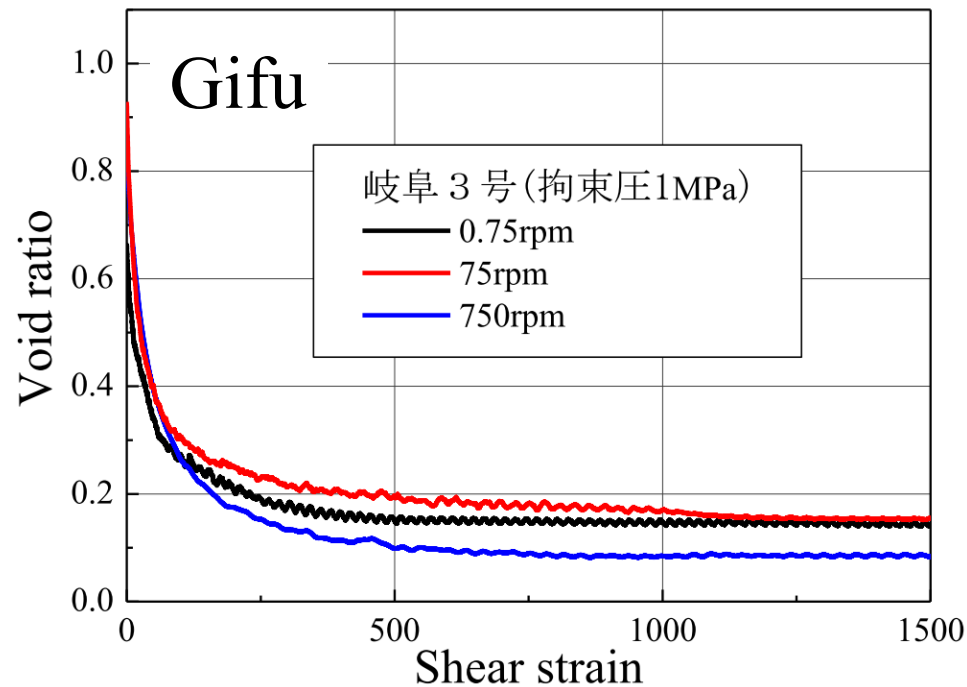
The device can apply wide range of shear rate
 $0.75(\text{rpm}):(\dot{\gamma}=0.21(1/\text{s})) \sim 750(\text{rpm}):(\dot{\gamma}=210(1/\text{s}))$

Shear stress – shear strain relation



The ratio of peak shear stress is not consistent with the ratio of single grain crushing stress
 → Effect of **friction** (shear without crushing) is included

