

# Mechanical and Numerical Modeling of Gas Hydrate Bearing Sediments

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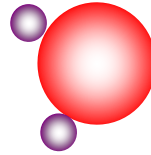
# Outline

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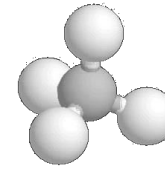
- ❑ **Brief introduction to Hydrate Bearing Sediments (HBS)**
- ❑ **Basic components of the proposed coupled THCM framework for HBS**
- ❑ **Simple benchmarks involving HBS**
- ❑ **HBS mechanical modeling**
- ❑ **Validation**
- ❑ **Final remarks**

# Gas Hydrate Bearing Sediments (HBS)

➤ Gas hydrates are solid ice-like materials that consists of guest gas molecules encased in a water matrix.



Water



+ Methane CH<sub>4</sub>

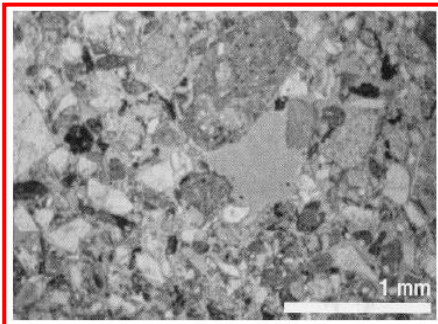


= Methane hydrate

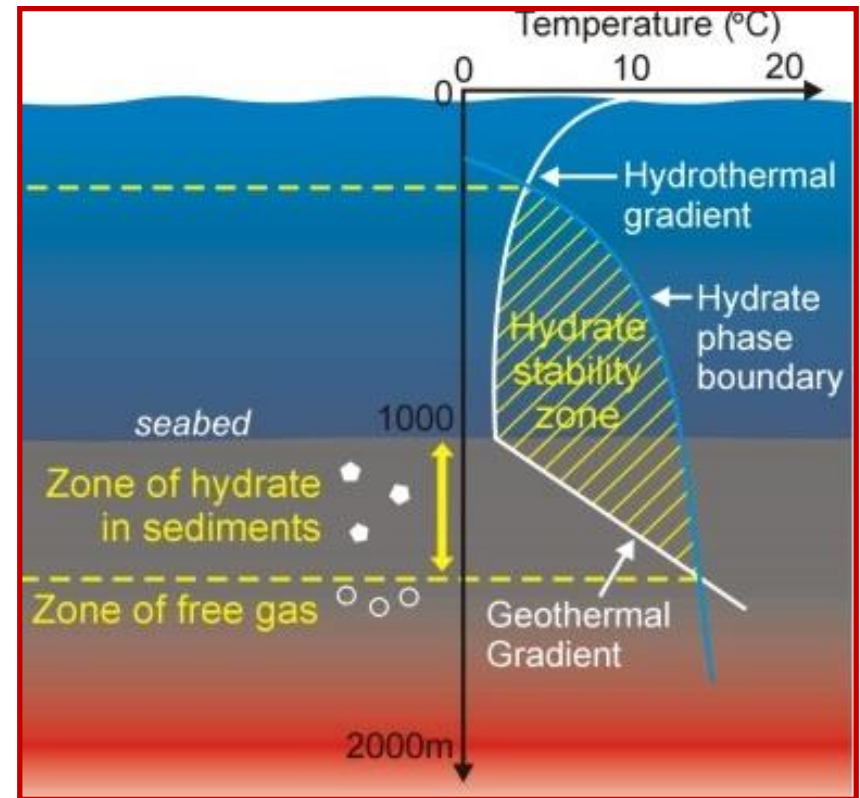
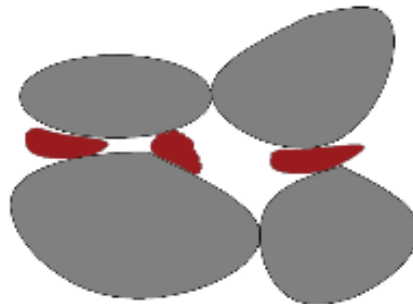
➤ Hydrate bearing sediments are typically found in:

- ✓ marine sediments
- ✓ permafrost

➤ Methane hydrates sediments are highly compacted (stable) under deposit conditions and are likely to behave as bonded sedimentary soils

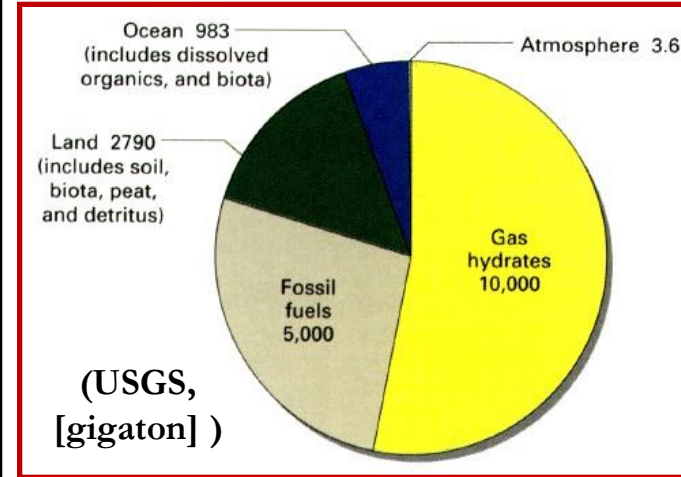
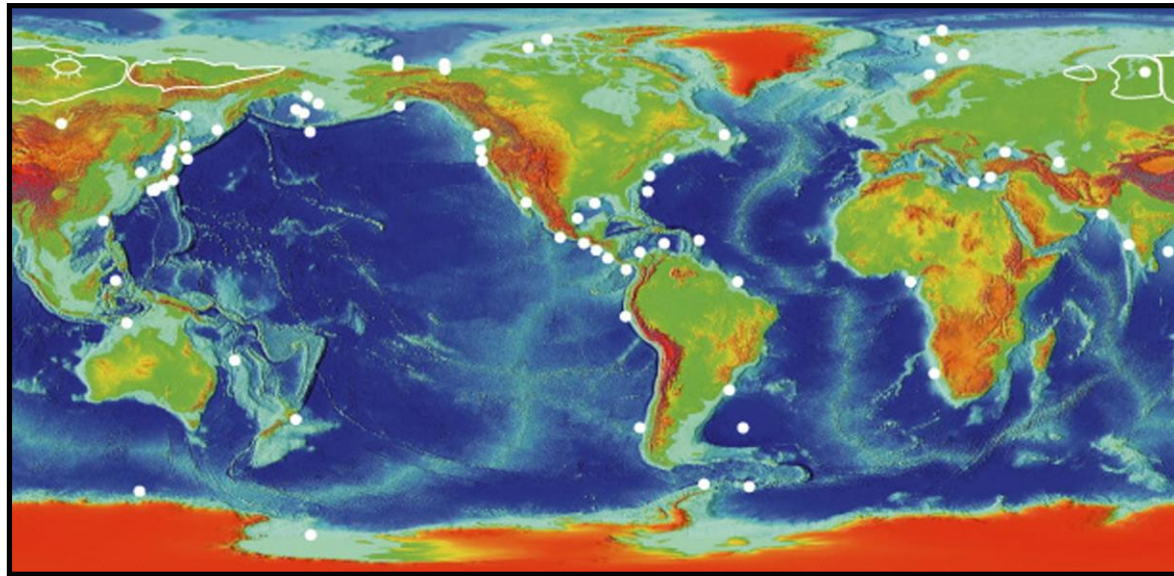


Soga et al. (2006)



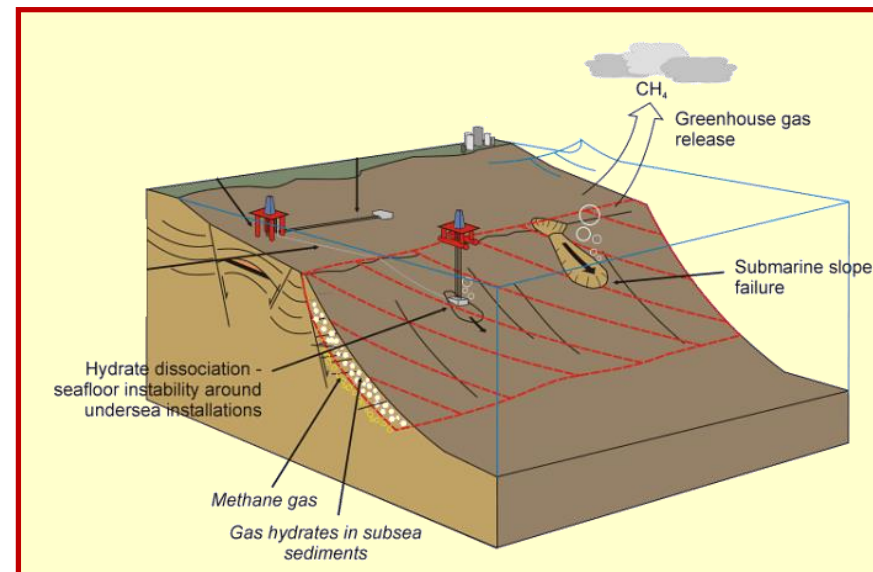


# Gas Hydrate Bearing Sediments (HBS)

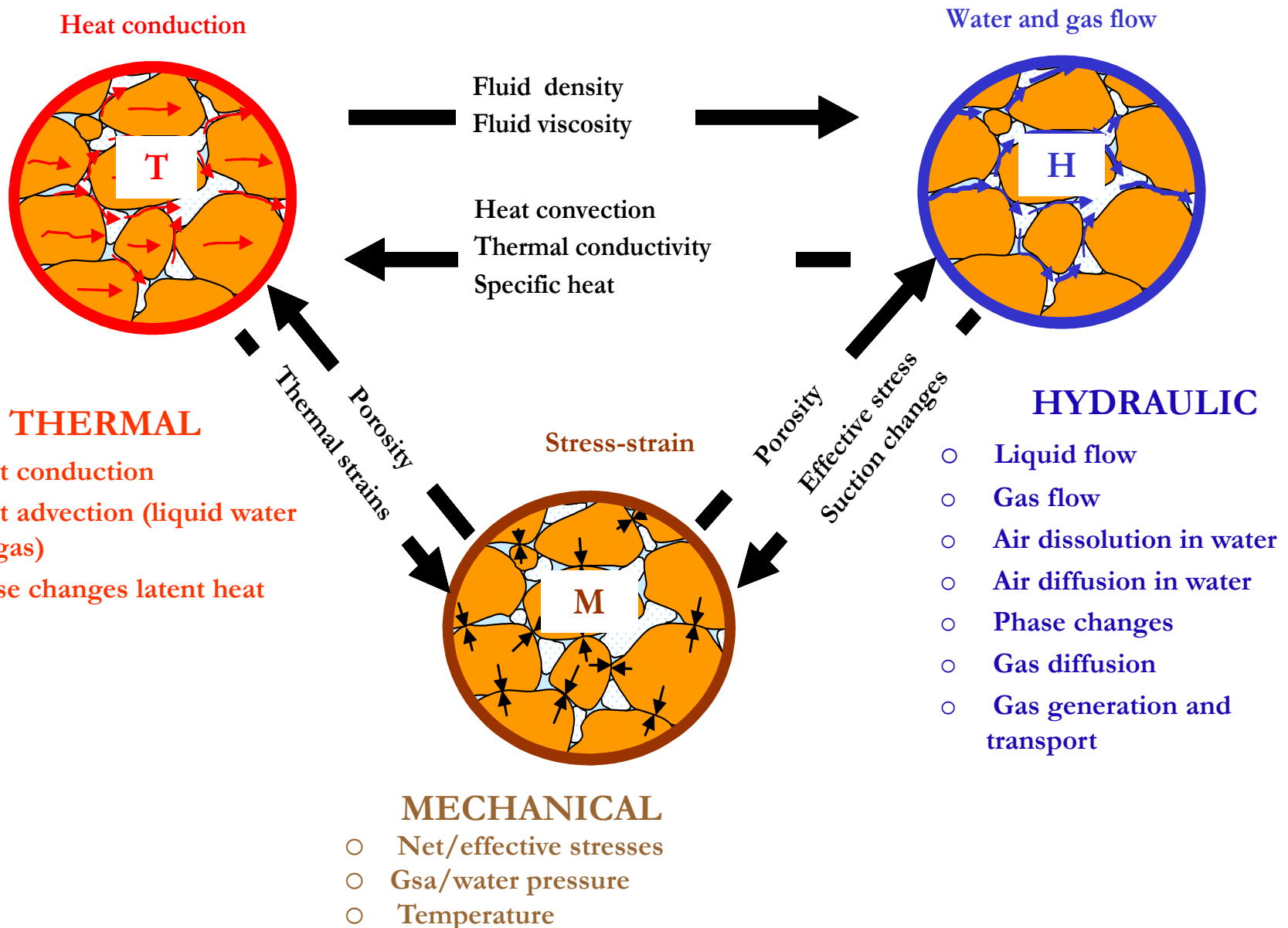


## Relevance:

- ✓ methane recovery, energy resource
- ✓ instability (boreholes and slopes)
- ✓ effect on submarine infrastructure
- ✓ climate change
- ✓ CO<sub>2</sub> sequestration

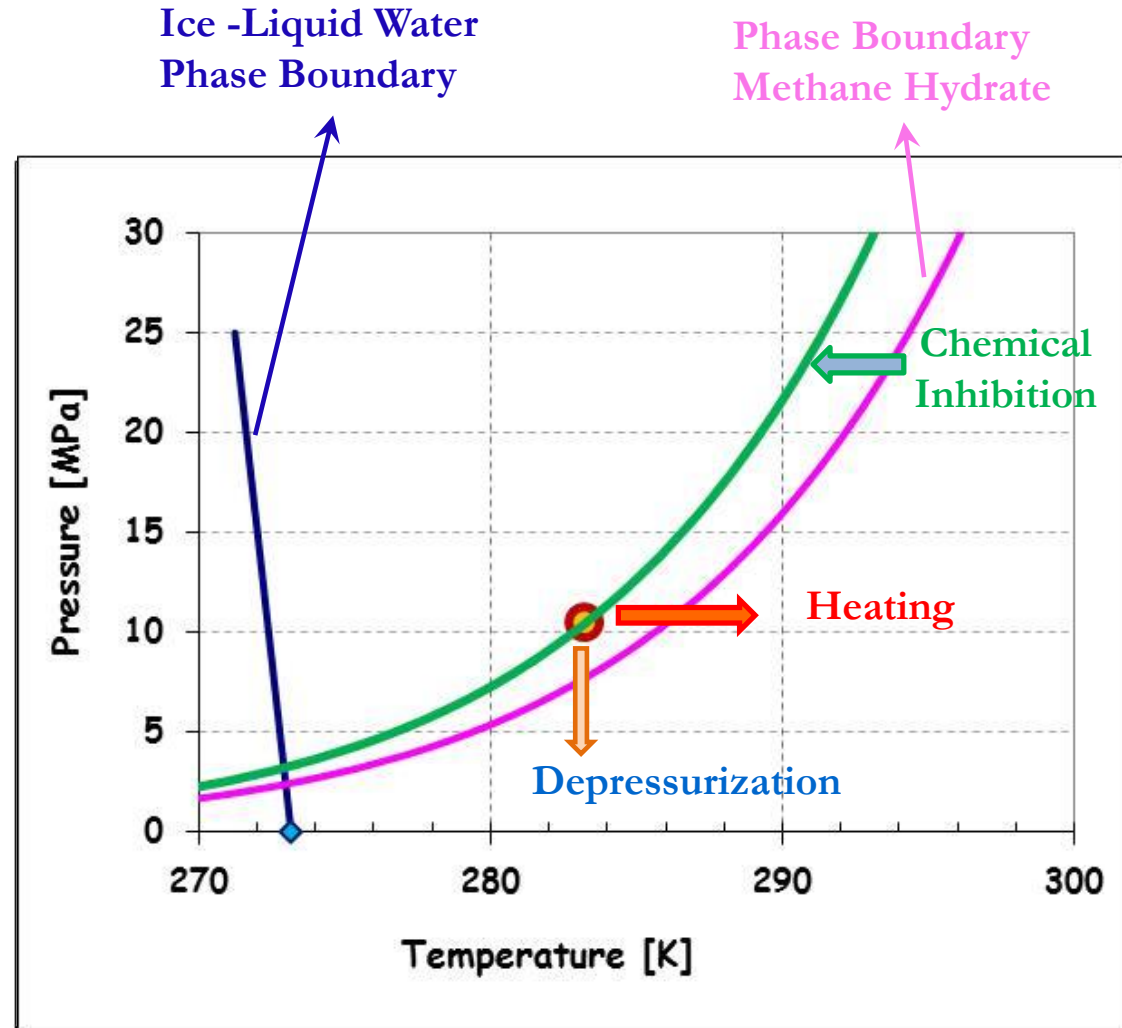
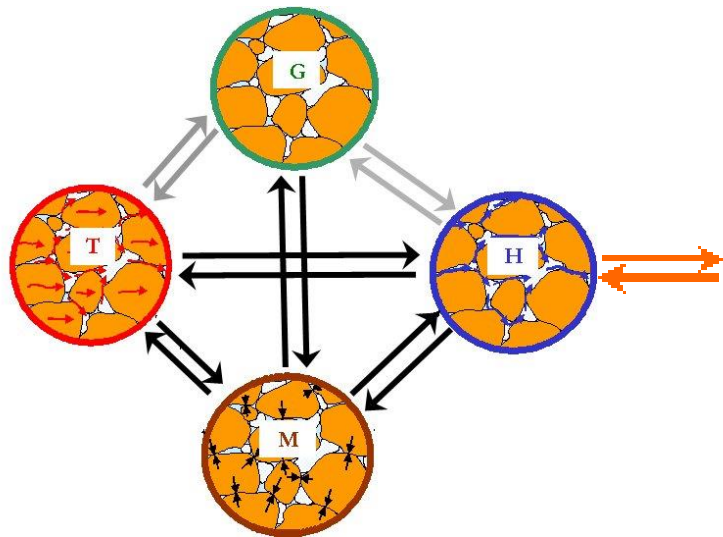