Void ratio evolution

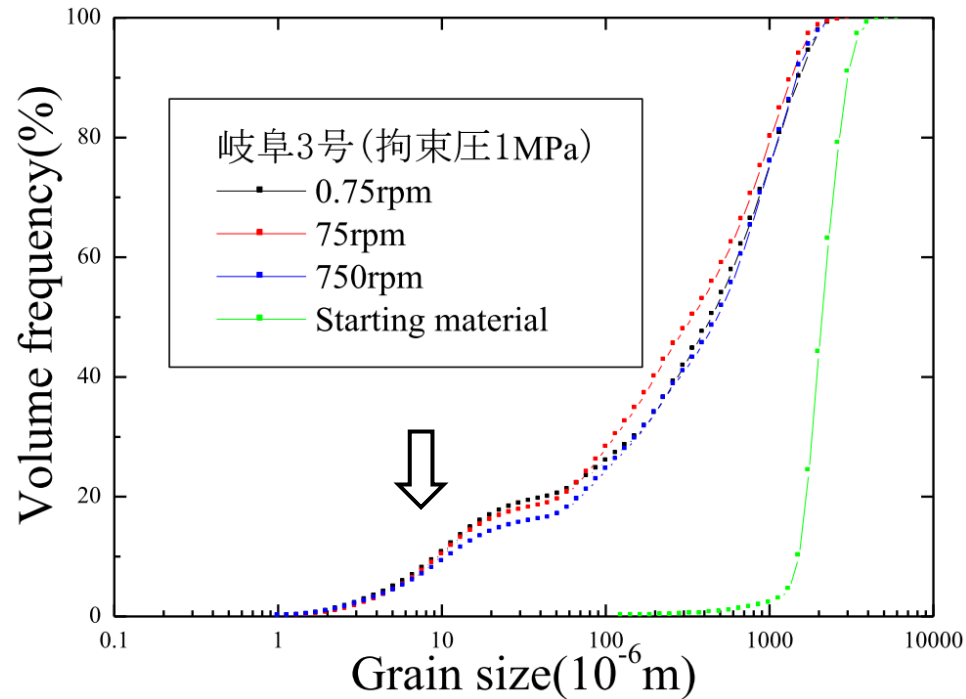


Void ratio reach the residual value (0.1 ~ 0.2)
after long shear (shear strain ~ 500)

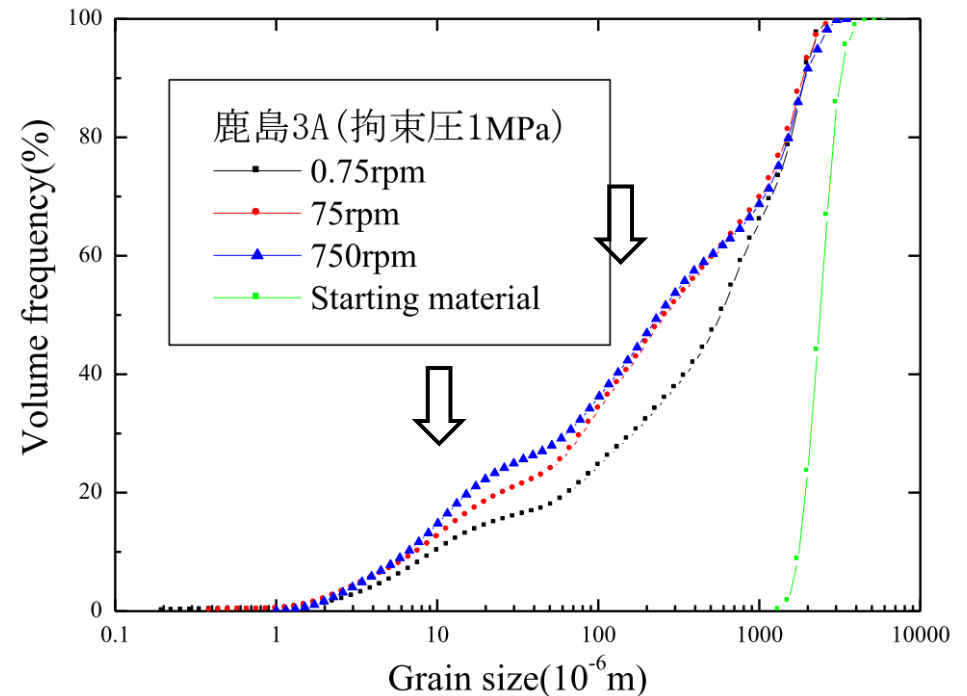
→ **Constant fabric change** provides constant opportunity
of crushing (**Force chain** must play an important role)

Grain size distribution (cumulative volume)

Gifu

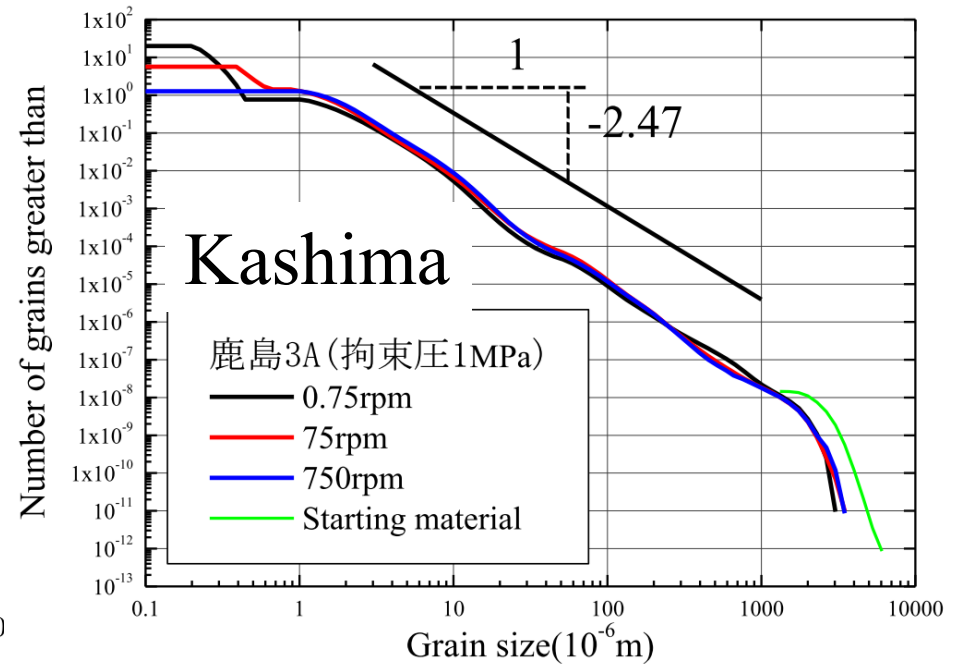
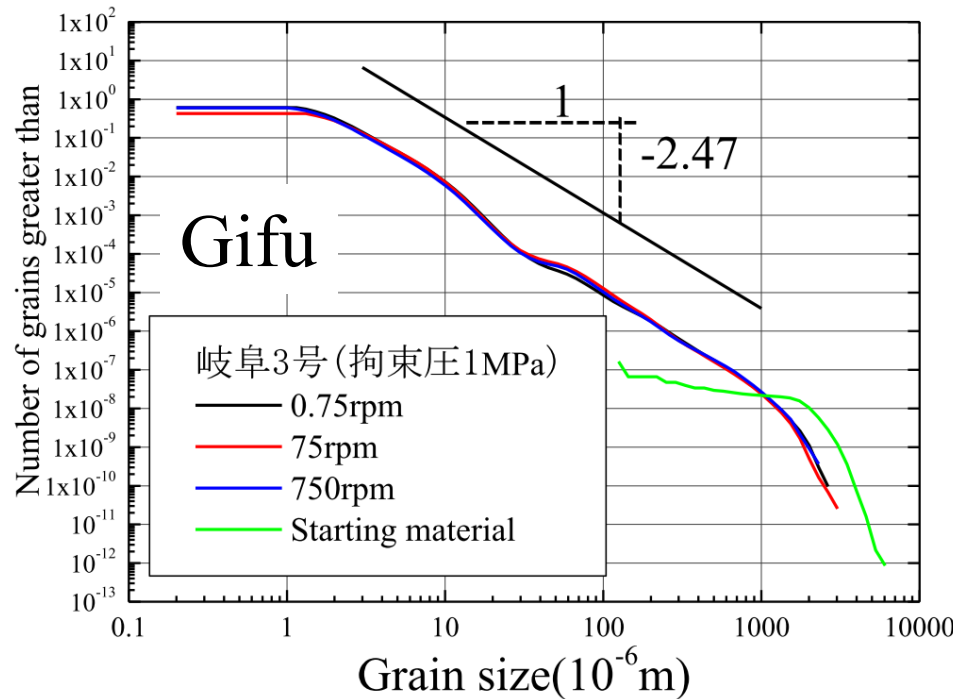


Kashima



The second and third peaks were observed.