# THM Coupled Phenomena



# HBS – Coupled Phenomena and Phase Boundaries

## Hydrate Dissociation



# Numerical Code

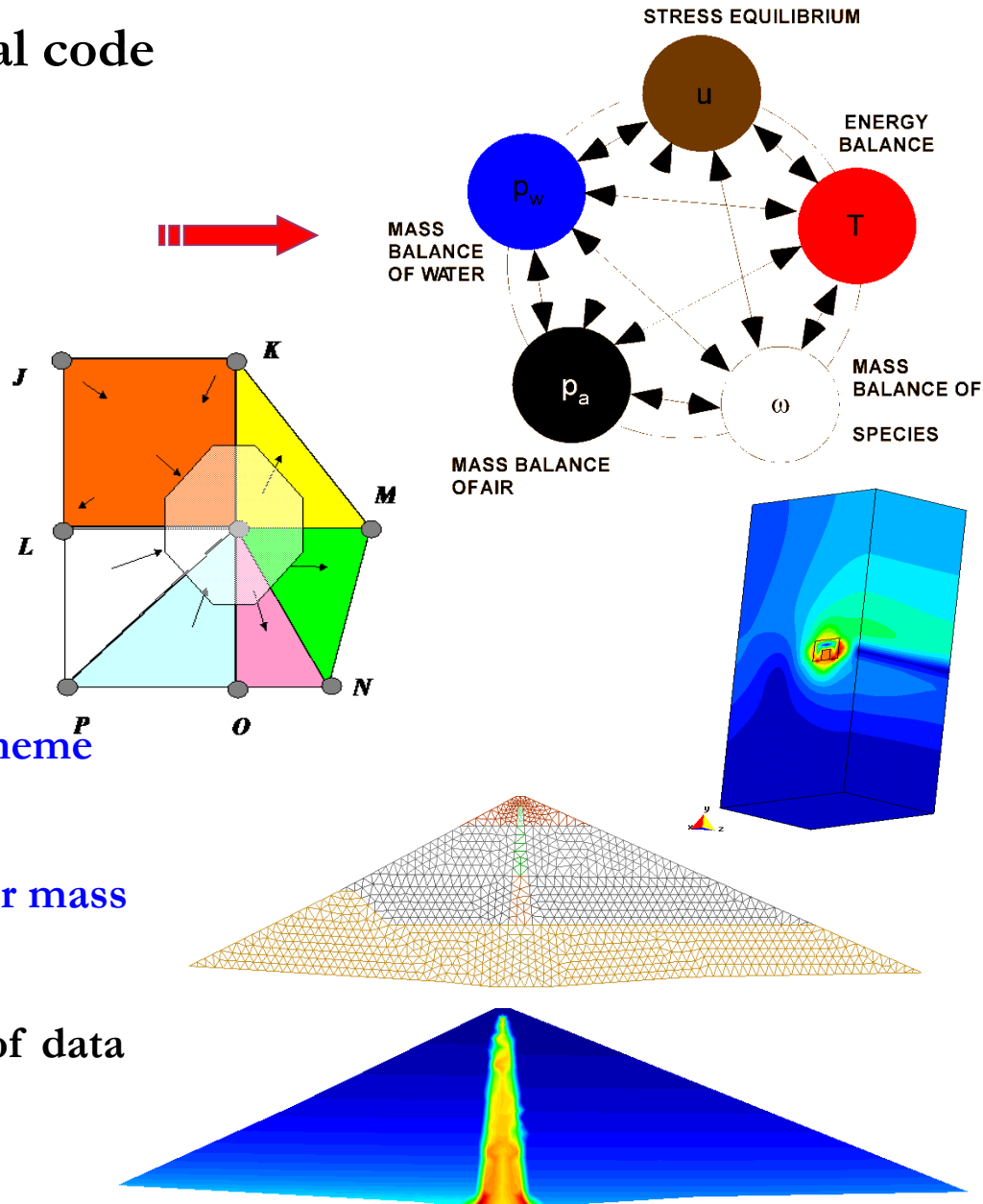
## ➤ CODE\_BRIGHT computational code

Olivella et al. (1996)

Coupled analysis in geological media

Interaction with the atmosphere

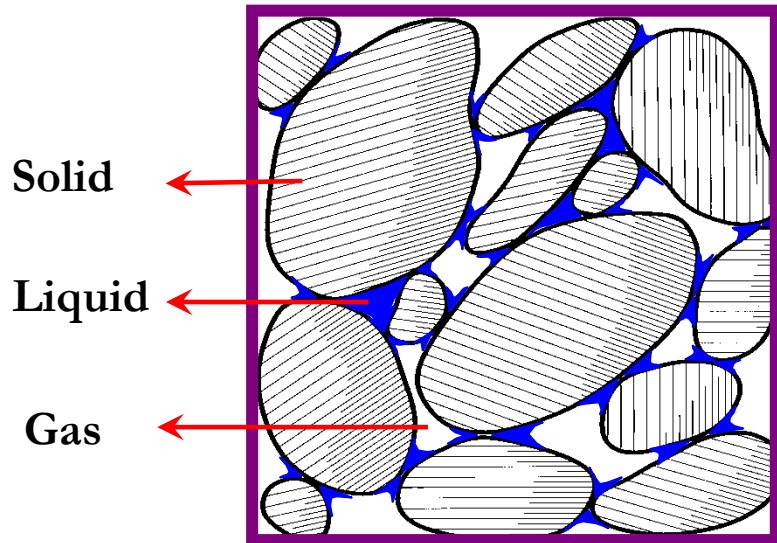
- ✓ Finite element in space
  - 1D, 2D and 3D elements
  - Monolithic coupling
  - Full Newton-Raphson
- ✓ Finite difference in time
  - Implicit time discretisation scheme
  - Automatic time advance
  - Mass conservative approach for mass balance equations
- ✓ User-friendly pre/post processing of data





# Coupled THM Formulation

## ➤ Phases and species



### Three phases:

- **solid** (*s*) : mineral
- **liquid** (*l*) : water + air dissolved
- **gas** (*g*) : mixture of dry air and water vapour

### Three species:

- **mineral** (-) : the mineral is coincident with solid
- **water** (*w*) : as liquid or evaporated in the gas phase
- **air** (*a*) : dry air, as gas or dissolved in the liquid phase

## ➤ Governing equations

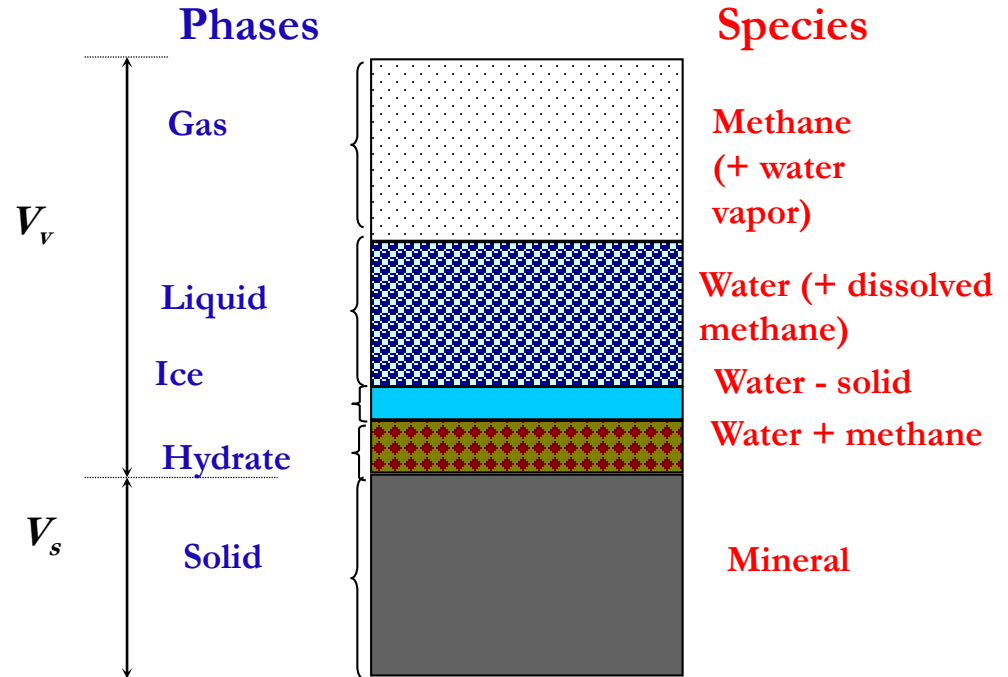
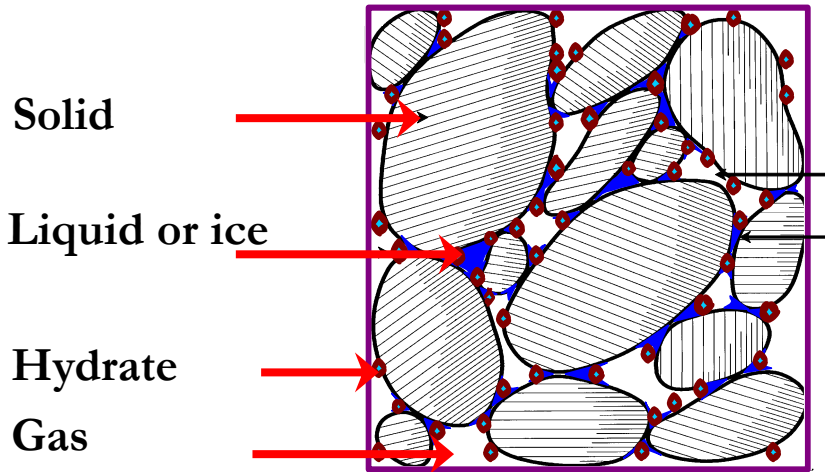
$$\phi = \frac{V_{pores}}{V_{total}} = \frac{V_{liquid} + V_{gas}}{V_{total}}$$

✓ Balance equations

$$S_l = \frac{V_{liquid}}{V_{pores}} = \frac{V_{liquid}}{V_{liquid} + V_{gas}} = 1 - S_g$$

✓ Equilibrium restrictions

# HBS - Species and Phases



Three main **species**:

- ✓ mineral
- ✓ water (w)
- ✓ methane (m)

The **species** are found in five **phases**:

- ✓ solid particles (s)
- ✓ liquid (l)
- ✓ gas (g)
- ✓ hydrates (h)
- ✓ ice (i)

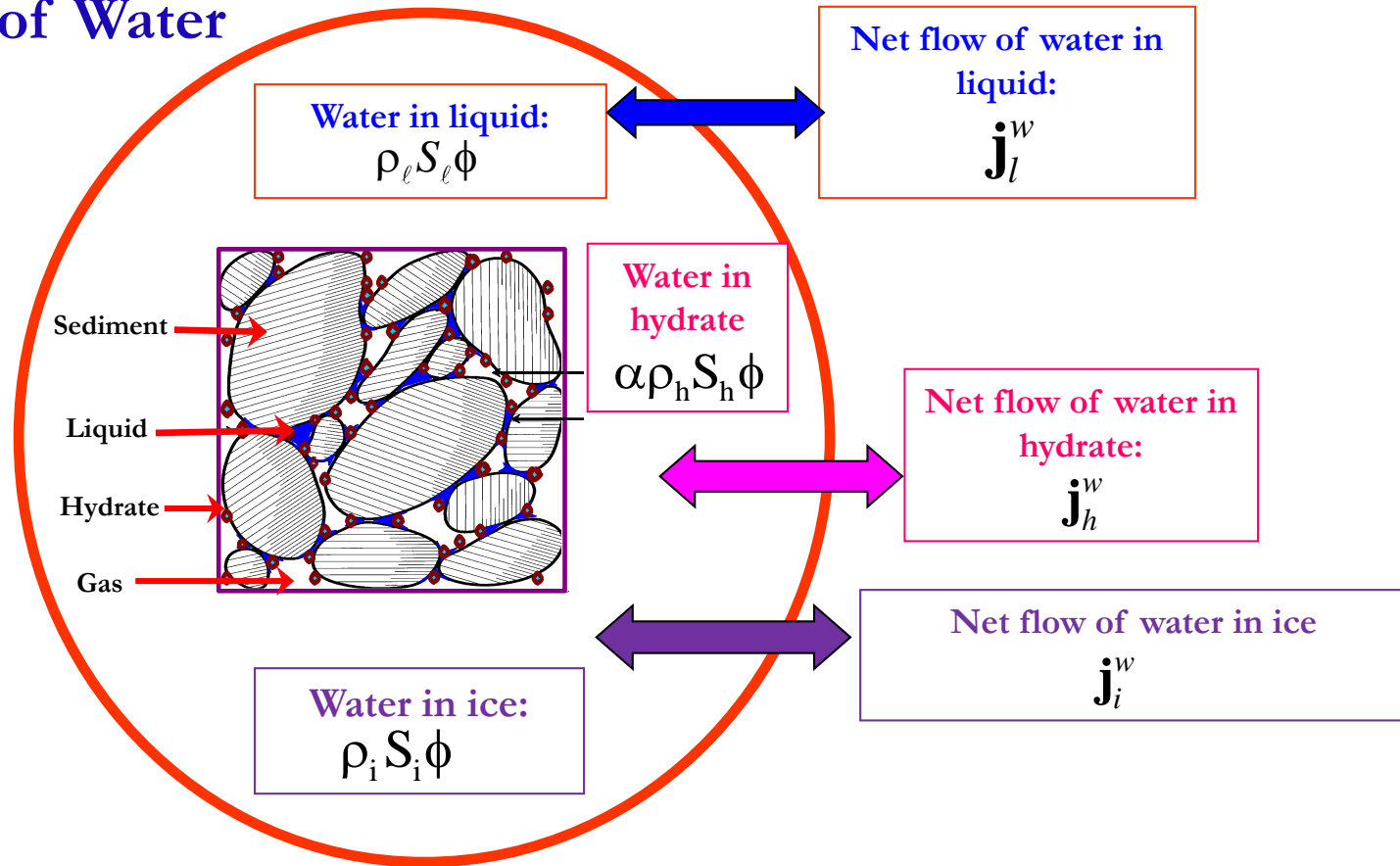
$$\phi = \frac{V_v}{V} = \frac{V_g + V_l + V_b + V_i}{V}$$

$$S_\beta = \frac{V_\beta}{V_v}$$

$$\beta = l, g, b, i \quad S_l + S_g + S_b + S_i = 1$$

# HBS - Mass Balance Equations

## Mass Balance of Water



$$\frac{\partial}{\partial t} \left( \text{mass of water in liquid, hydrate and ice (& gas) phases} \right) + \text{divergence} \left( \text{total flows of water} \right) = \left( \text{external supply of water} \right)$$

$$\frac{\partial}{\partial t} \left[ \underbrace{(\rho_\ell S_\ell + \alpha \rho_h S_h + \rho_i S_i)}_{\text{mass of water in liquid, hydrate and ice phases}} \phi \right] + \nabla \cdot \left[ \underbrace{\mathbf{j}_\ell^w + \mathbf{j}_h^w + \mathbf{j}_i^w}_{\text{total flows of water}} \right] = \mathbf{f}^w$$

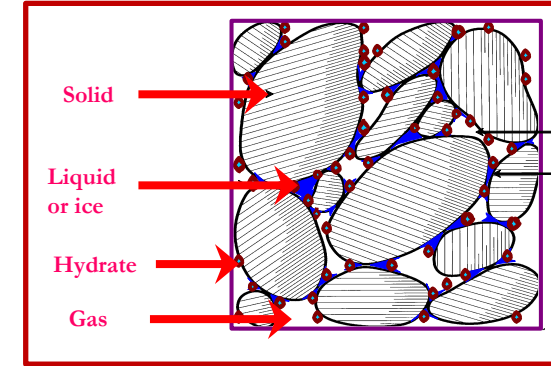
external supply of water

# HBS – Balance Equations

## Mathematical Formulation

### Mass of water ( $P_l$ )

$$\frac{\partial}{\partial t} \underbrace{[(\rho_\ell S_\ell + \alpha \rho_h S_h + \rho_i S_i) \phi]}_{\text{mass water per unit volume}} + \nabla \cdot \underbrace{[\rho_\ell \mathbf{q}_\ell + \rho_\ell S_\ell \phi \mathbf{v}]}_{\text{w in liquid}} + \underbrace{\alpha \rho_h S_h \phi \mathbf{v}}_{\text{w in hydrate}} + \underbrace{\rho_i S_i \phi \mathbf{v}}_{\text{w in ice}} = f^w$$



### Mass of methane ( $P_g$ )

$$\frac{\partial}{\partial t} \underbrace{[\rho_g S_g + (1-\alpha) \rho_h S_h]}_{\text{mass of methane per unit volume}} \phi + \nabla \cdot \underbrace{[\rho_g \mathbf{q}_g + \rho_g S_g \phi \mathbf{v}]}_{\text{m in gas}} + \underbrace{(1-\alpha) \rho_h S_h \phi \mathbf{v}}_{\text{m in hydrate}} = f^m$$

### Mass of solid ( $\phi$ )

$$\frac{\partial}{\partial t} \underbrace{[\rho_s (1-\phi)]}_{\text{mass mineral per unit volume}} + \nabla \cdot \underbrace{[\rho_s (1-\phi) \mathbf{v}]}_{\text{m in solid}} = 0$$

### Mass of solute ( $c_i$ )

$$\frac{\partial}{\partial t} \underbrace{(C_s S_\ell \rho_\ell \phi)}_{\text{mass s per unit volume}} + \nabla \cdot \underbrace{[\mathbf{D} \rho_\ell \nabla C_s]}_{\text{non advective flux of s}} + \underbrace{C_s \rho_\ell \mathbf{q}_\ell}_{\text{s in liquid}} + \underbrace{C_s \rho_i S_i \phi \mathbf{v}}_{\text{s in liquid}} = f^s$$

### Energy ( $T$ )

$$\frac{\partial}{\partial t} \underbrace{[e_s \rho_s (1-\phi)]}_{\text{energy per unit volume of the hydrate bearing sediment}} + \underbrace{(e_\ell \rho_\ell S_\ell + e_g \rho_g S_g + e_h \rho_h S_h + e_i \rho_i S_i) \phi}_{\text{energy per unit volume of the hydrate bearing sediment}} + \nabla \cdot \mathbf{i}_c + \nabla \cdot \underbrace{[e_\ell \rho_\ell (\mathbf{q}_\ell + S_\ell \phi \mathbf{v})]}_{\text{transport in } \ell} + \underbrace{e_g \rho_g (\mathbf{q}_g + S_g \phi \mathbf{v})}_{\text{transport in g}} + \underbrace{e_h \rho_h S_h \phi \mathbf{v}}_{\text{transport in h}} + \underbrace{e_i \rho_i S_i \phi \mathbf{v}}_{\text{transport in i}} + \underbrace{e_s \rho_s (1-\phi) \mathbf{v}}_{\text{transport in s}} = f^E$$

### Momentum ( $u$ )

$$\nabla \cdot \boldsymbol{\sigma} + \mathbf{b} = \mathbf{0}$$

# HBS - Constitutive Equations and Equilibrium Restrictions

EQUATION	VARIABLE NAME	VARIABLE
<b>Constitutive Equations</b>		
Fourier's law	conductive heat flux	$i_c$
Darcy's law	liquid and gas advective flux	$q_l, q_g$
Retention curve	liquid degree of saturation	$S_l, S_g$
Fick's law	vapour and air non-advective fluxes	$i_g^w, i_l^m$
Mechanical model	stress tensor	$\sigma$
Phase density	liquid density	$\rho_l$
Gases law	methane density	$\rho_g$
<b>Equilibrium Restrictions</b>		
Hydrate dissociation/formation	Hydrate Saturation	$S_h$
Ice thaw/formation	Ice Saturation	$S_i$
Henry's law	Methane dissolved mass fraction	$\omega_l^a$
Psychrometric law		



# HBS - Phases Properties and Phase Change

Phases	Mass Density $\rho$ [kg/m <sup>3</sup> ]	Specific Energy			Thermal Conductivity $\lambda$ [W/(m.K)]
		Expression	Latent heat L [J/g]	Specific heat c [J/(g K)]	
<i>liquid (water)</i>	998	$e_\ell = c_\ell (T - T_o)$	-	4.2	0.58
<i>ice</i>	917	$e_i = -L_{\text{fuse}} + c_i (T - T_o)$	334 <i>fusion</i>	2.1	2.3
<i>gas (methane)</i>	gas law (see text)	$e_g = c_g (T - T_o)$	-	1.9 V=const 2.5 P=const	0.05 (P-dependent <sup>a</sup> )
<i>hydrate</i>	910	$e_h = L_{\text{diss}} + c_h (T - T_o)$	339 <i>dissociation</i>	2.1	0.6 (T-dependent <sup>b</sup> )
<i>mineral</i>	2650	$e_s = c_s (T - T_o)$	-	0.7 <i>quartz</i>	8 <i>quartz</i>
			-	0.8 <i>calcite</i>	3 <i>calcite</i>

$$\mu_\ell [\text{Pa.s}] = 2.1 \cdot 10^{-6} \exp\left(\frac{1808.5 \text{ }^\circ\text{K}}{T}\right)$$

$$\mu_g [\text{Pa.s}] = 10.3 \cdot 10^{-6} \left[ 1 + 0.053 \frac{P_g}{\text{MPa}} \left( \frac{280 \text{ }^\circ\text{K}}{T} \right)^3 \right]$$

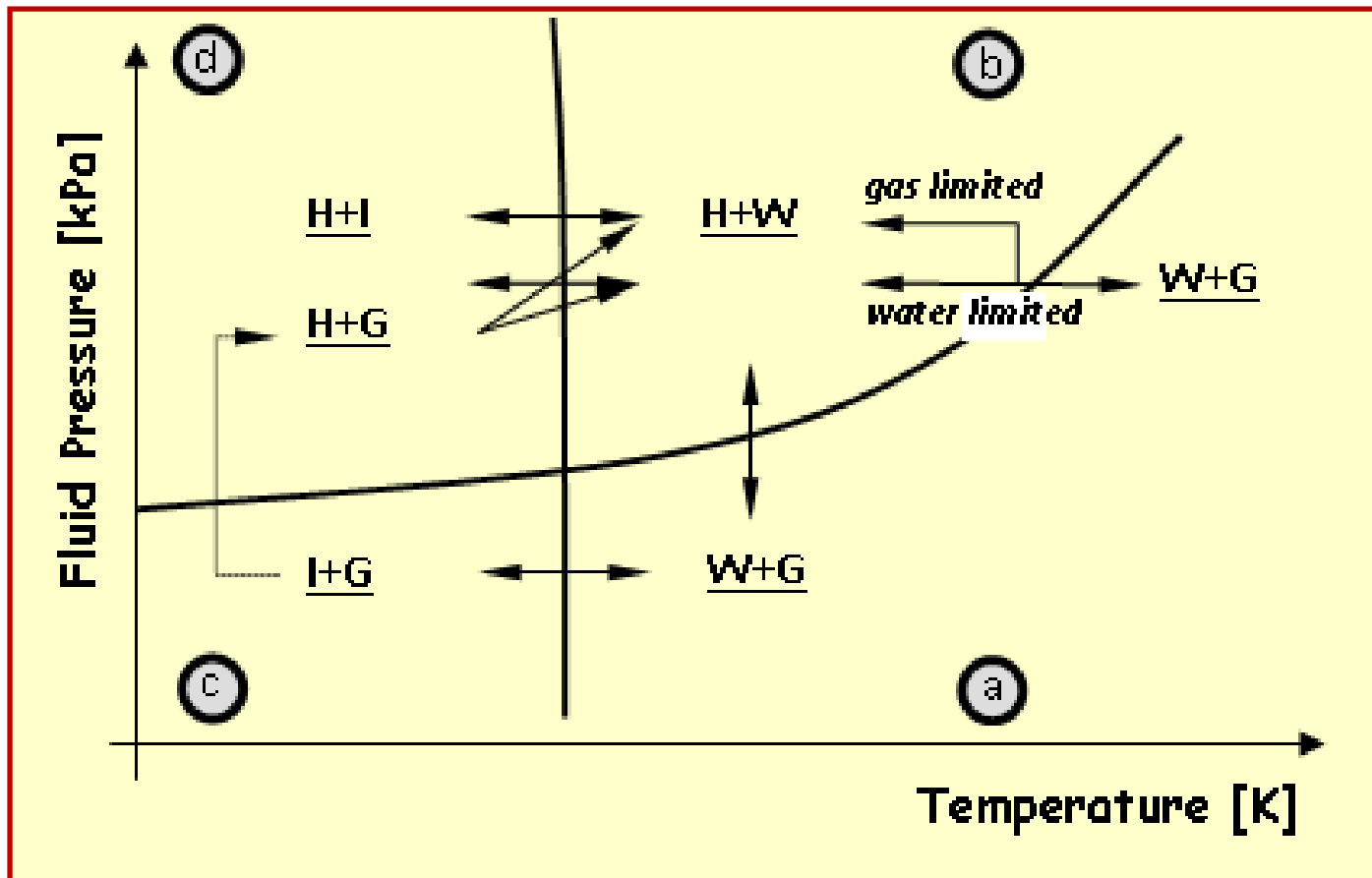
$$\rho_\ell = \rho_{\ell o} \left( 1 + \frac{P_\ell}{B_\ell} \right) \left[ 1 - \beta_{T_\ell} \left( \frac{T - 277 \text{ }^\circ\text{K}}{5.6} \right)^2 \right]$$

$$\lambda_{\text{hbs}} = \left[ (1 - \phi) \lambda_s^\beta + \phi (S_h \lambda_h^\beta + S_i \lambda_i^\beta + S_g \lambda_g^\beta + S_\ell \lambda_\ell^\beta) \right]^{1/\beta}$$

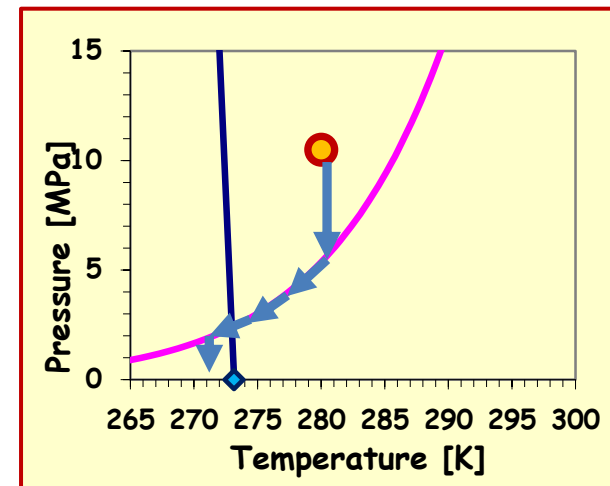
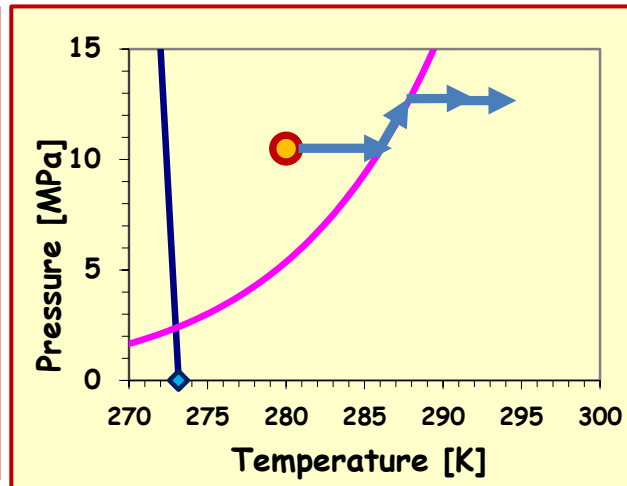
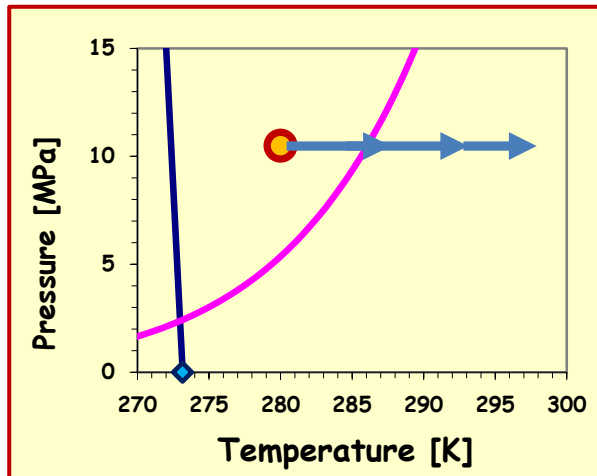
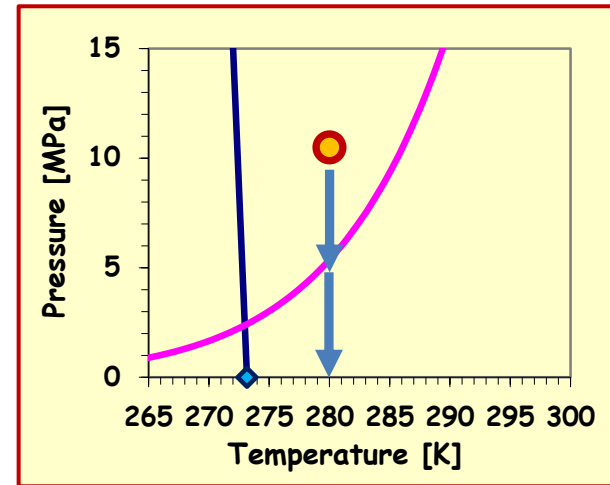
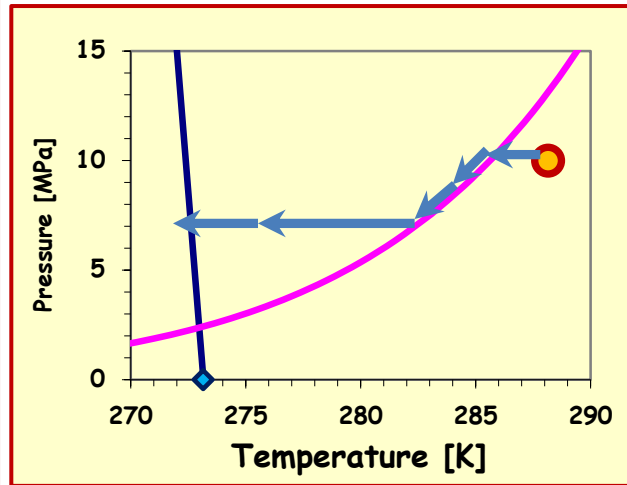
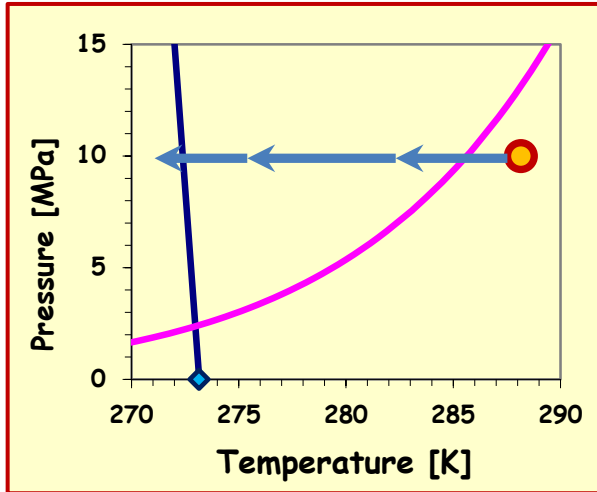
# HBS – Coupled Phenomena and Phase Boundaries

## ➤ PT Paths: Four Regions

- ✓ Phase boundaries for methane hydrate (H), gas (G), water (W) & ice (I).