Grain size distribution (cumulative number)



The power is similar to that in **Apollonian sphere packing**



[4] GSD evolution model

GSD evolution model

(1) Mono-disperse granular system (Radius=R)

$$e_0 = \frac{V_{void}}{V_{solid}}$$

(2) Fill the small grains (Radius=r) into the void with the same void ratio

$$e_0 = \frac{V_{void} - V_{solid}^S}{V_{solid}^S}$$

(3) Number ratio of small grains to large grains

$$\frac{N_S}{N_L} = \frac{V_{Solid}^S}{V_{Solid}^L} \left(\frac{R}{r} \right)^3 = \frac{e_0}{1 + e_0} \alpha^{-3} \quad \alpha = \frac{r}{R}$$

size ratio

GSD evolution model

$$\frac{N_S}{N_L} = \frac{V_{Solid}^S}{V_{Solid}^L} \left(\frac{R}{r} \right)^3 = \frac{e_0}{1 + e_0} \alpha^{-3}$$

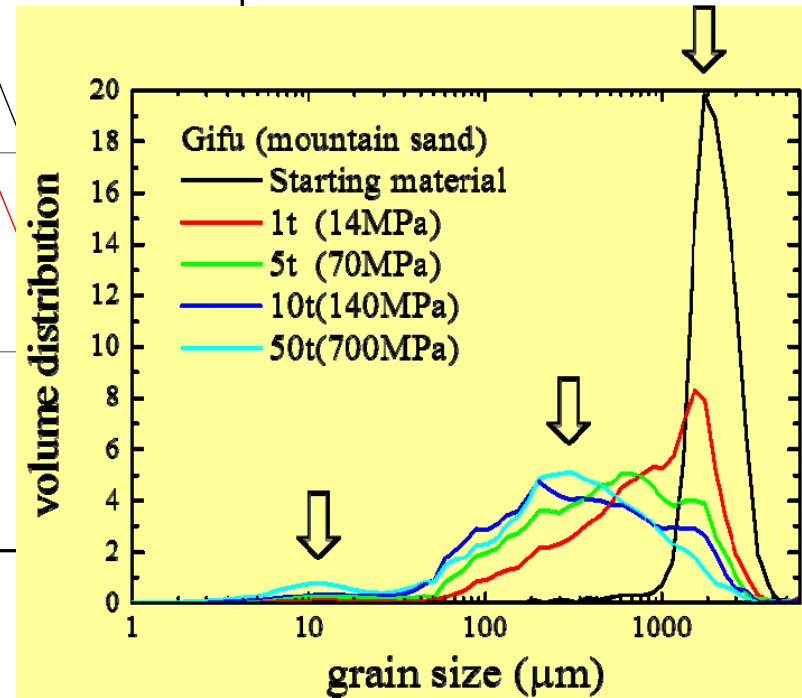
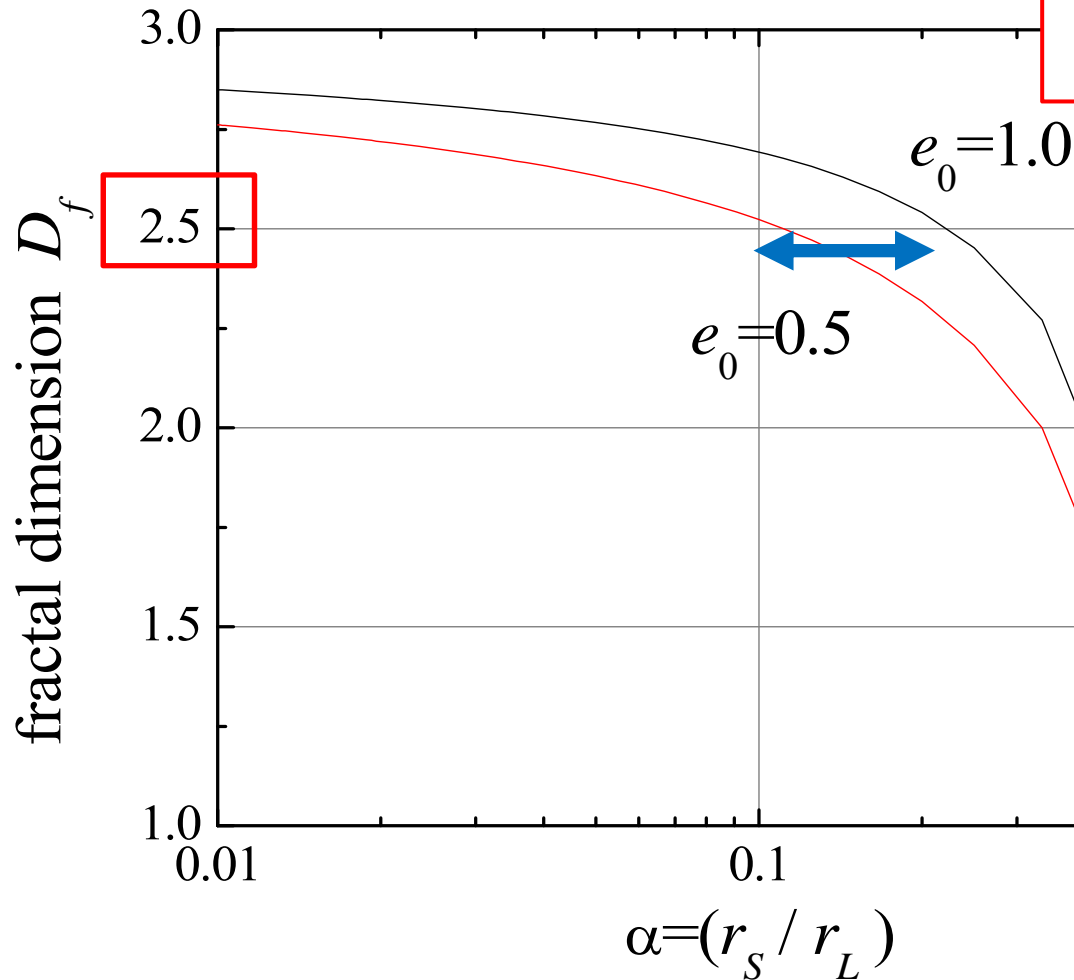
$$\log N_S - \log N_L = \log \left(\frac{e_0}{e_0 + 1} \right) - 3(\log r - \log R)$$