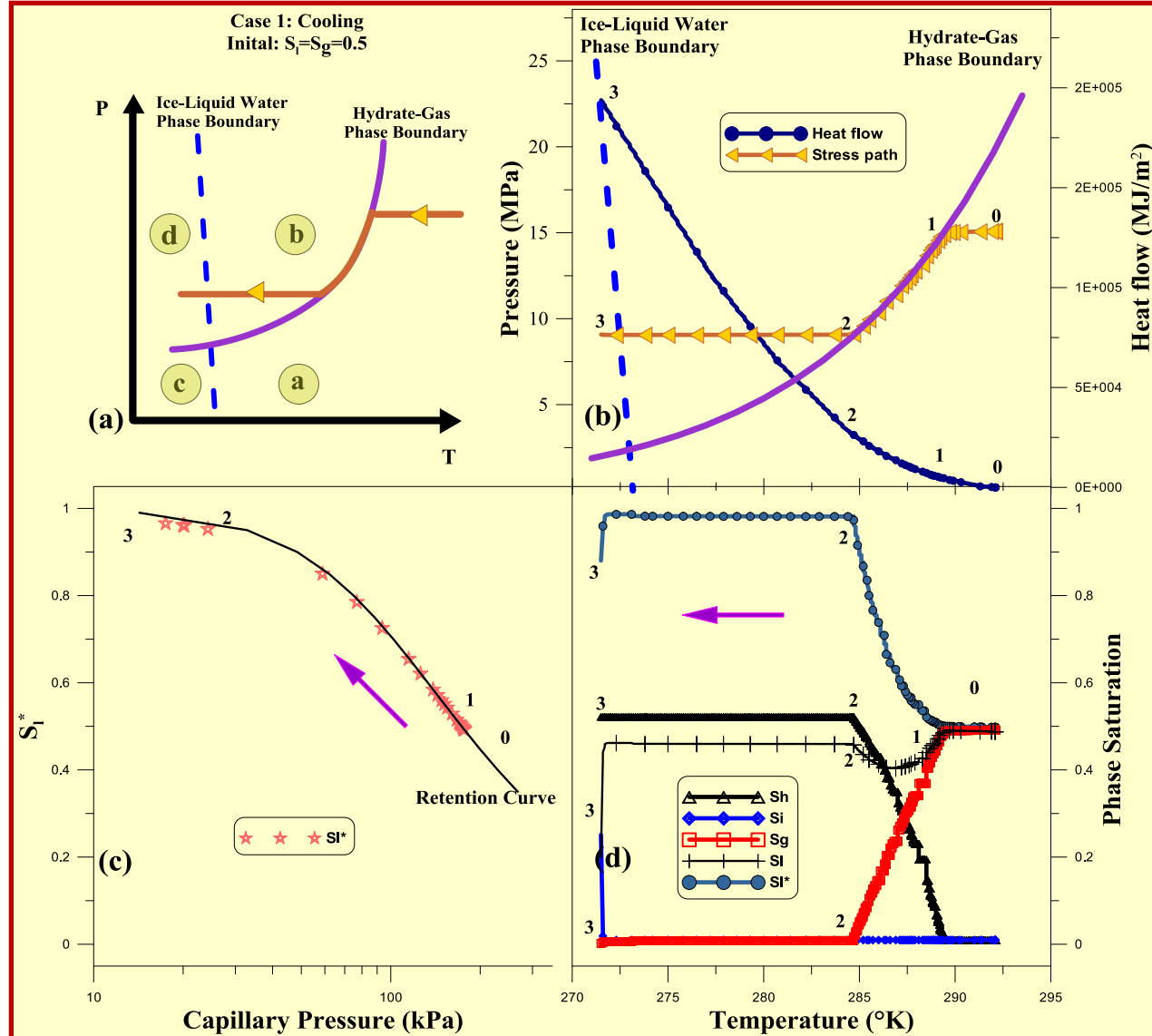


# Pressure-Temperature paths



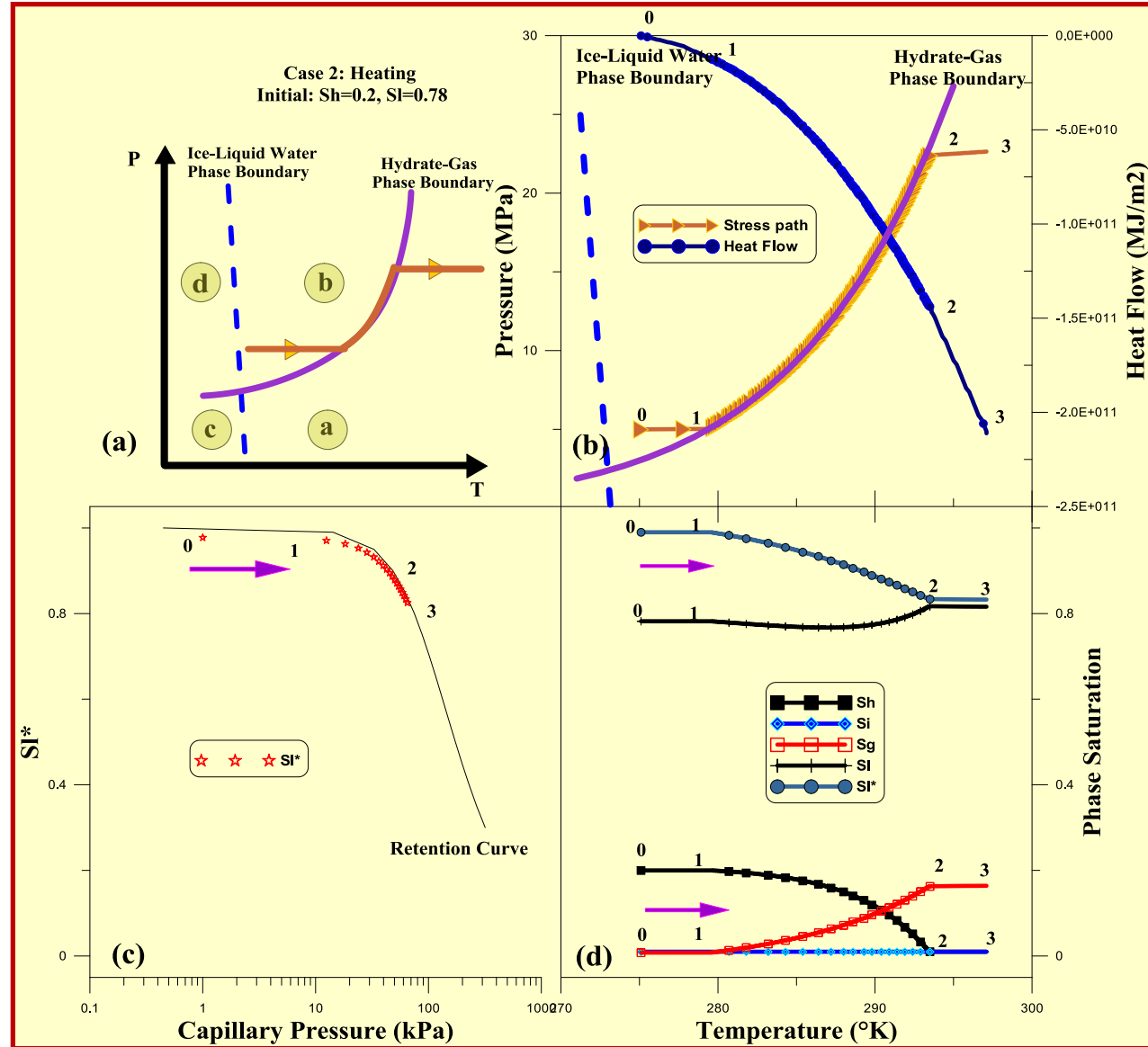
# Pressure-Temperature paths

## Hydrate formation



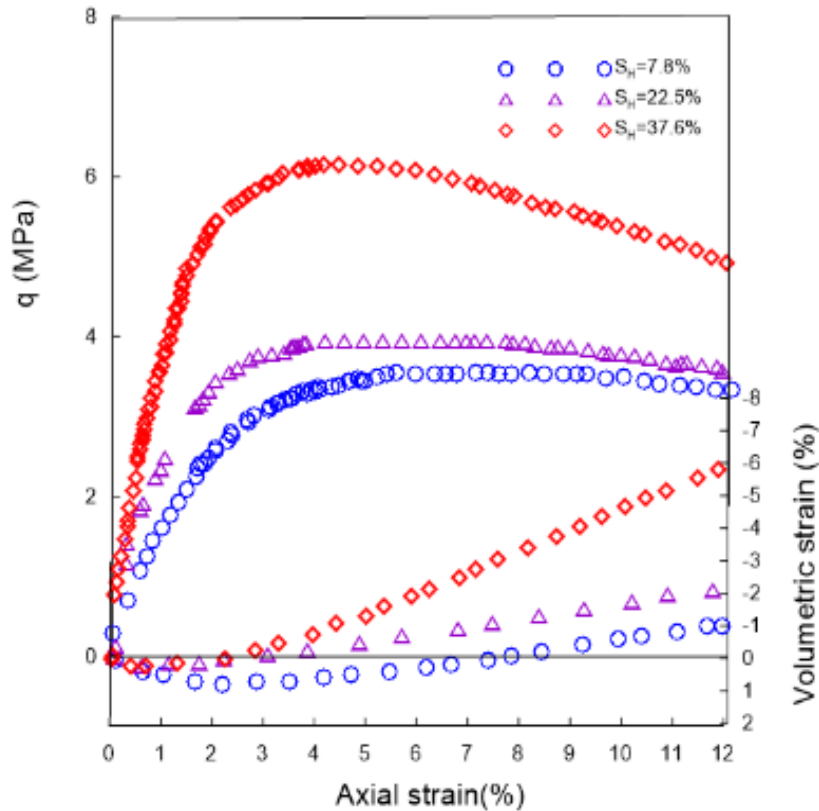
# Pressure-Temperature paths

## Hydrate Dissociation - Heating

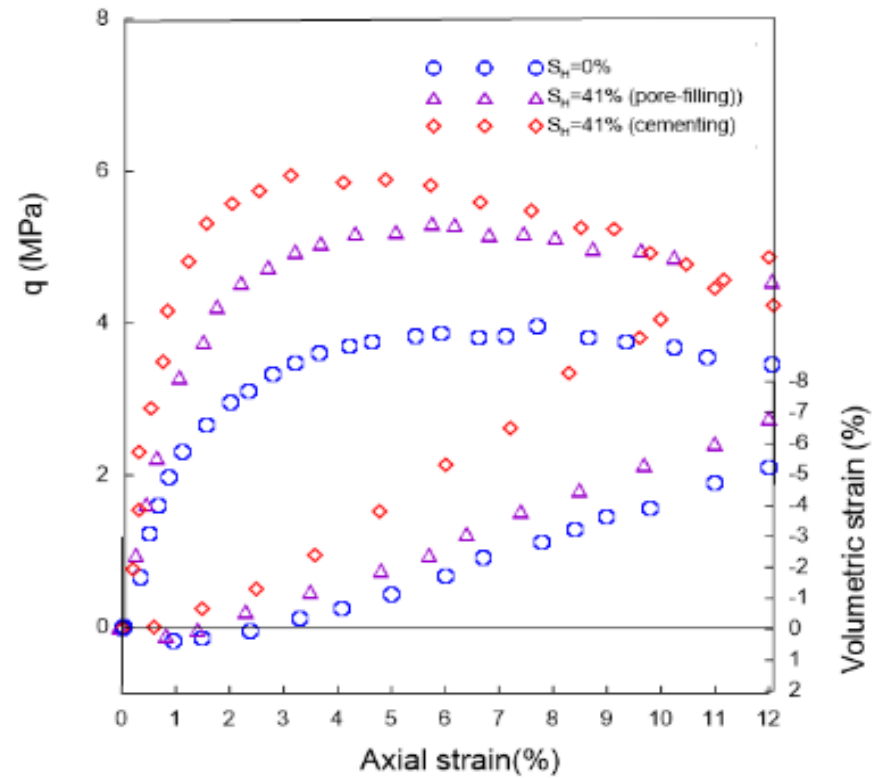




# HBS – Mechanical Behavior at Constant $S_h$

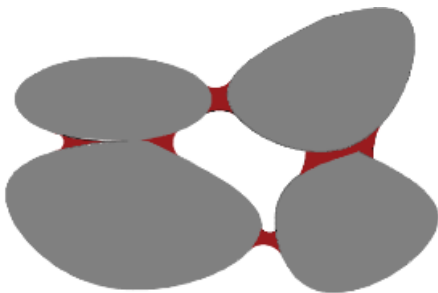


Masui et al. (2008)

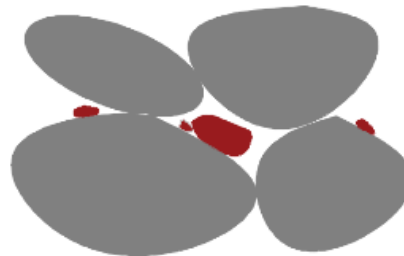


Masui et al. (2005)

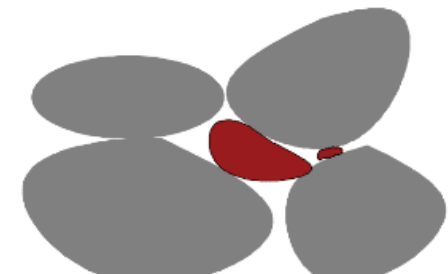
**Cementation**



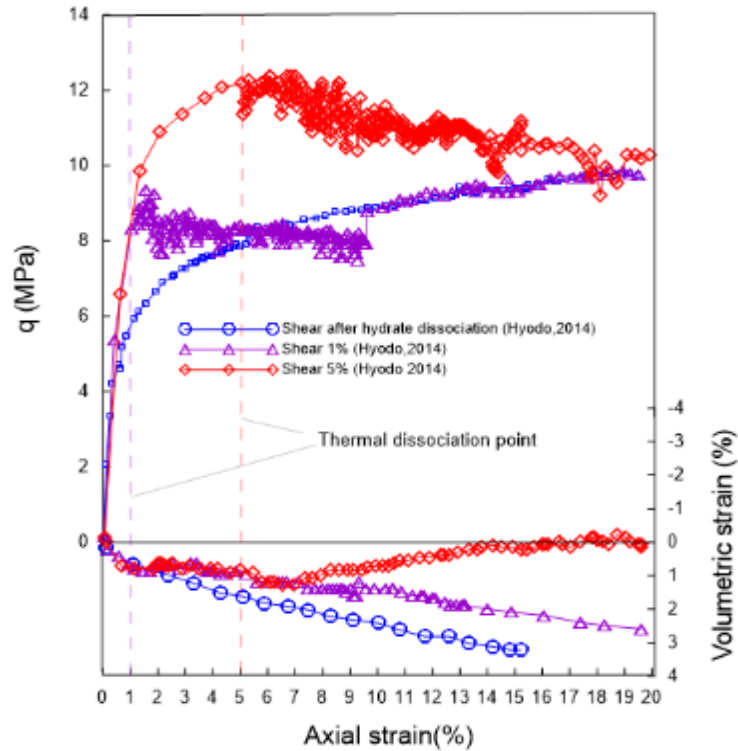
**Pore Filling**



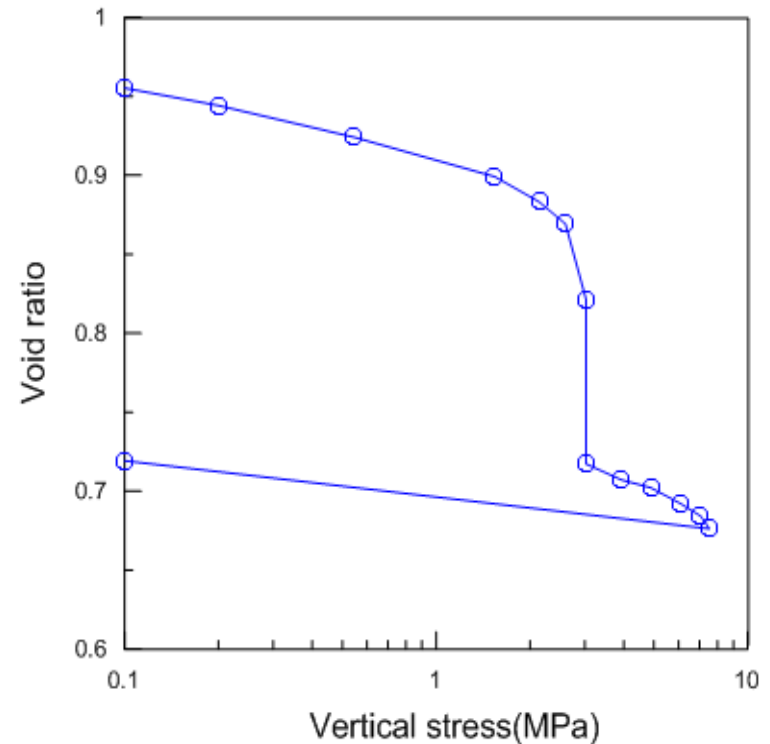
**Supporting Matrix**



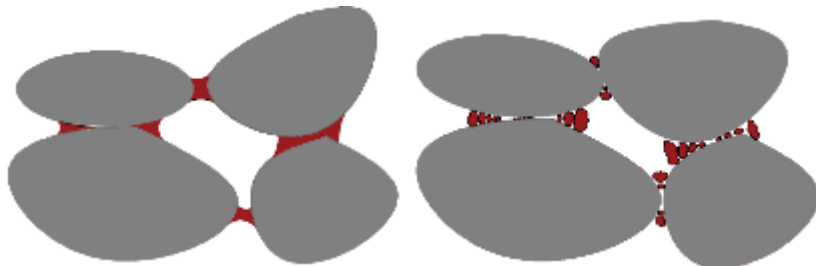
# HBS – Mechanical Behavior During Dissociation Under Stress



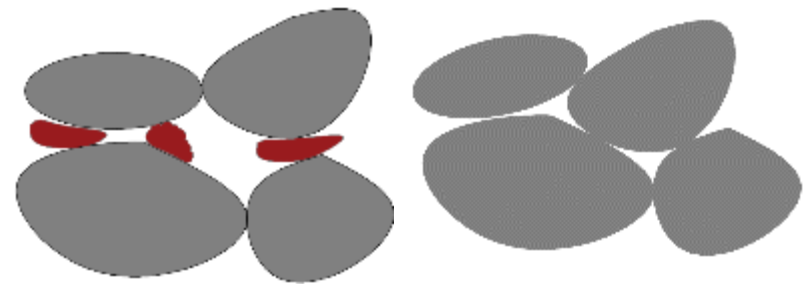
Hyodo et al. (2014)



Santamarina et al. (2015)



Shearing → Hydrate damage



Hydrate dissociation → Sediment collapse

# HBS – Mechanical Behavior

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## □ Some previous developments

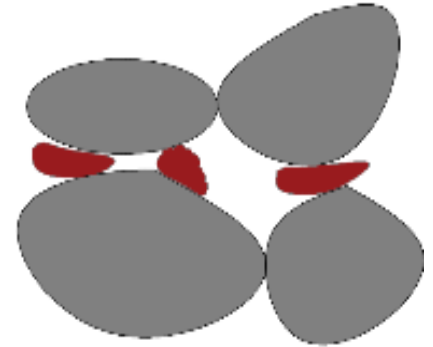
- Mohr–Coulomb based model
  - Rutqvist and Moridis (2007)
  - Klar, Soga and Ng (2010)
- Based on an elasto–viscoplastic framework
  - Kimoto, Oka, Fushita and Fujiwaki (2007).
  - Kimoto, Oka, Fushita (2010)
- Modified Cam-Clay based model
  - Sultan and Garziglia (2011)
  - Uchida, Soga and Yamamoto (2012)

# HBS – Mechanical Behavior

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## □ The mechanical behavior of HBS depends on

- Hydrate concentration
- Pore habit
- Stress level
- Stress history



## □ Hydrates in soils

- contribute to support the external applied stresses,
  - ✓ the strain partition concept is used to compute this contribution;
- alter the mechanical behavior of sediments, e.g. provide hardening and dilation enhancement
  - ✓ a critical state model for the sediment to account for these effects.

# HBS – Mechanical Model

## □ Strain partition concept

Proposed by Pinyol et al. (2007) for clayed cemented materials

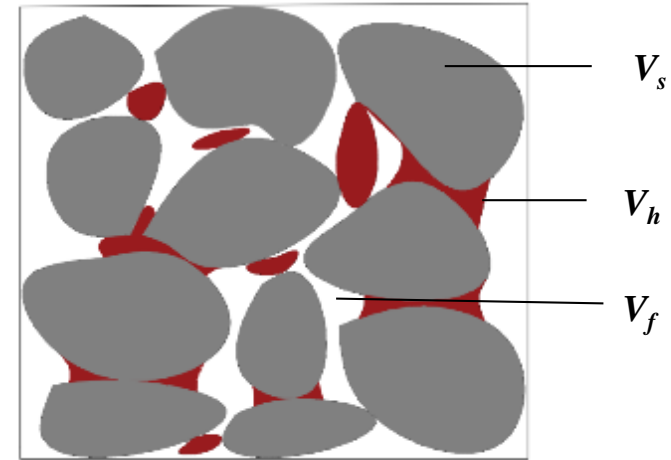
$$\boldsymbol{\varepsilon}^v = \boldsymbol{\varepsilon}_{ss}^v + C_h \boldsymbol{\varepsilon}_h^v$$

$$\boldsymbol{\varepsilon}^q = \boldsymbol{\varepsilon}_{ss}^q + C_h \boldsymbol{\varepsilon}_h^q$$

$$\boldsymbol{\varepsilon}_h^v = \chi \boldsymbol{\varepsilon}_{ss}^v$$

$$\boldsymbol{\varepsilon}_h^q = \chi \boldsymbol{\varepsilon}_{ss}^q$$

$$\boldsymbol{\varepsilon}_h = \frac{\chi}{1 + C_h \chi} \boldsymbol{\varepsilon}$$



## □ Hydrate Model

$$\boldsymbol{\sigma}_h = e^{-L} \mathbf{D}_{h0} \boldsymbol{\varepsilon}_h = \mathbf{D}_h \boldsymbol{\varepsilon}_h$$

$$C_h = \phi S_h$$

$$r_{(L)} = r_0 e^{r_1 L} = u_h$$

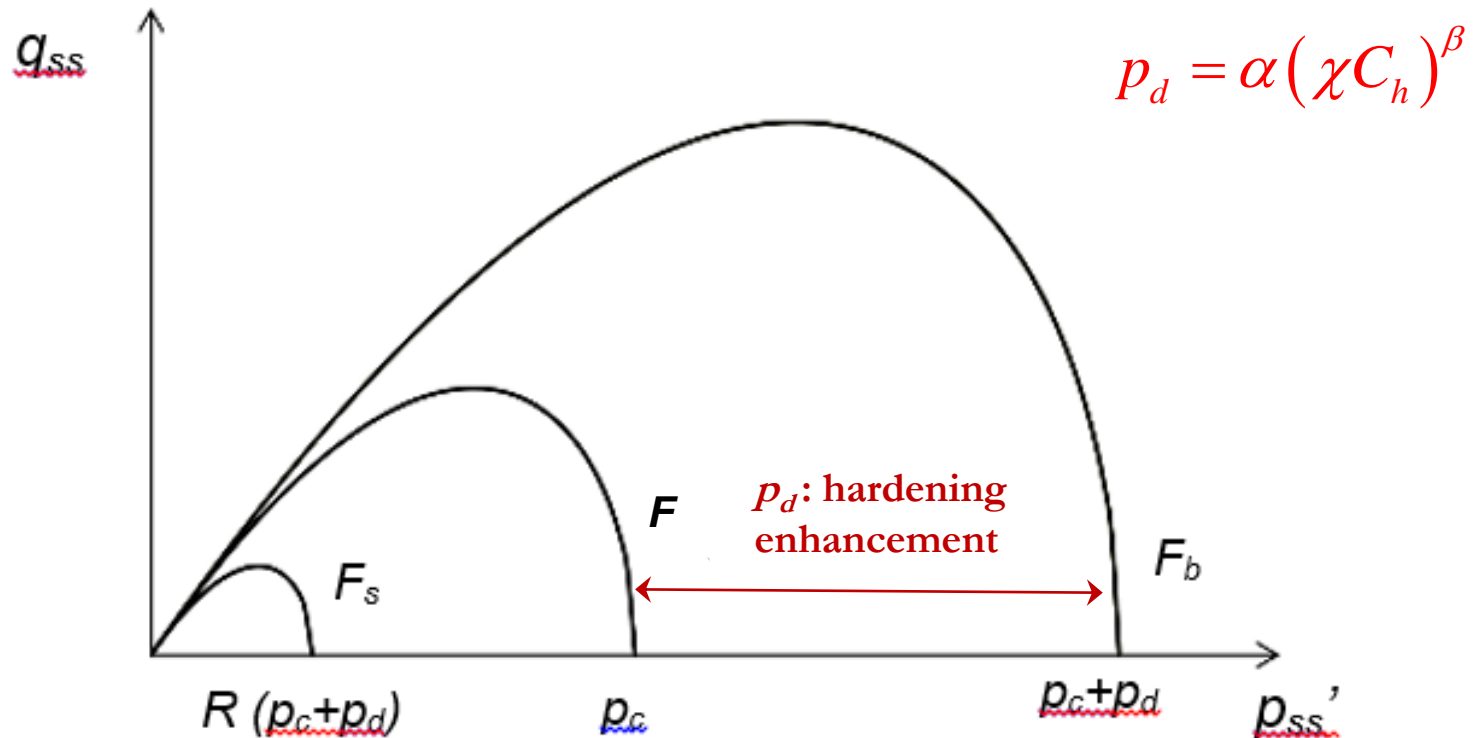
$$\chi = \chi_0 e^{-\frac{L}{2}}$$



# HBS – Mechanical Model

## □ Sediment Model

$$F_b = \frac{a}{M^2} q_{ss}^2 - 9\gamma \left\{ (p'_{ss})^2 - (p'_{ss})^n [(p_c + p_d)]^{2-n} \right\}$$



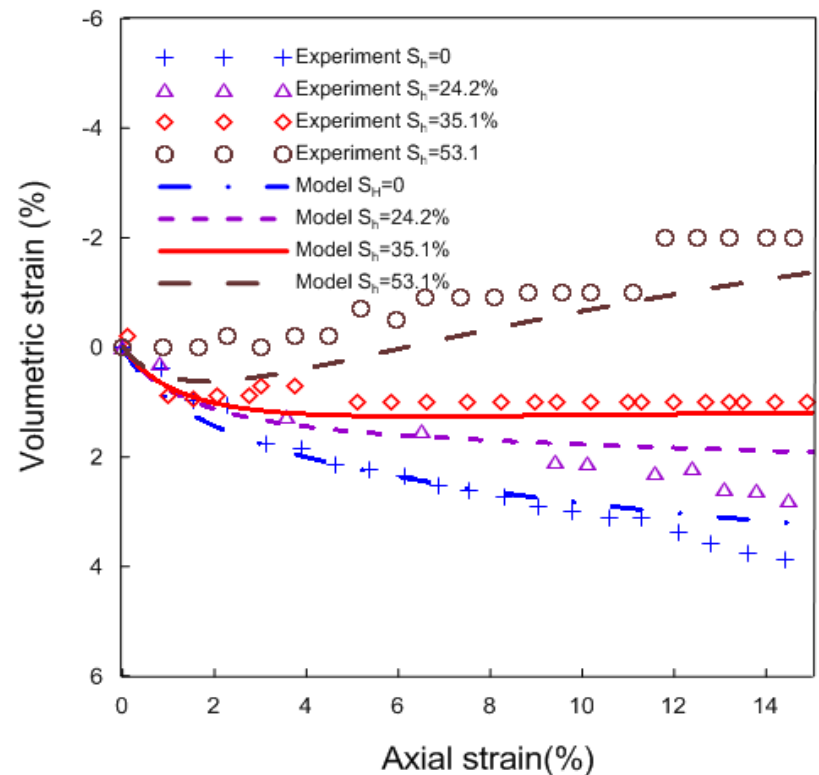
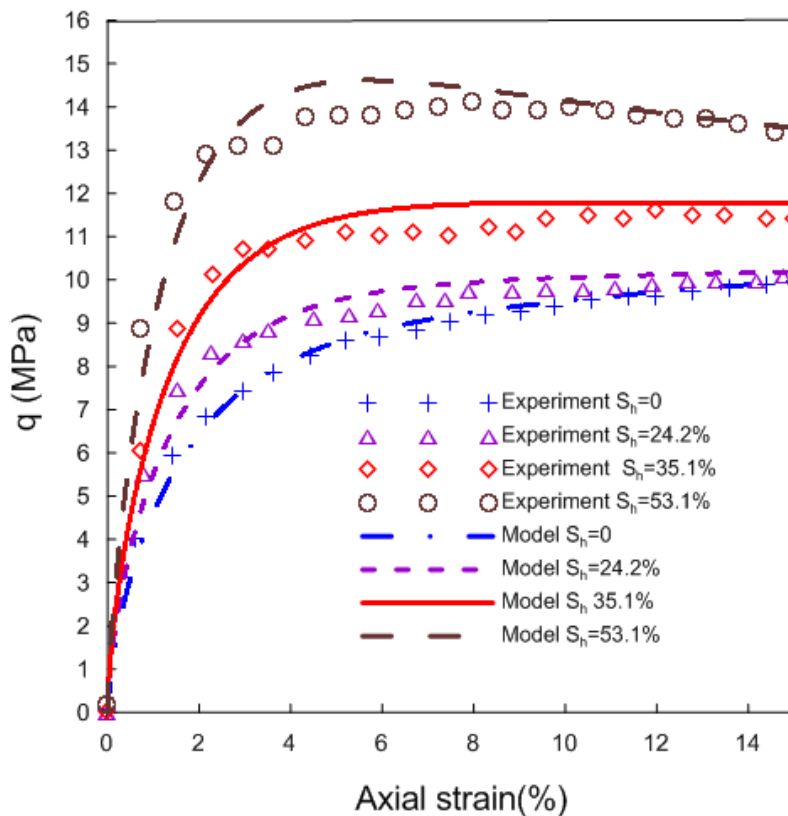
## □ Final Stress-Strain Relationships

$$d\boldsymbol{\sigma}' = d\boldsymbol{\sigma}'_{ss} + \frac{C_h \chi}{1 + C_h \chi} d\boldsymbol{\sigma}_h \quad d\boldsymbol{\sigma}' = \left[ \mathbf{D}_{ss} + \left( \frac{C_h \chi}{1 + C_h \chi} \right)^2 \mathbf{D}_h \right] d\boldsymbol{\varepsilon} + \left[ \mathbf{d}_{C_h} + \boldsymbol{\sigma}_h \left( \frac{C_h \chi}{1 + C_h \chi} - \left( \frac{C_h \chi}{1 + C_h \chi} \right)^2 \right) \right] dC_h$$