(4) Fractal dimension

$$D_f \equiv \frac{\log N_S - \log N_L}{\log r - \log R} = \frac{1}{\log \alpha} \log \left(\frac{e_0}{e_0 + 1} \right) - 3$$

GSD evolution model

$$D_f = \frac{1}{\log \alpha} \log \left(\frac{e_0}{e_0 + 1} \right) - 3$$



Consistent with Experimental observation
on the second and the third peaks

GSD evolution model

(5) Void ratio after adding the small grains

$$e_1 = \frac{e_0}{\frac{e_0 + 1}{e_0} + 1} \quad \Rightarrow \quad 1 + \frac{1}{e_k} = \left(1 + \frac{1}{e_0} \right)^{k+1}$$

Recursive equation

(6) Breakage stress for the smallest grain determines the 1D compression stress (McDowell and Bolton 2008)

$$\sigma_0 = C \cdot r_{\min}^{-3/m}$$

m : Weibull modulus

$5 < m < 10$ in various materials

MODELLING

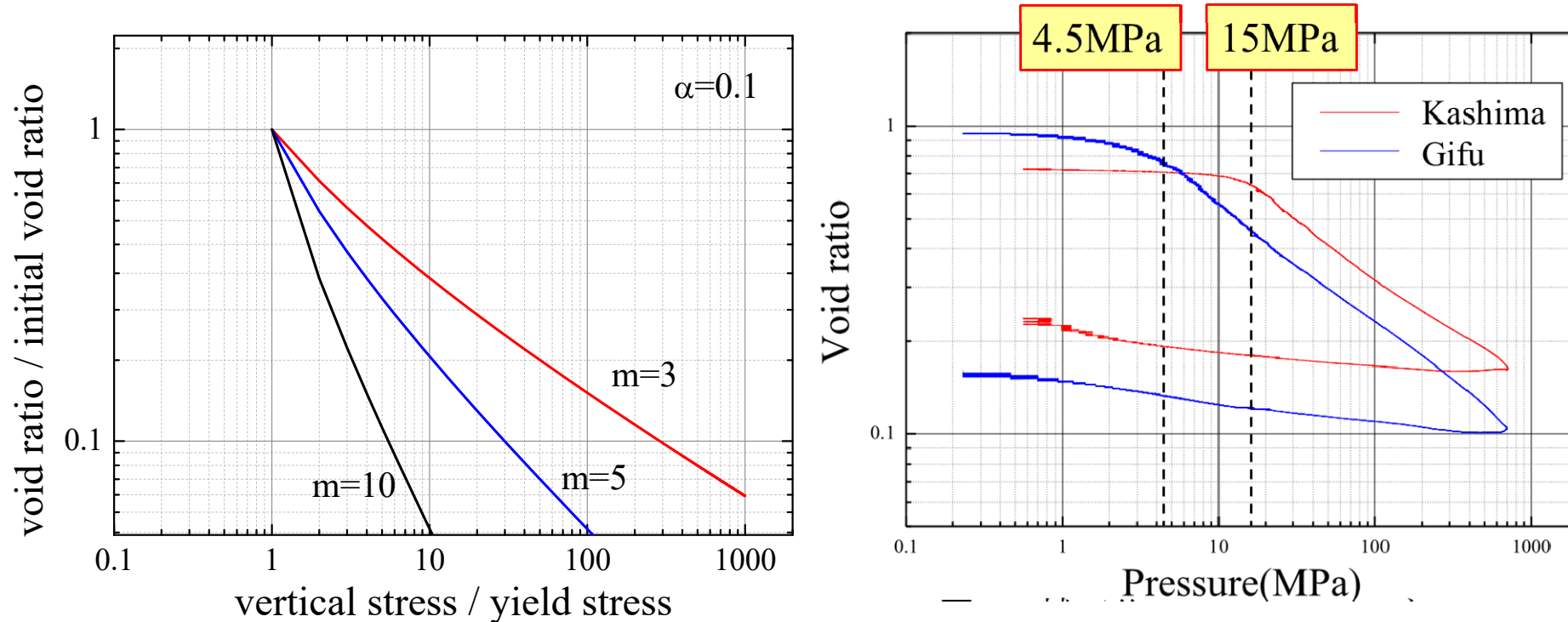
(7) Final relation between e and p

$$r_k = \alpha^k R \quad \Rightarrow \quad k = \frac{\log \frac{r_k}{R}}{\log \alpha}$$

$$\sigma_Y = C \cdot \left(\frac{r_k}{R} \right)^{-3/m} \quad \Rightarrow \quad -\frac{3}{m} \log \left(\frac{r_k}{R} \right) = \log C - \log \sigma_Y$$

$$\therefore \frac{\log \left(1 + \frac{1}{e_k} \right)}{\log \left(1 + \frac{1}{e_0} \right)} = k + 1 = \frac{\log \frac{r_k}{R}}{\log \alpha} + 1 = \frac{-\frac{m}{3} (\log C - \log \sigma_Y)}{\log \alpha} + 1$$

Validation



The model can provide more or less **linear** relation in $\log(e)$ - $\log(p)$ curve.