# HBS – Mechanical Model Validation

## Effect of Hydrate Saturation – Constant $S_h$

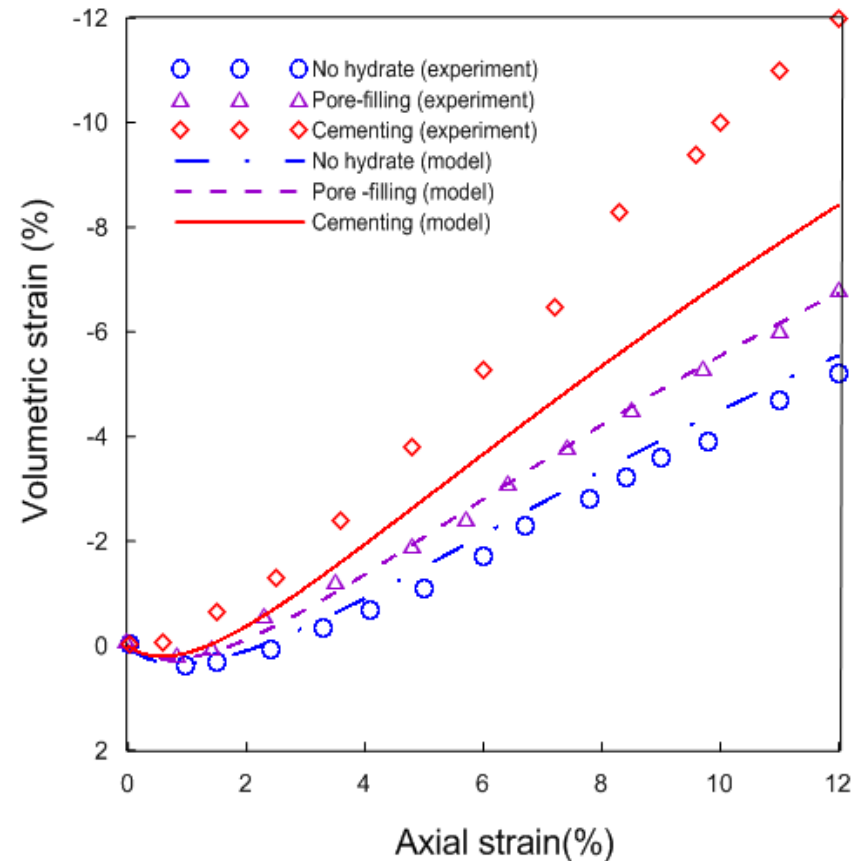
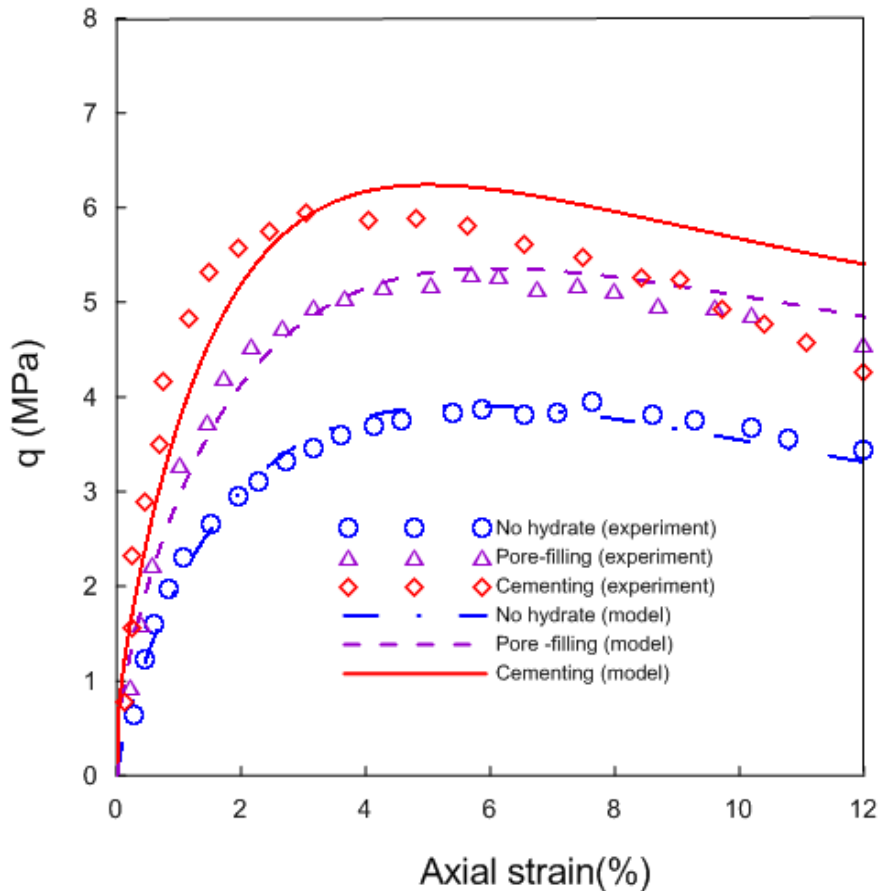
### Synthetic Samples – Triaxial Conditions



# HBS – Mechanical Model Validation

## Effect of Pore Habit – Constant $S_h$

### Synthetic Samples – Triaxial Conditions

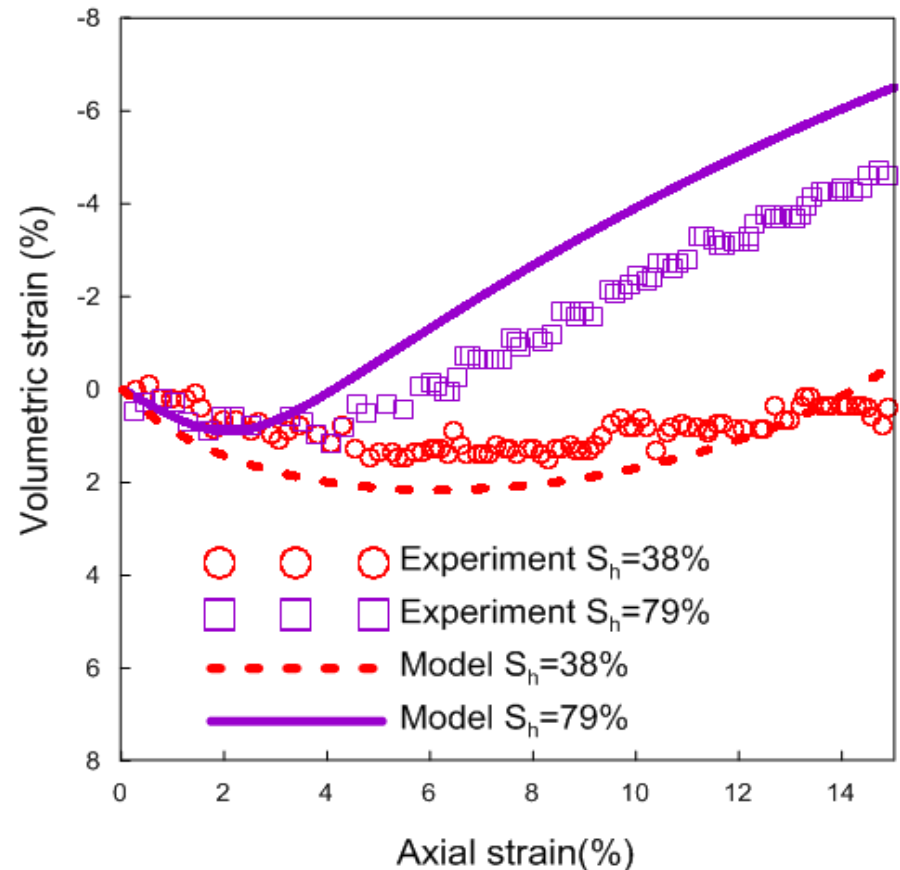
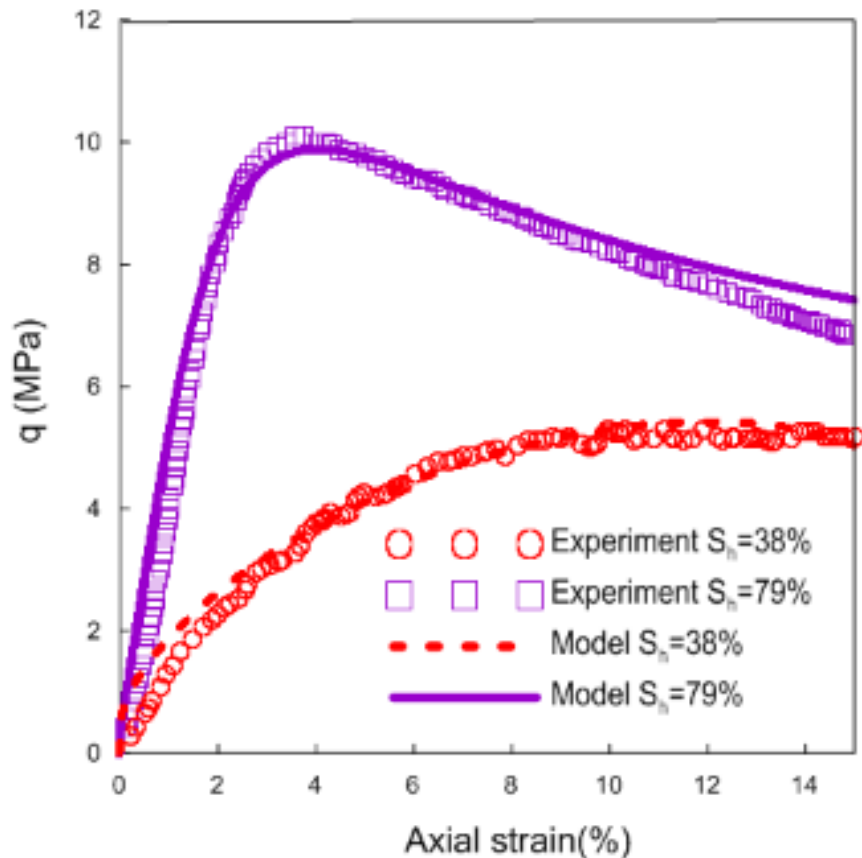


Experimental data from Masui et al. (2005)

# HBS – Mechanical Model Validation

## Effect of Hydrate Saturation – Constant $S_h$

### Natural Samples – Triaxial Conditions

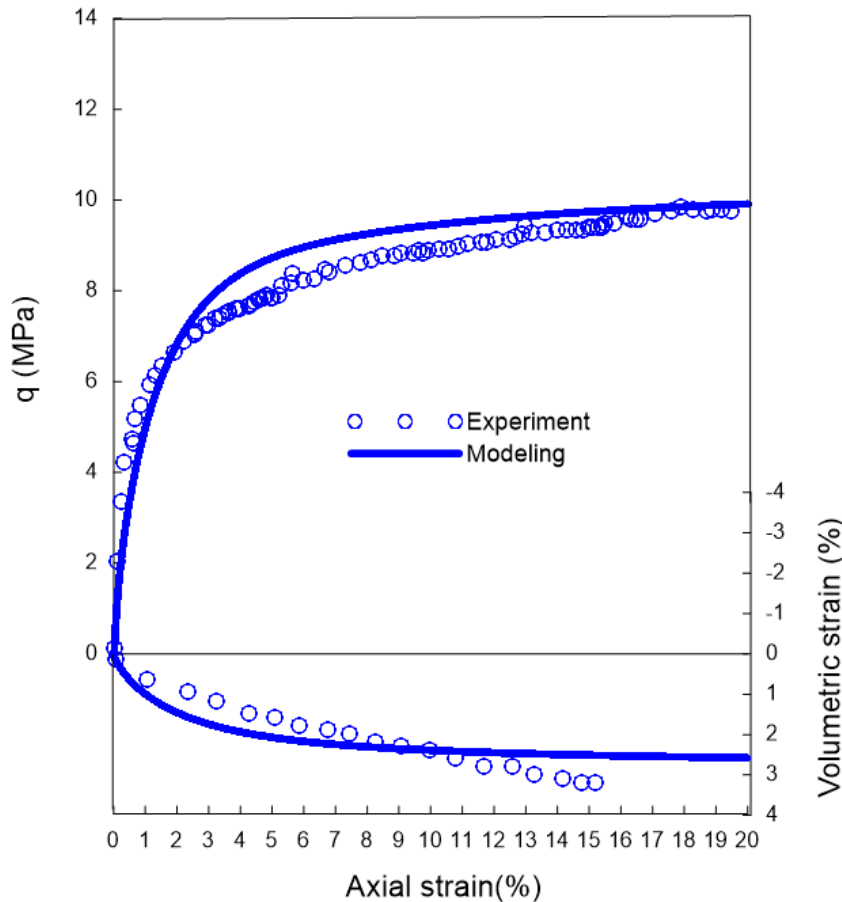


Experimental data from Joneda et al. (2015)