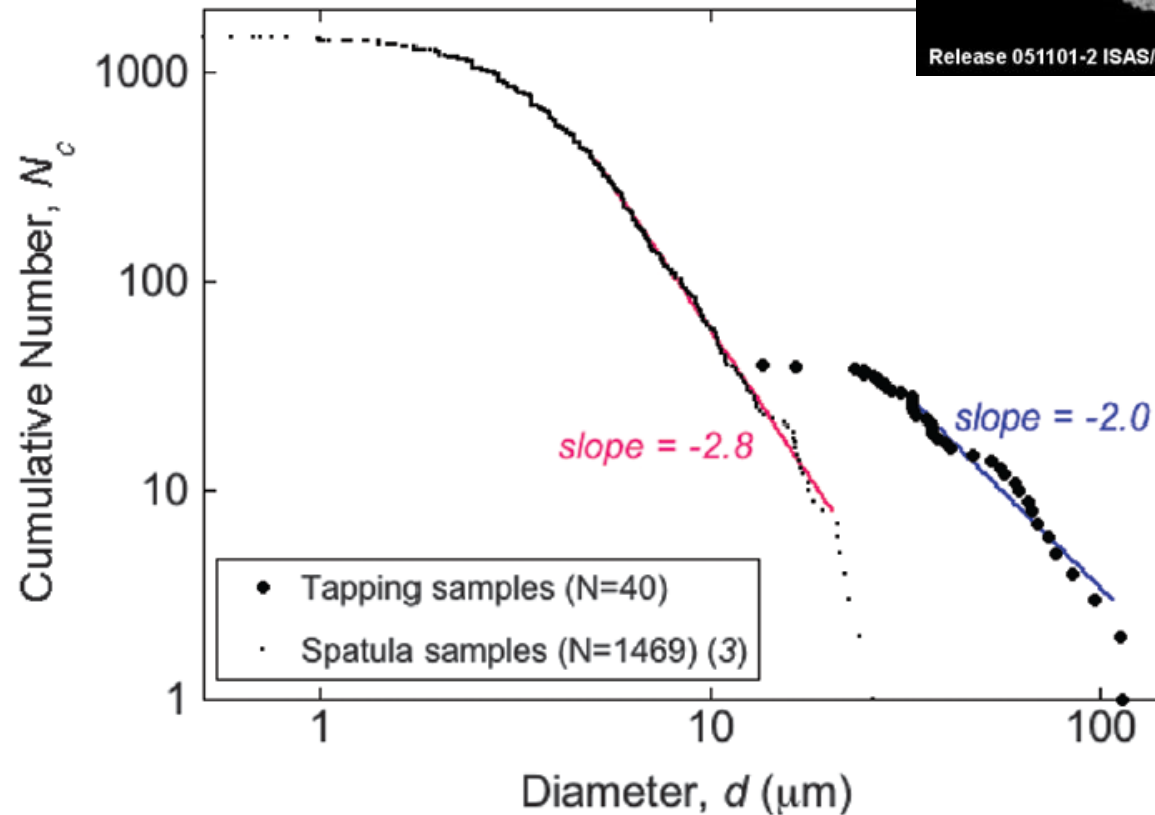
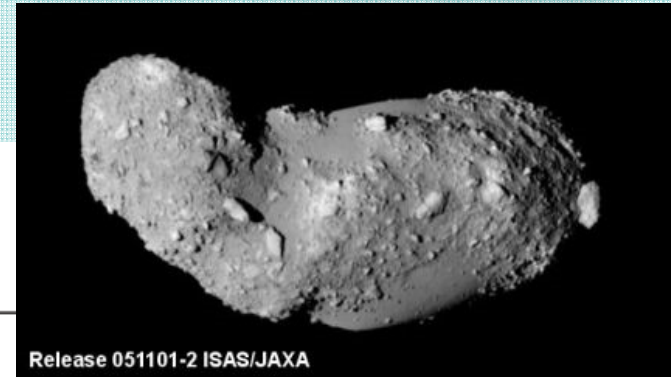
The power becomes consistent with the ODC test when $m=3$ (with $\alpha=0.1$)