# HBS – Mechanical Model Validation

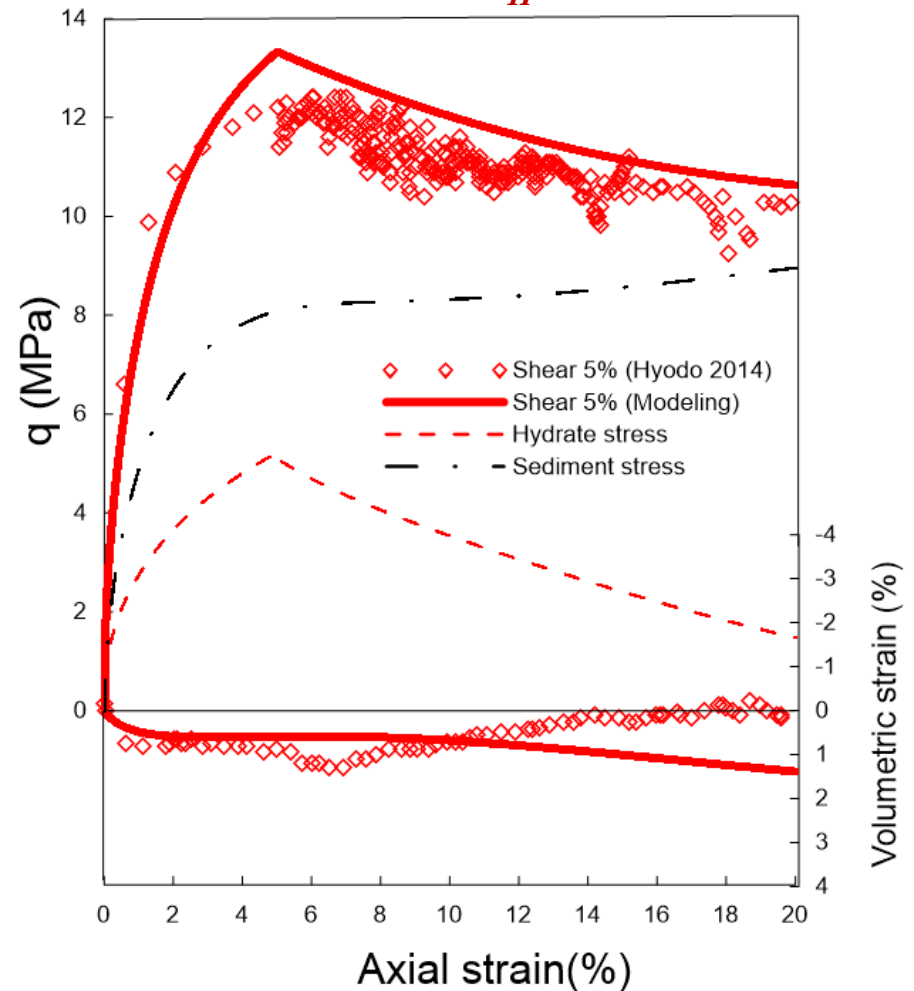
## Effect of Hydrate Dissociation

Synthetic Samples – Triaxial Conditions -  $S_h \sim 48\%$



Already dissociated sediment

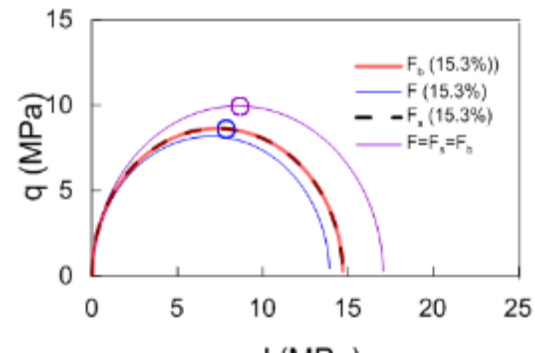
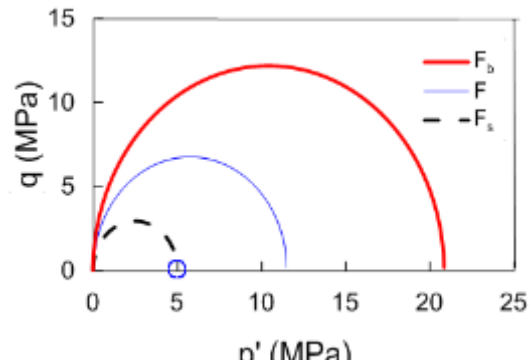
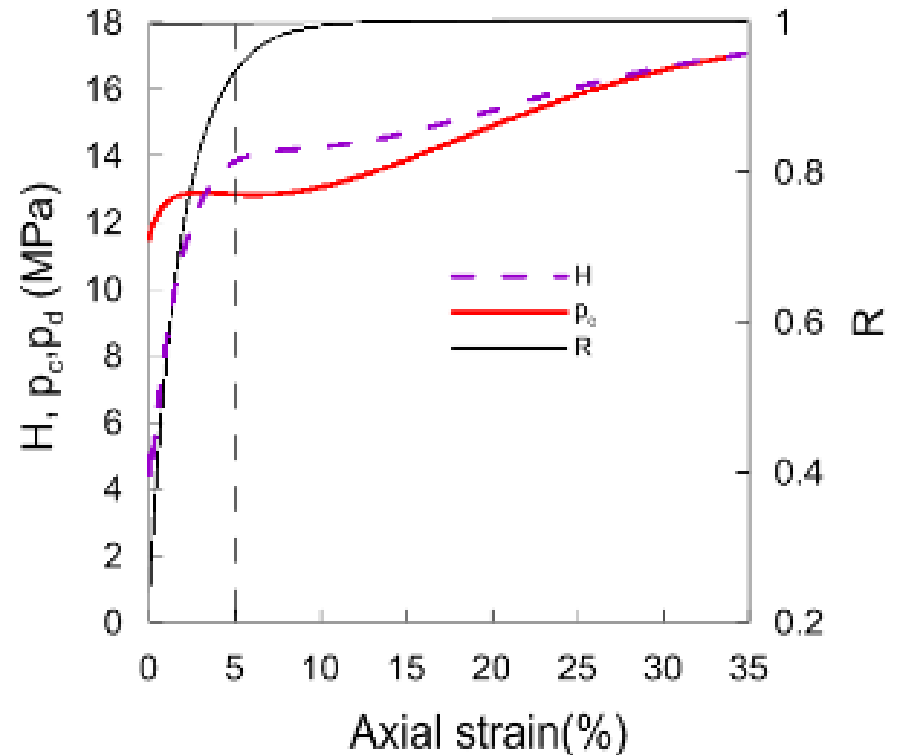
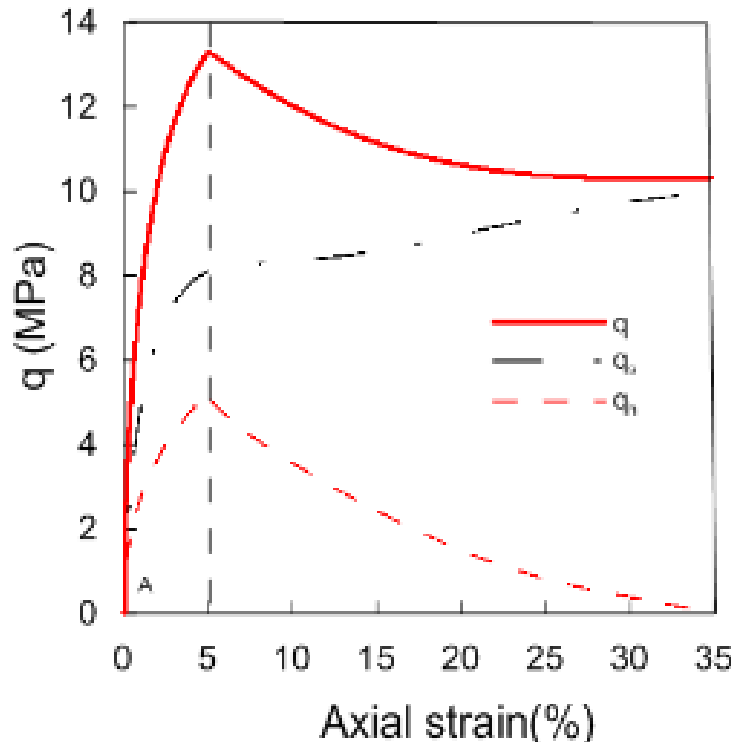
Experimental data from Hyodo et al. (2014)



Dissociation induced at  $\varepsilon_a = 5\%$ .

# HBS – Mechanical Model Validation

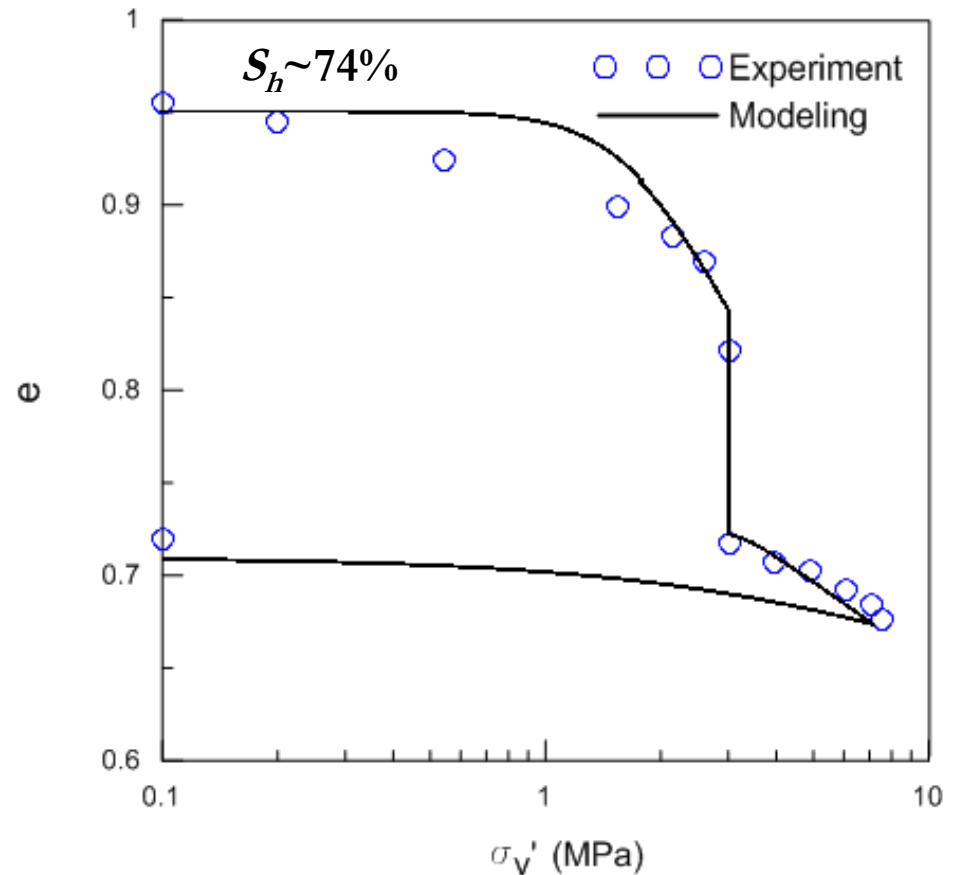
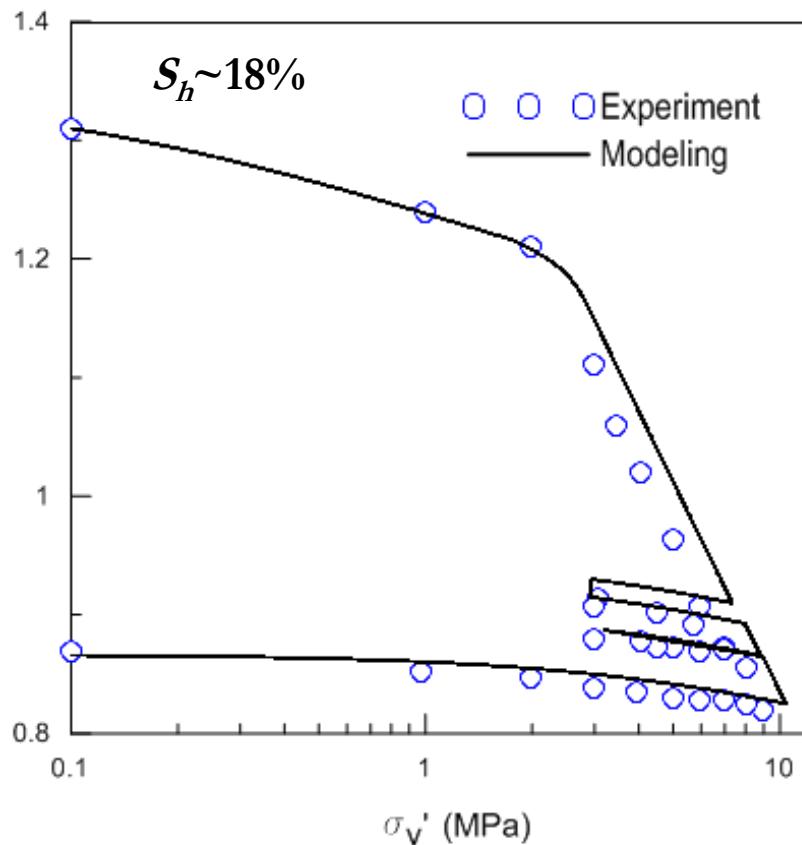
## Effect of Hydrate Dissociation



# HBS – Mechanical Model Validation

## Effect of Hydrate Dissociation

### Natural Samples – Oedometric Conditions

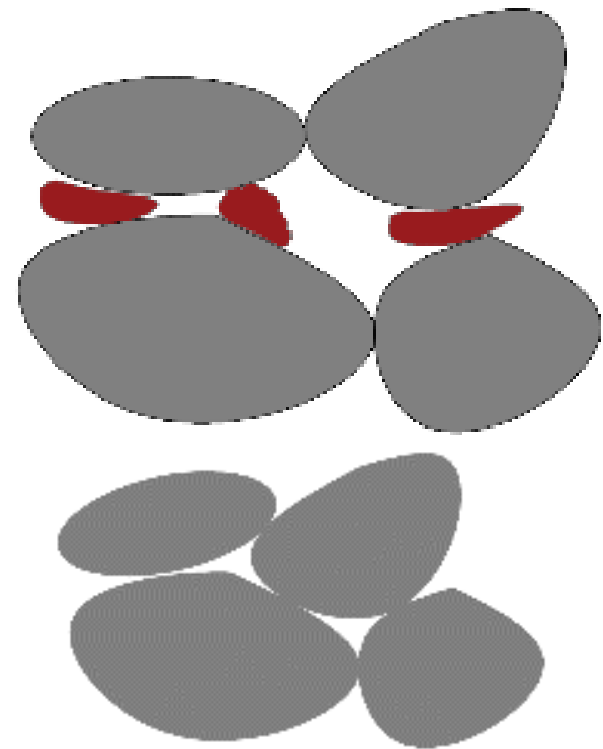
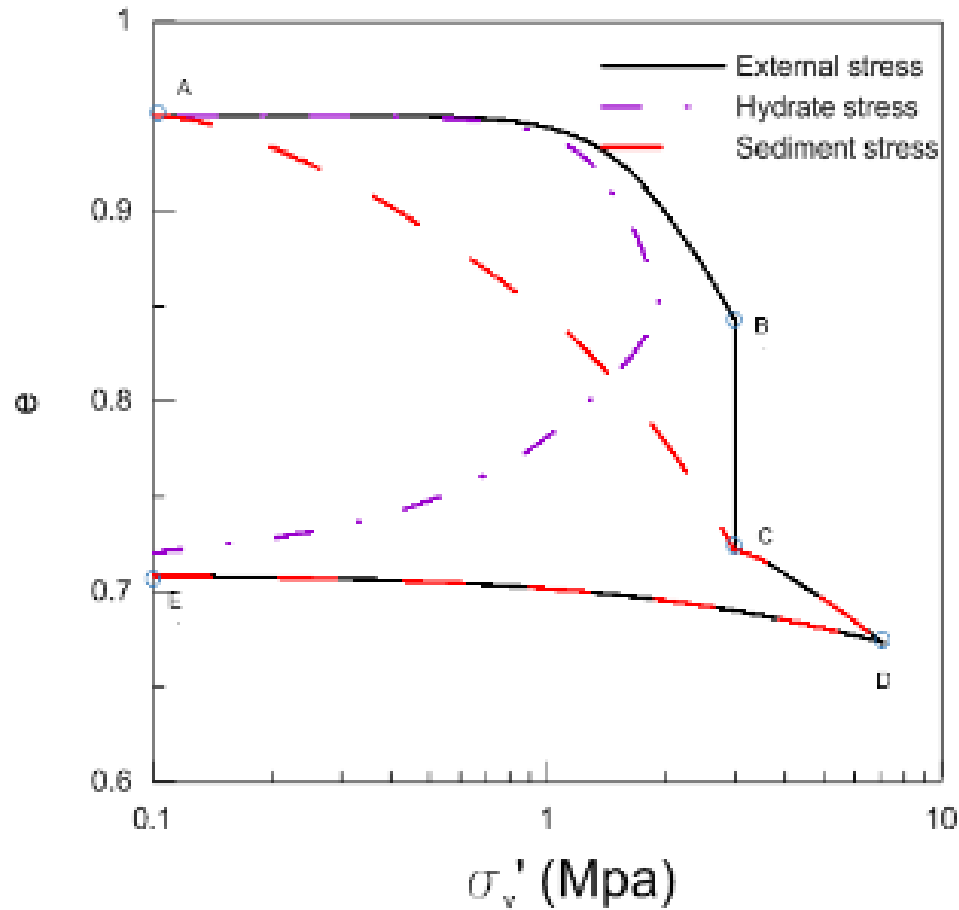


Experimental data from Santamarina et al. (2015)

# HBS – Mechanical Model Validation

## Effect of Hydrate Dissociation

### Oedometric Conditions





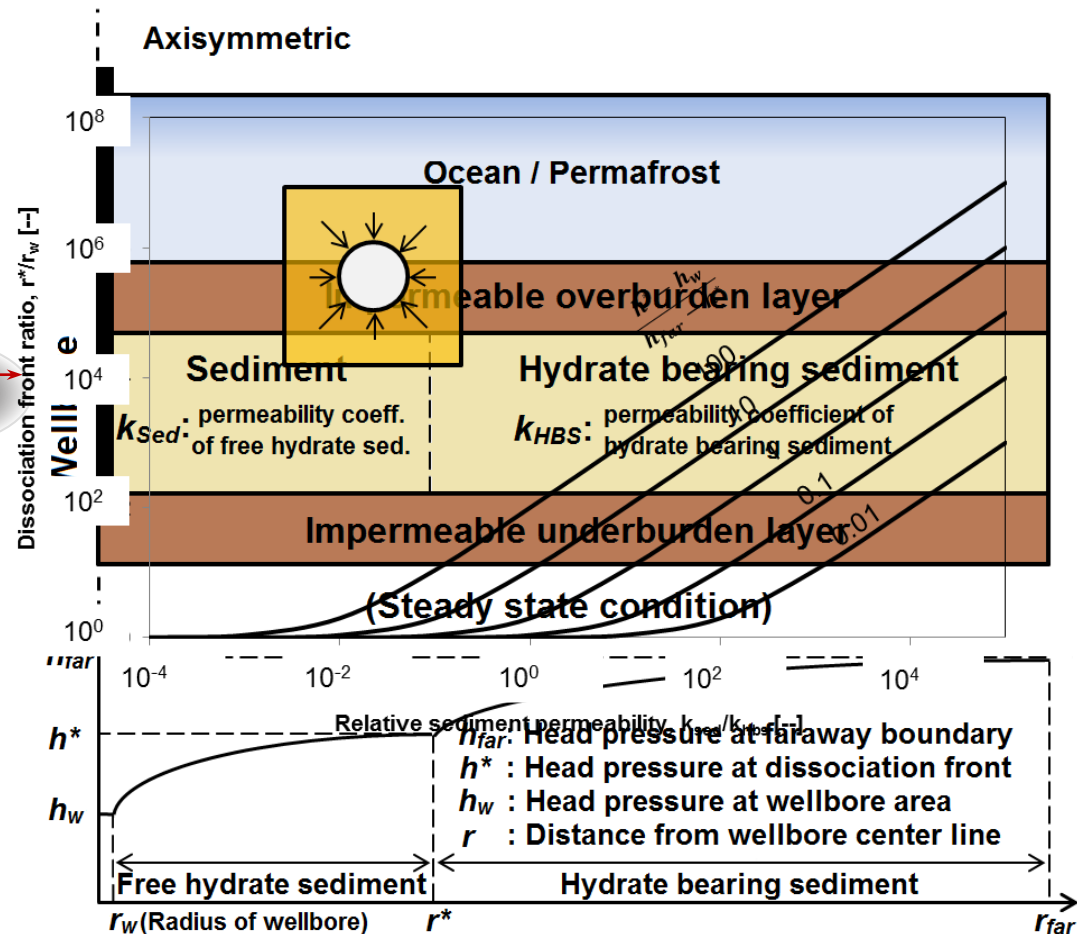
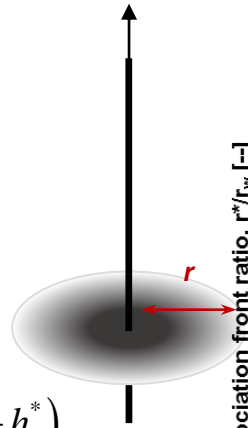
# Code verification