[5] Some observation in planetary science

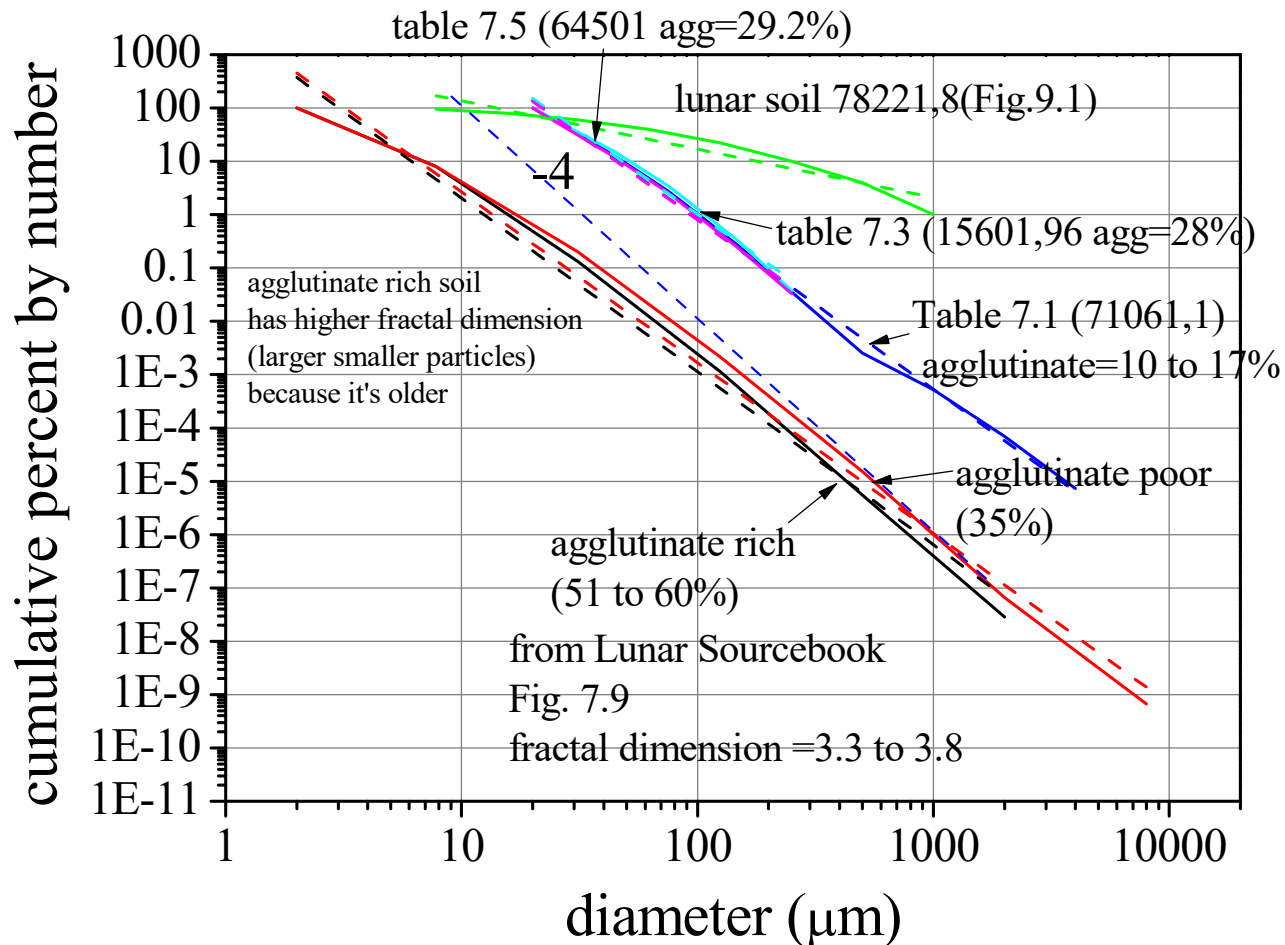
Itokawa regolith

(Tsuchiyama *et al.*, *Science* 2011)



The power $D_f = -2.0 \sim -2.8$ close to ODC test

Lunar soil GSD (*Lunar Sourcebook*)



$$D_f = -3.0 \sim -4.0$$

Higher fractal dimension due to **repeated comminution?**
 or due to **unconfined comminution?**

Conclusions

- (1) Single grain crushing test
 - (2) One dimensional compression test
 - (3) Rotary shear test
- were performed with the common geomaterials.

We observe

- * Linear $\log(e)$ - $\log(p)$ curve in plastic compression
- * Steady peaks and fractal nature in grain size distribution

The proposed model reproduce the above observation.