➤ Maximum gas production from HBS by depressurization

✓ Analytical solution – Cylindrical radial flow - Steady state conditions

$$q = -\frac{2\pi kH(h_2 - h_1)}{\ln\left(\frac{r_2}{r_1}\right)}$$

$$\frac{2\pi k_{Sed} H (h^* - h_w)}{\ln\left(\frac{r^*}{r_w}\right)} = \frac{2\pi k_{HBS} H (h_{far} - h^*)}{\ln\left(\frac{r_{far}}{r^*}\right)}$$



$$r^* = \left( r_w r_{far} \left( \frac{k_{Sed}}{k_{HBS}} \right) \left( \frac{h^* - h_w}{h_{far} - h^*} \right) \right)^{-1} \left( 1 + \left( \frac{k_{Sed}}{k_{HBS}} \right) \left( \frac{h^* - h_w}{h_{far} - h^*} \right) \right)$$

# Code verification

## ➤ Maximum gas production from HBS by depressurization

### ✓ Analytical solution – Cylindrical radial flow - Steady state conditions

- 2D axisymmetric model
- A single vertical well producing
- Very fine grid (2503 elements)

$$k_{HBS} = k_{sed} (1 - S_h)^N$$

$$k_{HBS} = 1 \times 10^{-12} \text{ m}^2$$

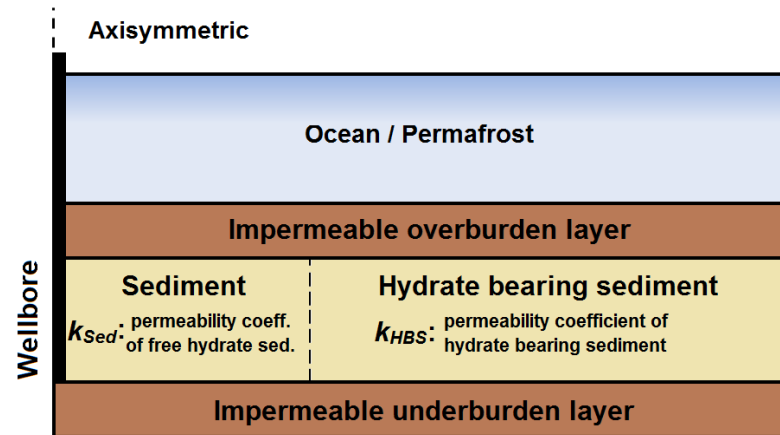
$$S_b = 0.5$$

$$\phi = 0.40$$

$$L = 1.20 \text{ km}$$

$$b = 0.40 \text{ m}$$

$$r_w = 0.1 \text{ m}$$

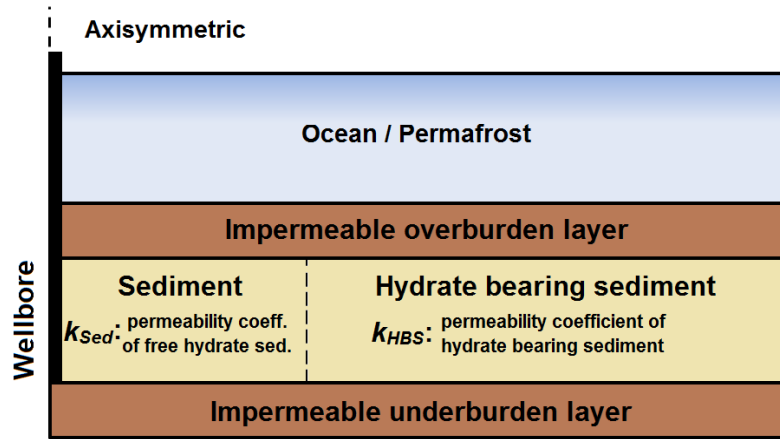


Case	$h_{far}$ (m)	$h_w$ (m)	$T$ (°C)	$\frac{h^* - h_w}{h_{far} - h^*}$
A	1020	306	12	7.14
B	1224	306	12	2.14
C	1224	510	12	1.44
D	1224	306	10	0.91

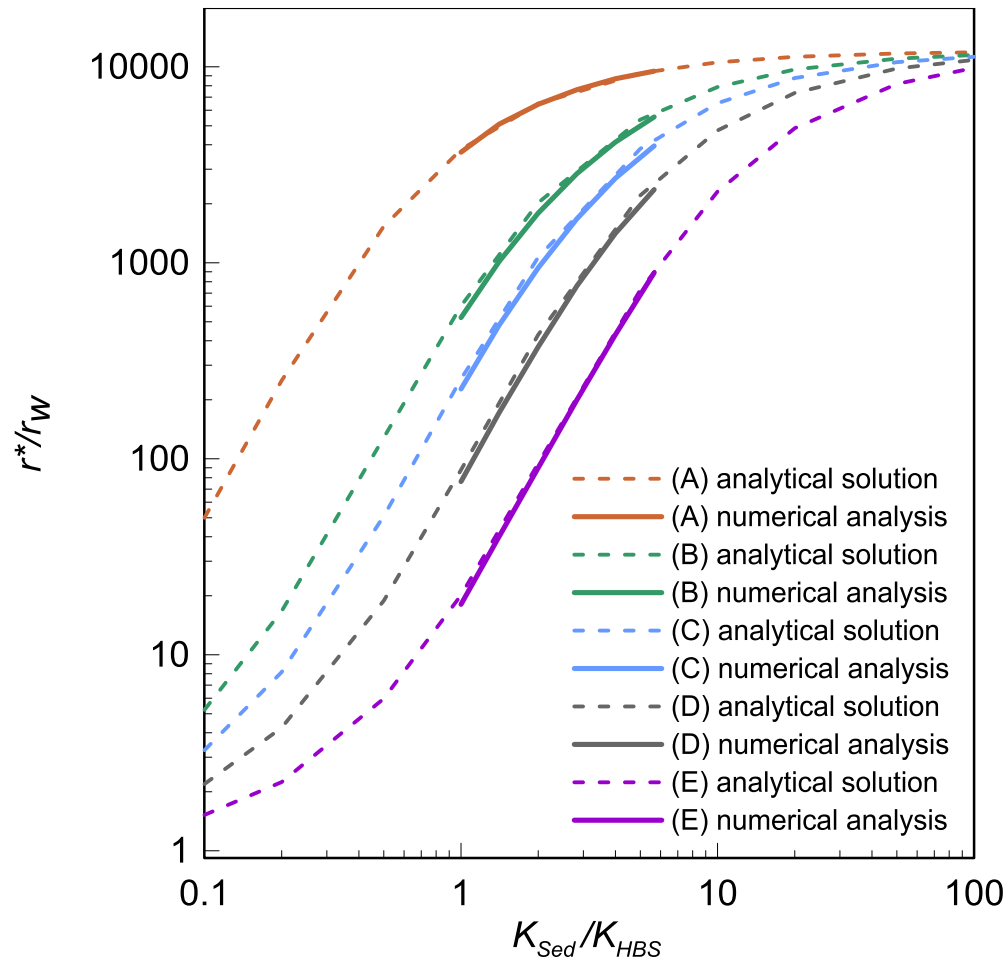
# Code verification

➤ Maximum gas production from HBS by depressurization

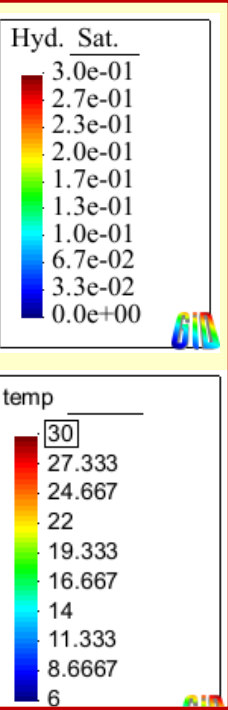
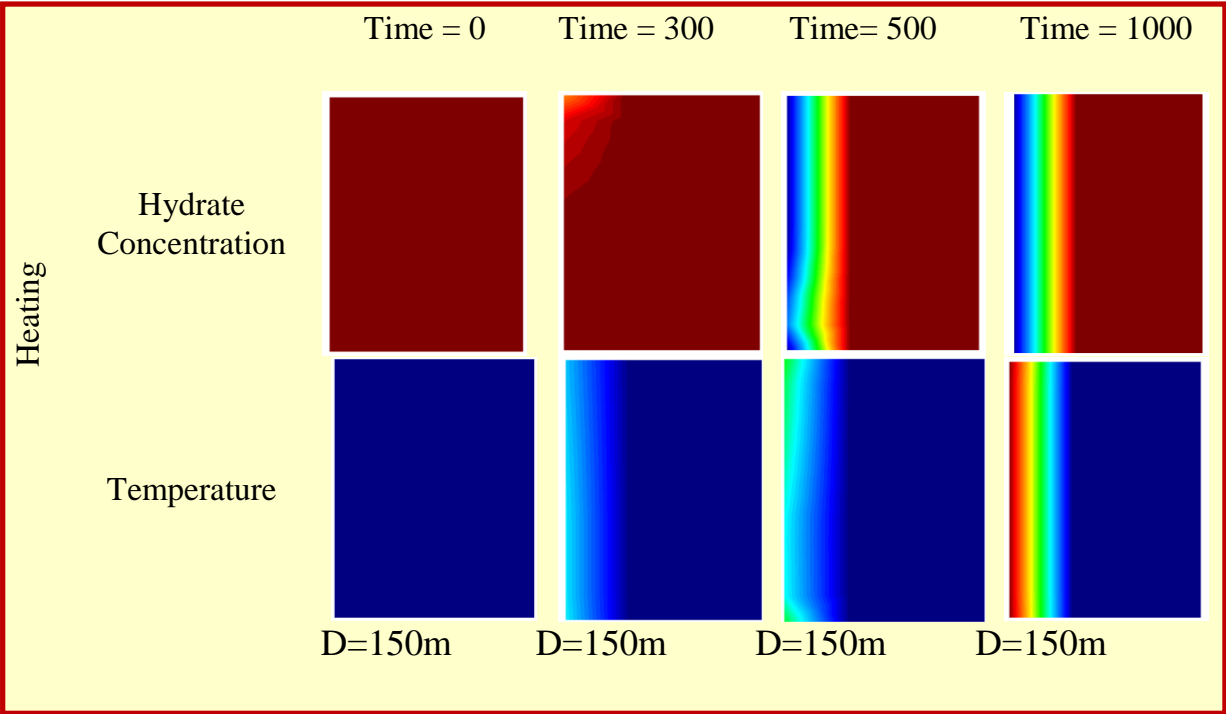
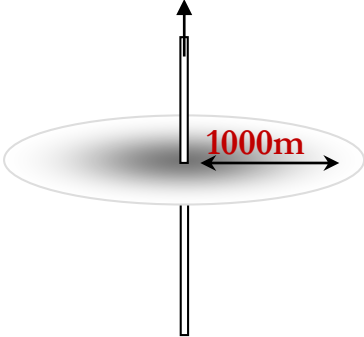
✓ Analytical solution – Cylindrical radial flow - Steady state conditions



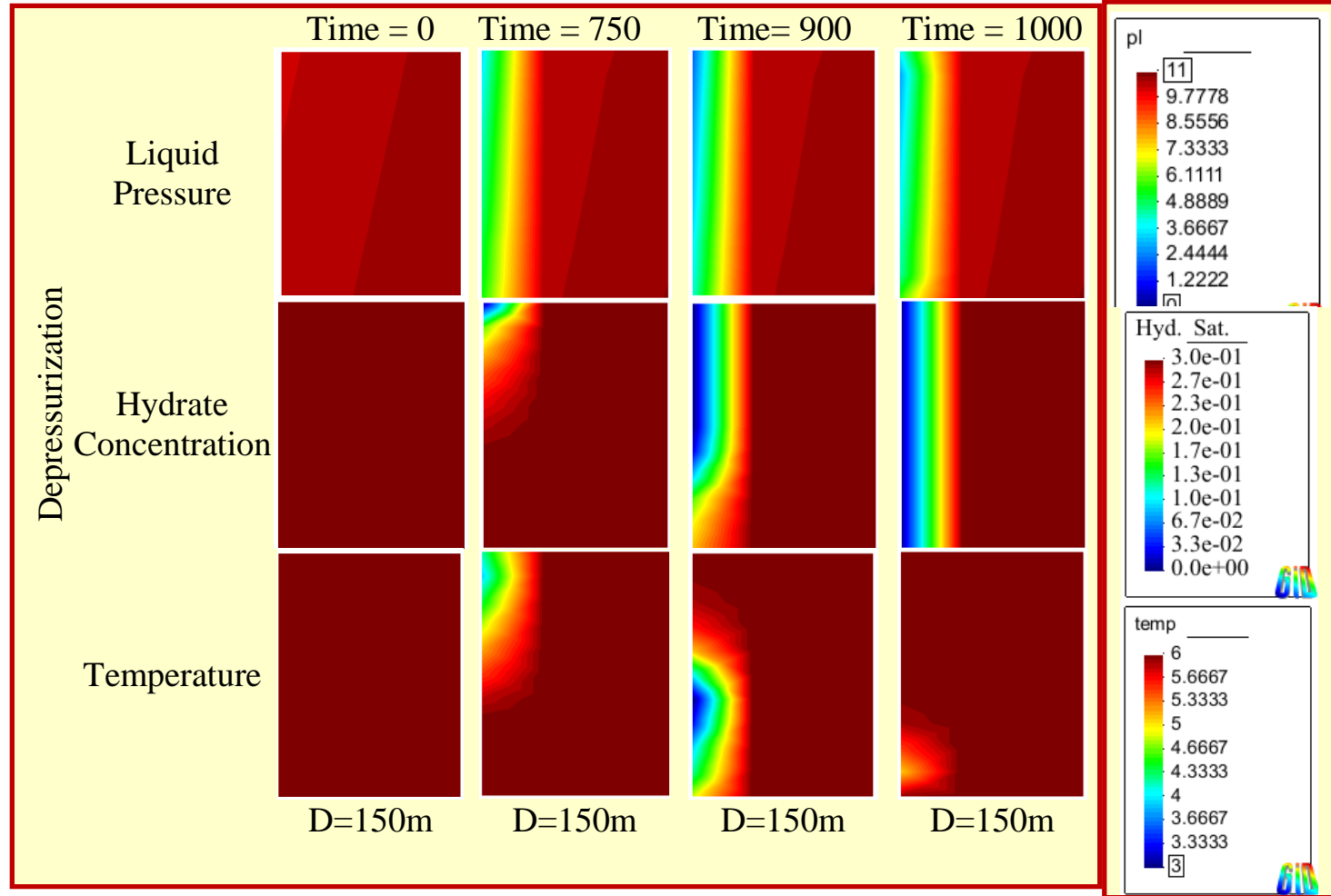
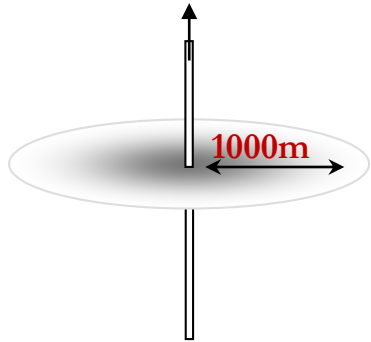
Case	$h_{gr}$ (m)	$h_w$ (m)	$T$ (°C)	$\frac{h^* - h_w}{h_{gr} - h^*}$
A	1020	306	12	7.14
B	1224	306	12	2.14
C	1224	510	12	1.44
D	1224	306	10	0.91



# Gas Production in-situ by Heating



# Gas Production in-situ by Depressurization



# Summary and conclusions

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- In this work we present a **coupled THCM formulation** for modeling the behavior of gas hydrates bearing sediments.
- The proposed approach incorporates the **fundamental physical and chemical phenomena** that control the behavior of gas hydrates bearing sediments.
- The **FE program CODE\_BRIGHT** has been adapted to incorporate the main balance and constitutive equations related to problems involving gas hydrate sediments.
- An **advanced mechanical model** for HBS has been proposed and **validated**.
- **Cases studies**, at actual scale, modeling the different strategies for gas (methane) production has been analyzed, showing the **potential of the proposed approach** to model these kinds of problems

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**Award No.: DE-FE0013889.**

